

**Outils d'interprétation des aquifères non-uniformes
en essai de pompage :
des "diagnostic plots" aux séquences de dimensions
d'écoulement**

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Séminaire RQES, 12 décembre 2019



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Review papers

Using flow dimension sequences to interpret non-uniform aquifers with constant-rate pumping-tests: A review

Anouck Ferroud*, Silvain Rafini, Romain Chesnaux

Research Group R2Eau, Cen

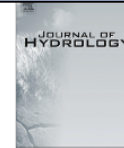


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Research papers

Insights on pumping well interpretation from flow dimension analysis: The learnings of a multi-context field database

Anouck Ferroud*, Romain Chesnaux, Silvain Rafini

Research Group R2Eau, Centre d'étu



HydrogeolJ (2017) 25:877–894
DOI 10.1007/s10040-017-1560-x



TECHNICAL NOTE

A numerical investigation of pumping-test responses from contiguous aquifers

Silvain Rafini¹ · Romain Chesnaux¹ · Anouck Ferroud¹

Hydrogeology Journal
<https://doi.org/10.1007/s10040-019-01972-7>



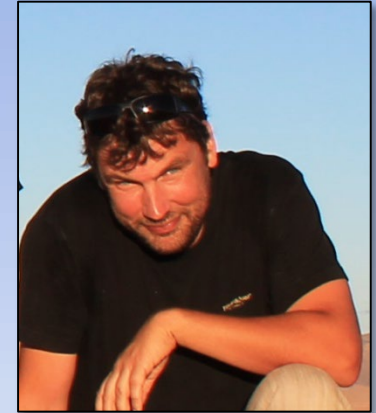
PAPER

Drawdown log-derived analysis for interpreting constant-rate pumping tests in inclined substratum aquifers

Anouck Ferroud¹ · Romain Chesnaux¹ · Silvain Rafini¹



Related publications



GW resources management : current challenges

- **Hydrogeology practitioners** must cope with new contexts:
 - **Sustainability management of GW resources systems that are increasingly pressured** (population growth and industrial development)
 - Emerging advanced fields: geothermal energy, radionuclide in situ repositioning, carbon sequestration etc.
 - Stricter legislations for water quality and environmental impacts assessment

- This requires to provide **refined investigation tools in routine applications**

- **Better assessing the complex nature of real aquifers**

- **Conceptual models accounting for heterogeneous flow conditions**

GW resources management : common practices

➤ Routinely used models are **overly simplified**

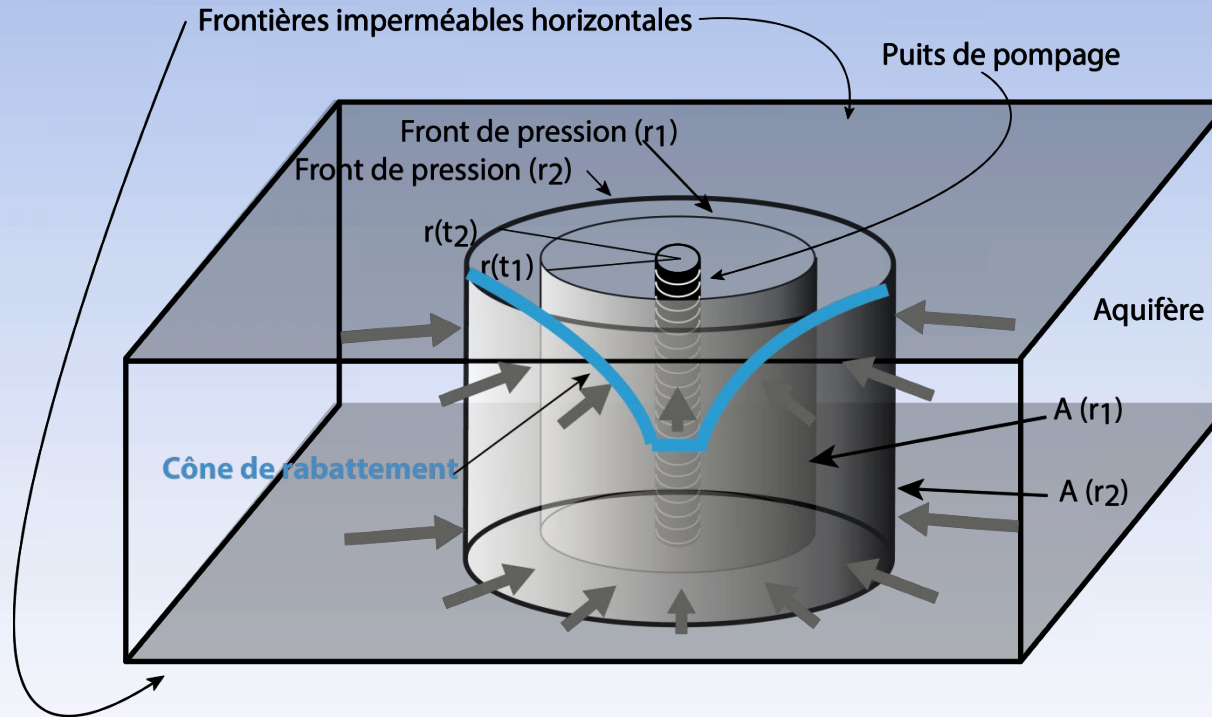
- Curve-matching with **Theis** type-curves (Theis, 1935)
- Theis-derived models: **Cooper and Jacob** semilog plot (Cooper and Jacob 1946)

➤ **Theis model (80 years-old...!)**

- First-order solution to the hyperbolic transient-diffusivity problem,
- Assumes **perfectly homogeneous and uniform domain**
- **Radially symmetric flow geometry** into an *Infinite Acting Radial Flow* (IARF) domain

GW resources management : common practices

Theis-like aquifer (also referred to as IARF model) :
radially symmetric flow geometry



Unable to render any heterogeneity of flow occurring into real aquifers
Produce gross assessments of the hydraulic properties
Overestimation and underestimation the hydraulic properties of specific hydraulic objects into the aquifer

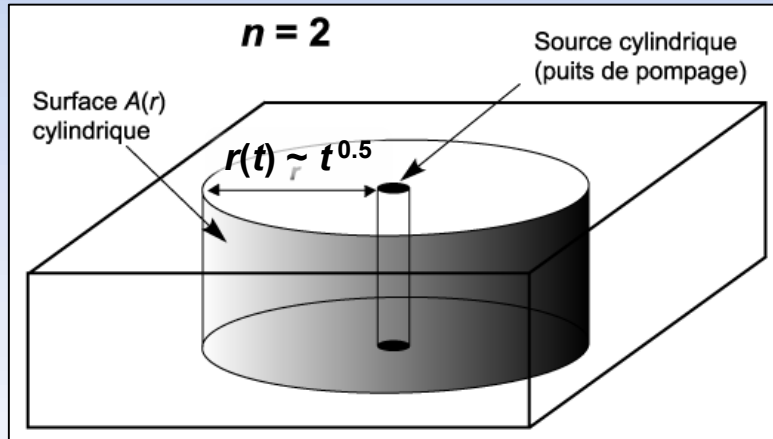
Radially symmetric flow geometry vs radial flow regime

Radially symmetric flow geometry (Theis-like) is a specific type of radial flow regime
Radial flow regime refers to the **transient growth of the cross-flow area $A(r)$**

Theis = IARF

Cylindrical shapes, radially symmetric flow
(homogeneous isotropic aquifer)

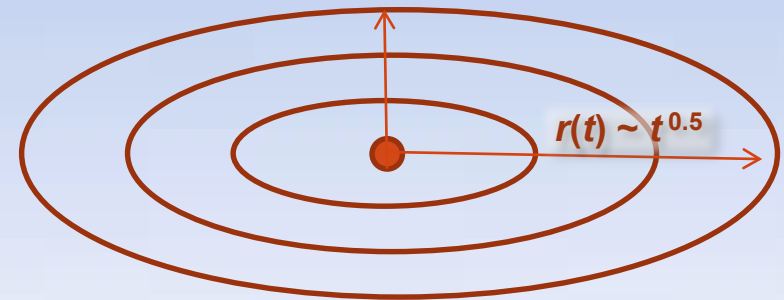
$$A(r) \sim r \rightarrow \text{Radial}$$



Non-Theis

Elliptical
(homogeneous anisotropic aquifer)

$$A(r) \sim r \rightarrow \text{Radial}$$



Any shape (heterogeneous aquifer)

$$A(r) \sim r \rightarrow \text{Radial}$$



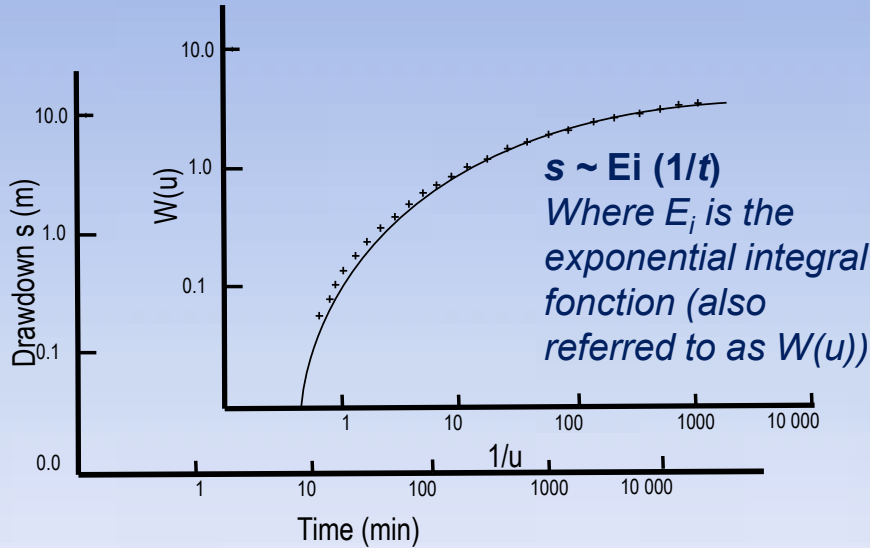
Radial flow strictly means $A(r) \sim r$

It does not refer to any specific symmetry of flow lines, a priori

The diagnostic response of a radial flow regime

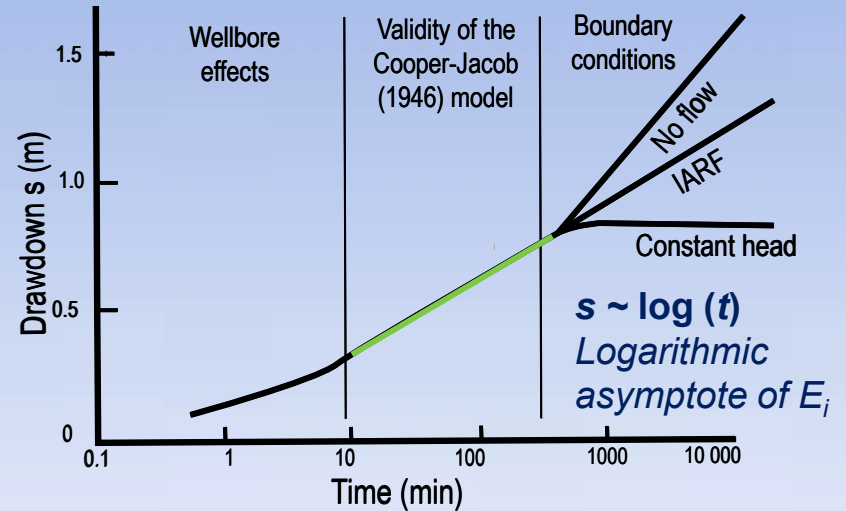
Theis (1935)

Curve-fitting on log-log time-drawdown plots



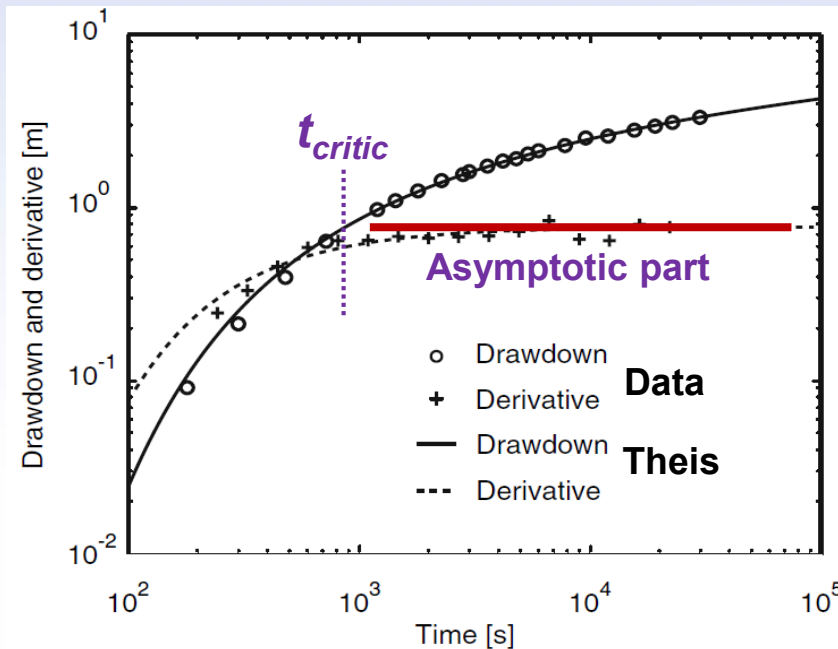
Cooper-Jacob (1946)*

Straight line on semi-log time-drawdown plots



* for large t or small $r \rightarrow$ at the source, practically from pumping test's beginning

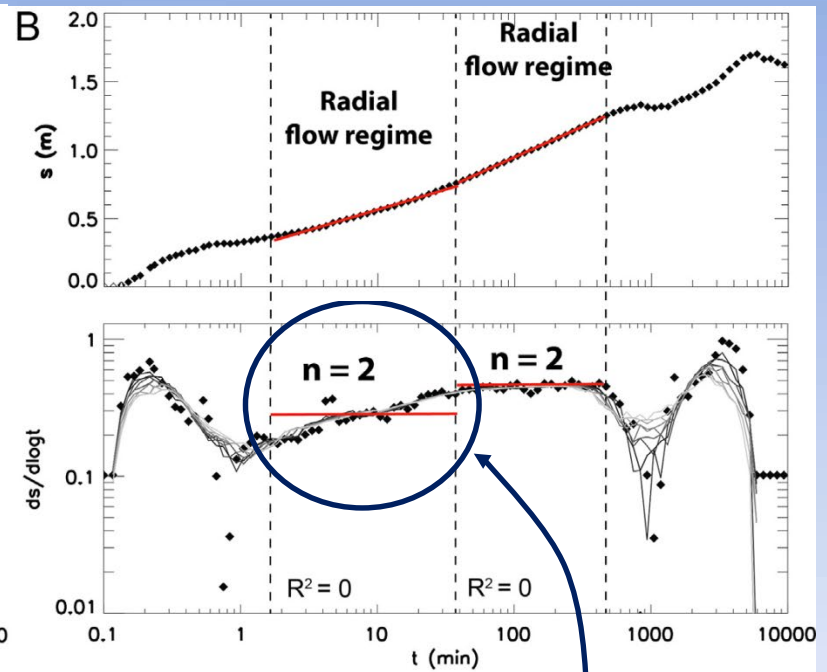
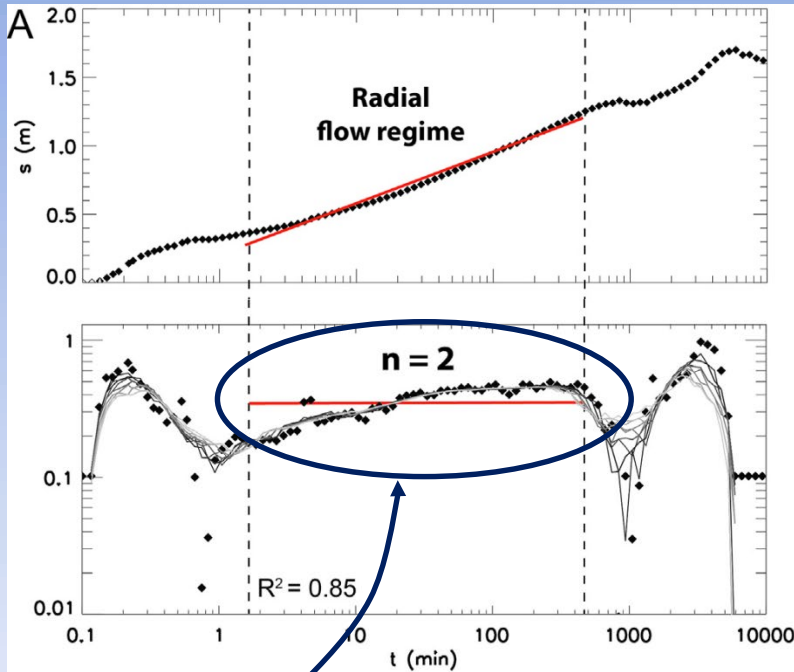
Drawdown log-derivative $ds/d\log(t)$



Plateau elevation = $2.3Q/4\pi T$

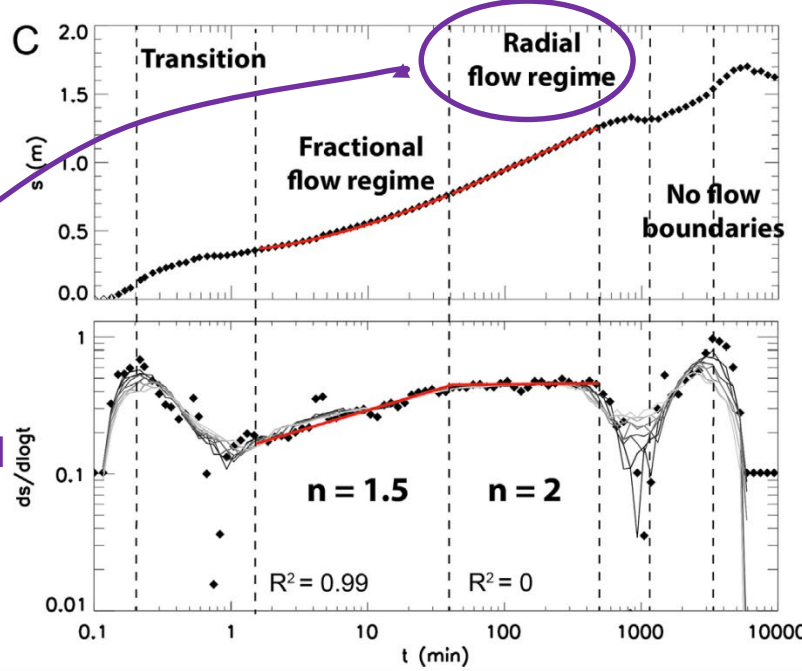
The derivative signature of a radial flow regime is a horizontal plateau (slope = 0)

Verifying the validity of the Theis model with the log-derivative signal



Mis-use of the Theis model
The multistage hydrodynamic response is grossly approximated

This can be used only into this time-window
The assumption of a radial flow regime is verified



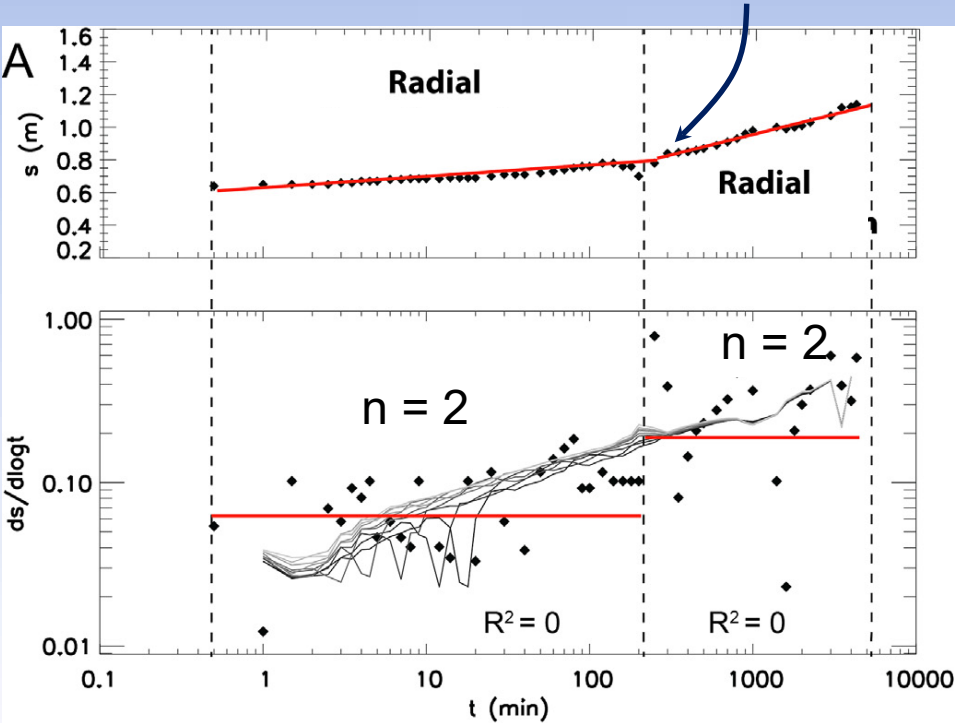
Mis-use of the Theis model
The assumption of a radial flow regime is not valid

Diagnostic non-radial flow regime ($n = 1.5$): the hydrodynamics of the aquifer is governed by a strongly inclined high-conductivity fault

Verifying the validity of the Theis model with the log-derivative signal

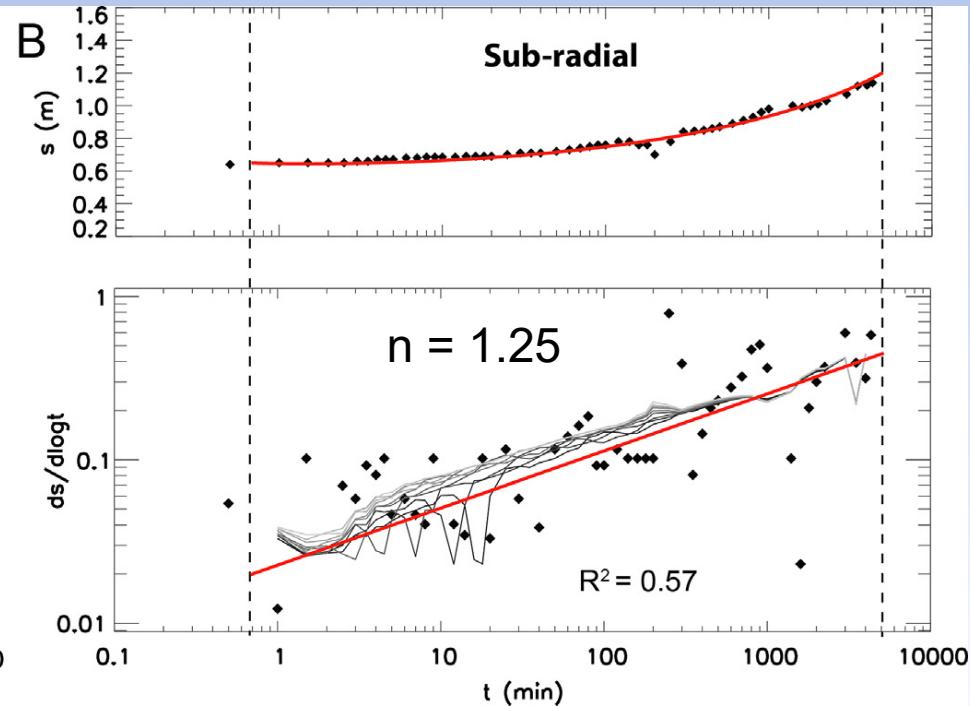
Interpretation with Theis-derived model

Impermeable boundary ?



Interpretation with fractional flow model

→ No !



Noisy derivative time serie → Bourdet differentiation

Mis-use of the Theis model

The assumption of a radial flow regime *is not* valid

GW resources management : validity of the common practices

Prior validation of a radial flow regime should be done in routine applications before applying Theis or CJ methods

Poorly assessed by straight lines in CJ semi-log plots → derivative plots

1. **How valid is the Theis model** in nature? To which degrees does it diverge from reality ?
2. **How significant** are the qualitative and quantitative **errors** induced on aquifers characterisation ? Practical implications ?
3. **How to assess the validity** of the Theis model in routine applications ?
4. **If not in a radial flow regime... what are the alternative models ?**

Some consequences of applying overly idealized interpretation model

- Simplifying the behaviour of the system to an extreme degree and disregarding the real geometry of flow;
- Ignoring the presence of several 1D, 2D or 3D hydraulic objects with non-equal properties, which may actually be governing the aquifer's global hydrodynamics at various pumping times;
- In low-conductivity contexts, overestimating by several order of magnitude the hydraulic properties of the pumped domain and missing the presence of distal and/or discrete conductive domains which may be exerting a predominant role in supplying water to the well over some pumping time-windows;
- Missing the presence of boundaries, or erroneously interpret non-existent ones;
- Globally, dismissing most of the diagnostic potential of the time-drawdown signal;
- Poorly assessing the impact of pumping an aquifer :
 - Erroneous sustainable pumping rates
 - Erroneous delineation of WHPA (wellhead protection areas)
 - Misunderstanding of the risk from potential contaminant source inventory,
 - Misunderstanding of the incidence on various objects into the environment, wet zones, etc.

FLOW REGIMES ANALYSIS

Real-world pumping tests databases

Compilations from various geological contexts : hard-rock (magmatic, sandstones, limestones), granular (unconsolidated sediments, fluvial channels)

- Rafini (2008) : compilation of **41 constant-rate pumping tests**
 - 80% exhibited multi-stage responses
 - **Radial regime occurs in 17% of the 41 datasets**

- Ferroud et al. (2018) : compilation of **69 constant-rate pumping tests**
 - 88% exhibited multi-stage responses
 - **Radial regime = 31% (of 121 interpreted flow regimes)**

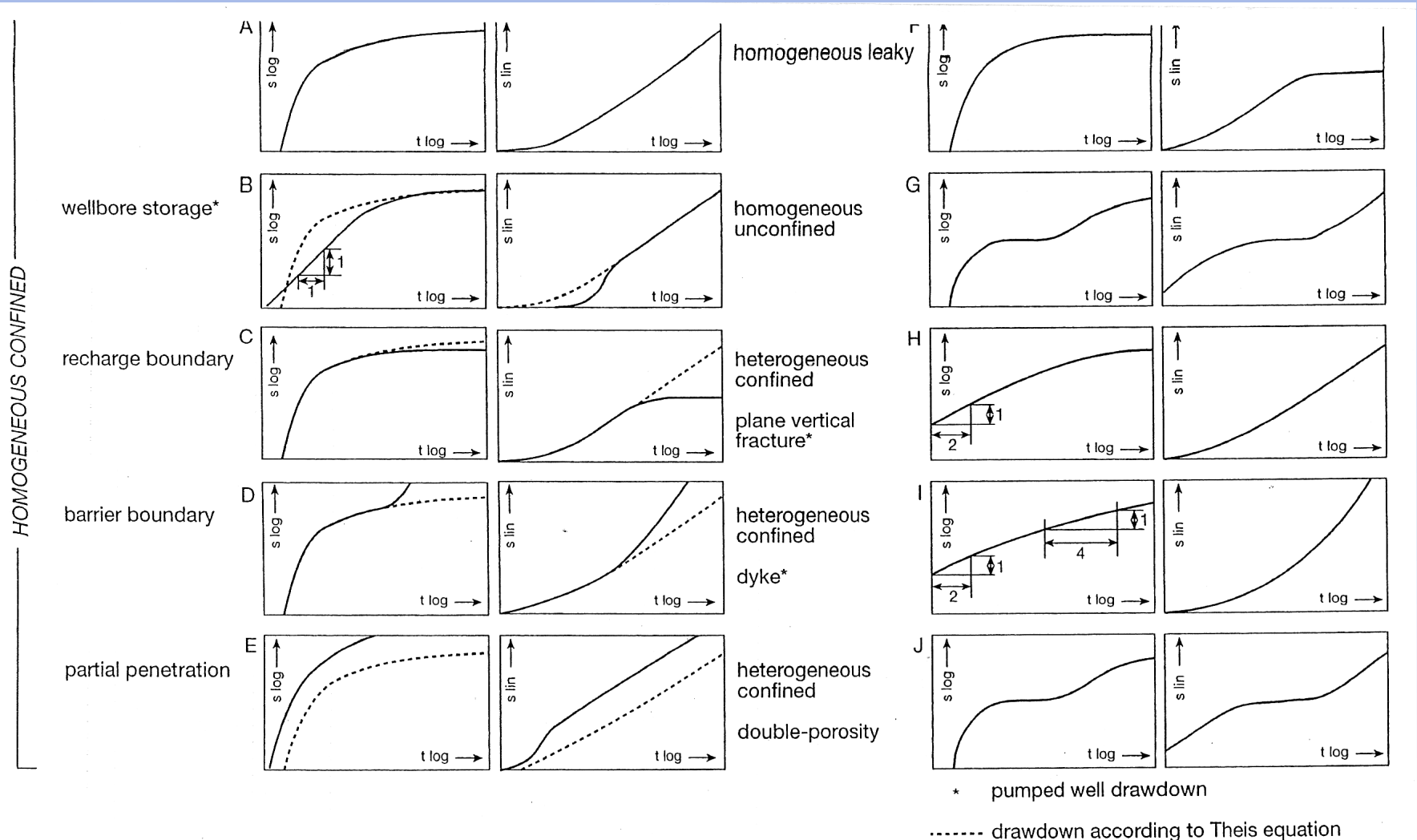
- **Since the early 80's**, numerous authors* have reported
 1. that the **flow regimes occurring in real media are actually much more complex and diversified** than the unique radial flow regime
 2. That **This model is unable to accurately reproduce the obtained responses** in many occurrences

- **70's, 80's, 90's** : Oil&Gas + GW researches produced **numerous analytical models** accounting for **heterogeneous flow** into various reservoir configurations

- **Provides numerous diagnostic diagrams : *diagnostic plots approach***

* (Audouin et al., 2008; Ferroud et al., 2018; Kuusela-Lahtinen et al., 2003; Leveinen, 2000; Lods and Gouze, 2004; Maréchal et al., 2004; Odling et al., 2013; Verbovšek, 2011, 2009 ; Bourdet et al, 1983)

Diagnostic plots approach



Two fundamental historic breakthrough developments

1. The *derivative analysis* (Bourdet et al, 1983)

Use of Pressure Derivative in Well-Test Interpretation

Dominique Bourdet, * SPE, **J.A. Ayoub**, SPE, and **Y.M. Pirard**, ** SPE, Flopetrol-Johnston Schlumberger

Summary. A well-test interpretation method based on the analysis of the time rate of pressure change and the actual pressure response is discussed. A differentiation algorithm is proposed, and several field examples illustrate how the method simplifies the analysis process, making interpretation of well tests easier and more accurate.

Introduction

The interpretation of pressure data recorded during a well test has been used for many years to evaluate reservoir characteristics. Static reservoir pressure, measured in shut-in wells, is used to predict

ogeneous formations reveals the good definition obtained with derivative plots, and the distinction between currently used interpretation models is clearly shown.

2. The flow dimension theory (Barker, 1988)

WATER RESOURCES RESEARCH, VOL. 24, NO. 10, PAGES 1796-1804, OCTOBER 1988

A Generalized Radial Flow Model for Hydraulic Tests in Fractured Rock

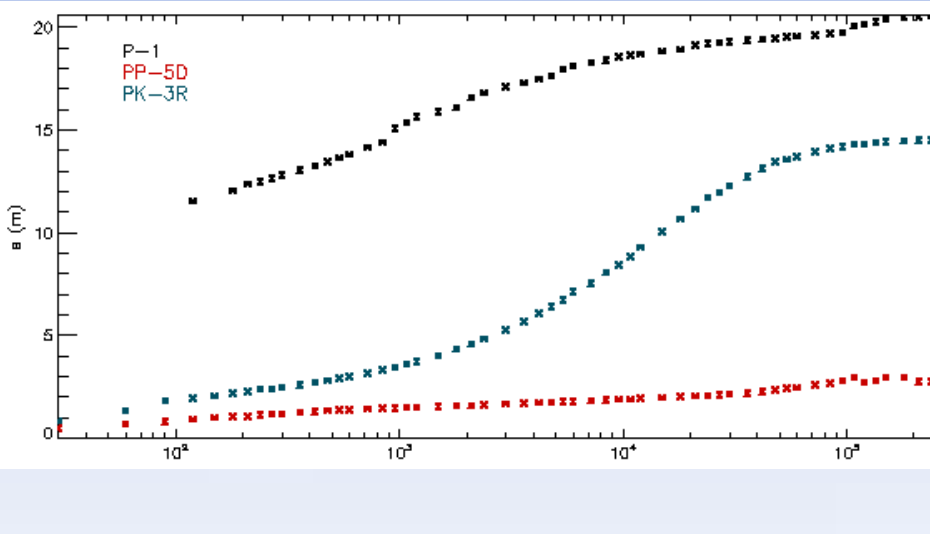
J. A. BARKER

British Geological Survey, Wallingford, Oxfordshire, United Kingdom

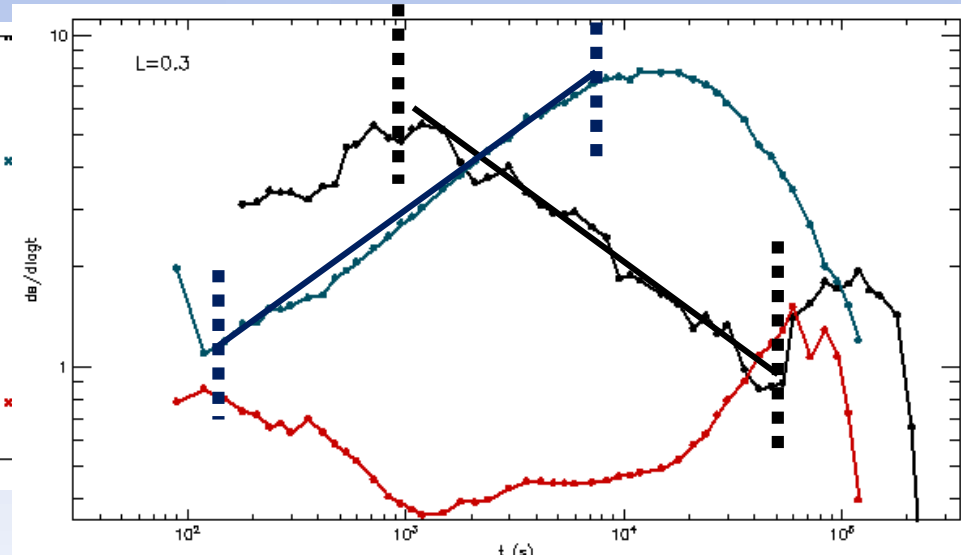
Models commonly used for the analysis of hydraulic test data are generalized by regarding the

1st historic breakthrough development :
the derivative analysis (eg. Bourdet et al, 1983)

Drawdown semi-log plot



Log-derivative $ds/d\log t$ bilog plots



The derivative signal provides a drastic gain in sensitivity

Makes it possible to distinguish between changes in flow regime caused by subtle variations in aquifer conditions

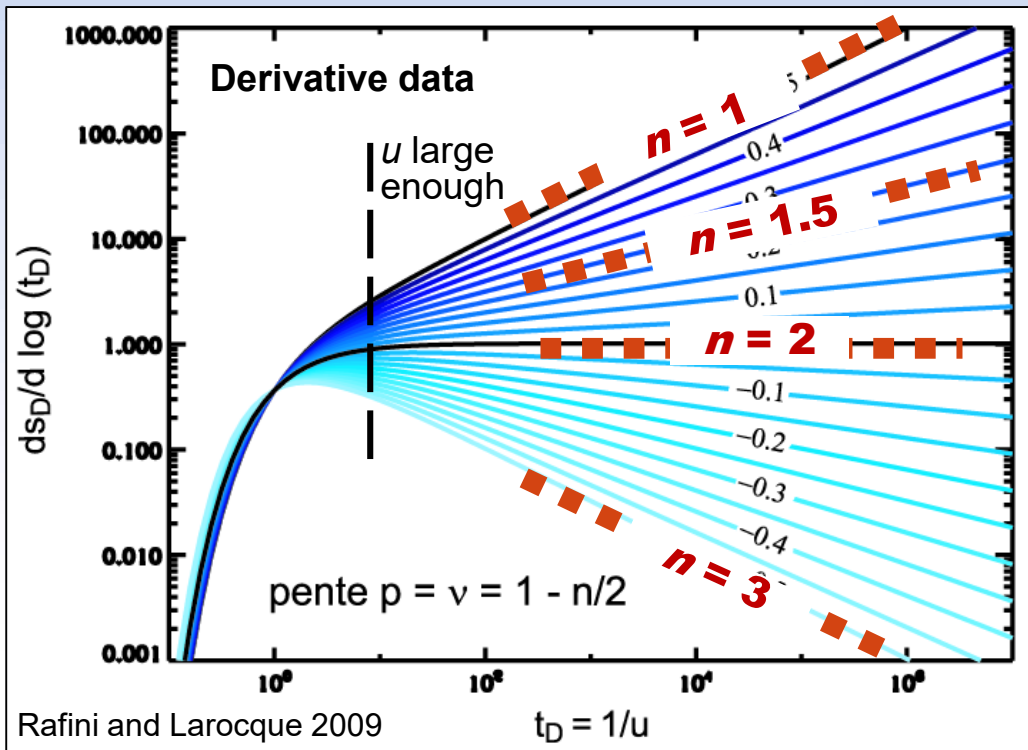
→ **identification of several successive flow regimes**

Constant slope of the derivative signal = **hydrodynamic stable flow regime**
(Barker, 1988)

2nd historic breakthrough development : the flow dimension theory (Barker, 1988)

The basics of a new formalism :

1. The flow regime is defined by a new parameter : the **flow dimension n**
2. Radial flow regime : $A(r) \sim r$; Generalized Radial Flow (GRF) regimes : $A(r) \sim r^{n-1}$
3. n reflects the **transient evolution of the frontal cross-flow area $A(r)$ at distance r**
4. n is obtained by a **direct reading*** of the log-derivative slope p : $n = 2 - 2(p)$
5. **Stable $n =$ hydrodynamically settled flow regime**



$n = 2$, radial flow regime =
drawdown log-rate is constant

$n < 2$ flow regimes
= drawdown log-rate is increasing
Aquifers with a limited potential

$n > 2$ flow regimes
= drawdown log-rate is decreasing
High potential aquifers

* for large u , i.e., large t or small $r \rightarrow$ at the source, practically from pumping test's beginning

2nd historic breakthrough development : the *flow dimension theory* (Barker, 1988)

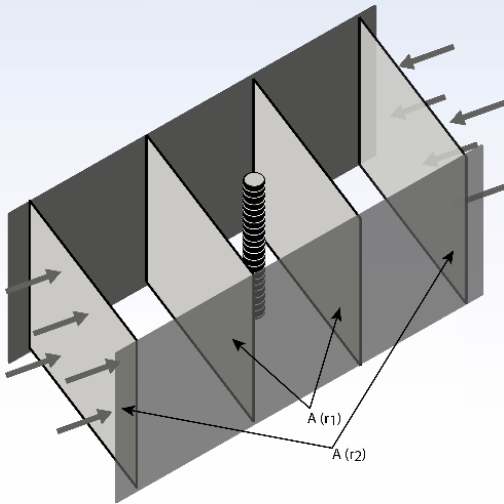
Barker's GRF model and flow dimension theory provide with a **universal relationship between** :

1. **The drawdown rate**, which is given by the log-derivative signal : $n = 2 - 2p$
2. **The expansion rate of the frontal cross-flow area $A(r)$** (depressurization front pulse), which is unknown and relates to conceptual models : $A(r) \sim r^{(n-1)}$

Ideal geometrical shapes

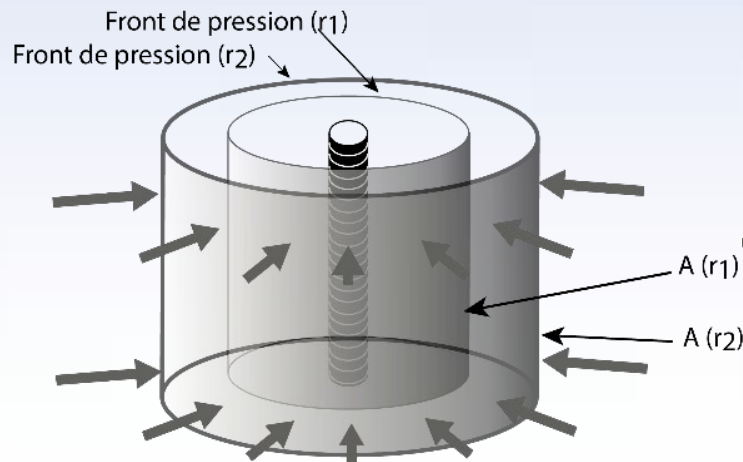
Linear regime: corridor

$$n = 1 ; A = \text{constant}$$



Radial regime: cylinders

$$n = 2 ; A \sim 2\pi r$$



Spherical regime: spheres

$$n = 3 ; A \sim 4\pi r^2$$

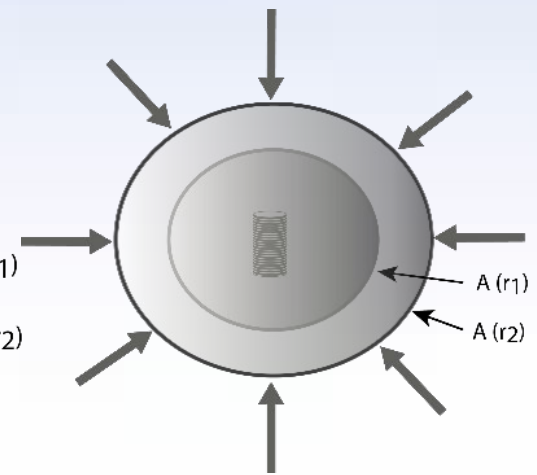
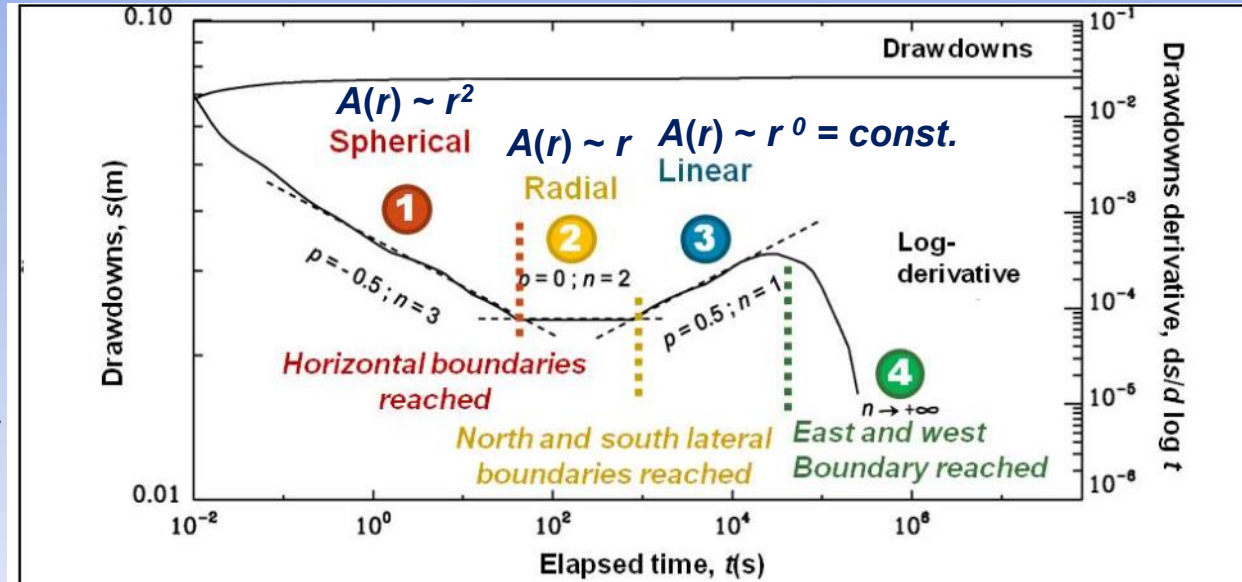


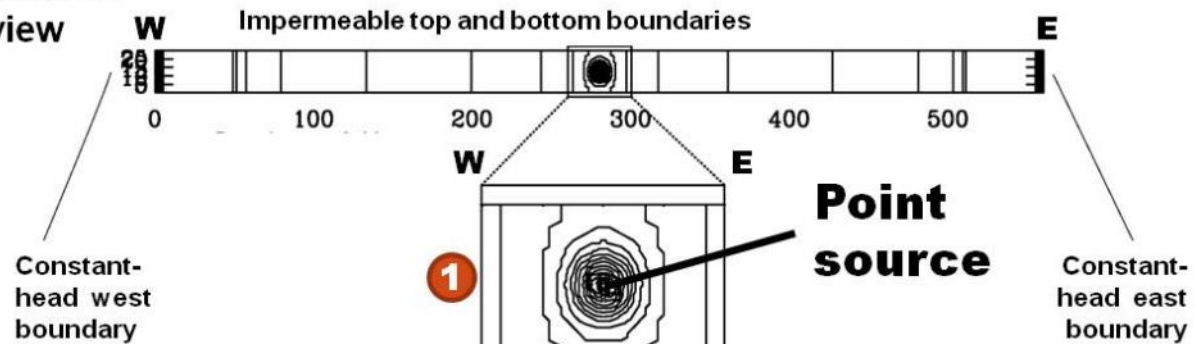
Illustration of the relationships between the propagation of the cross flow frontal area A and the flow dimension n (and the drawdown rate)

$$A(r) \sim r^{(n-1)}$$

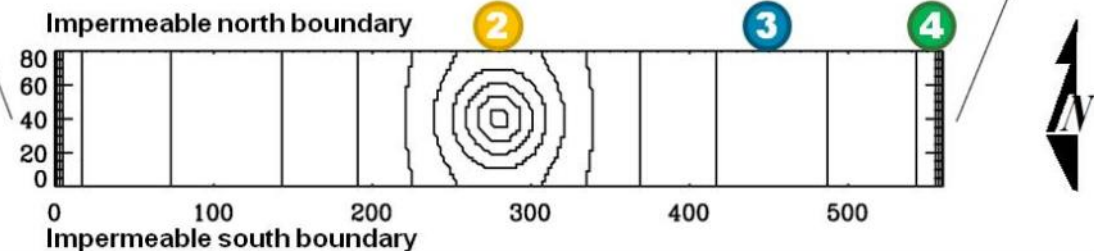
Numerical simulation of a point source into an elongated aquifer (homogeneous isotropic medium)



Section view



Map view

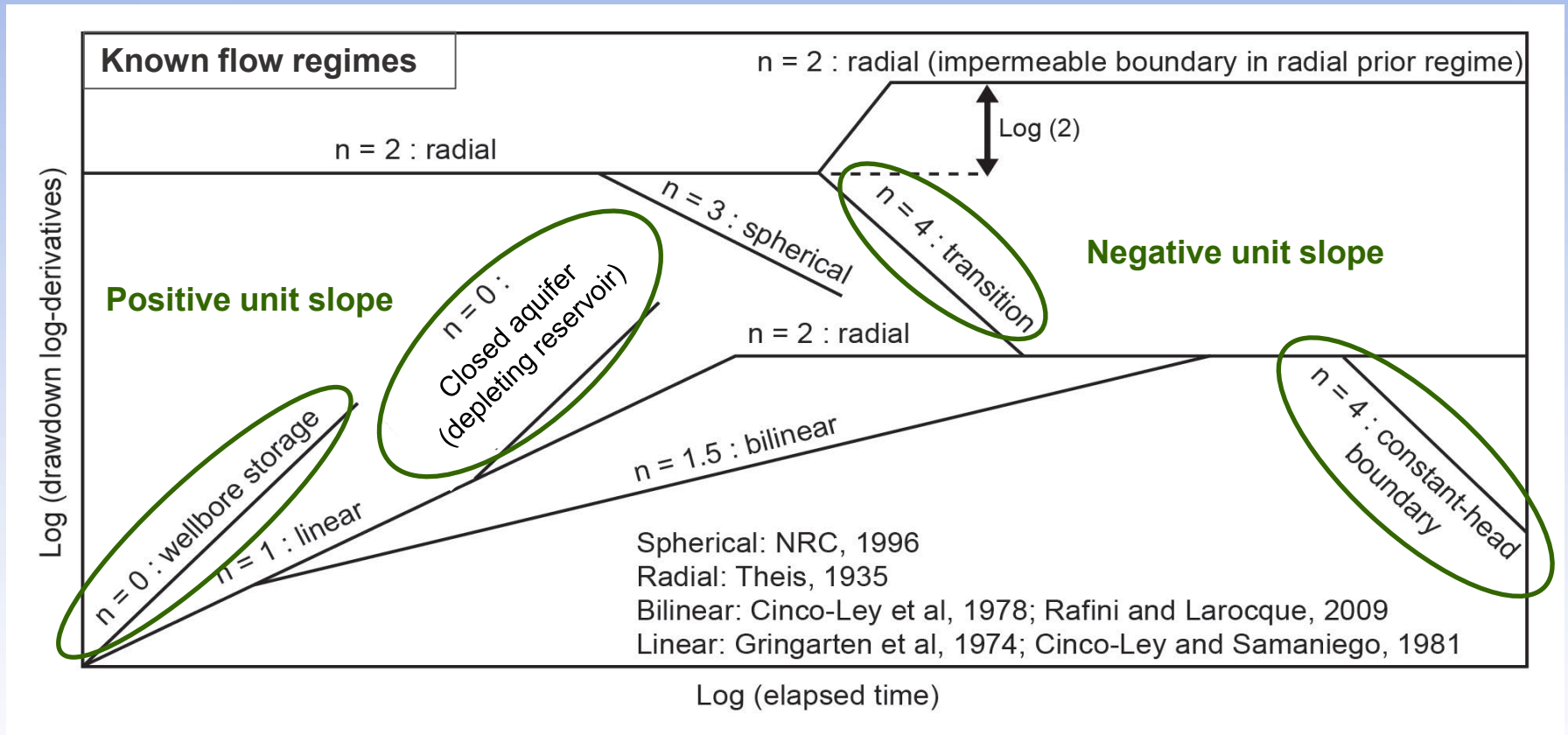


Evolution of n during the pumping test : scan of hydraulic conditions in the aquifer

n sequence : 3 – 2 – 1 – inf (spherical – radial – linear – inf)

The interpretable elementary flow regimes

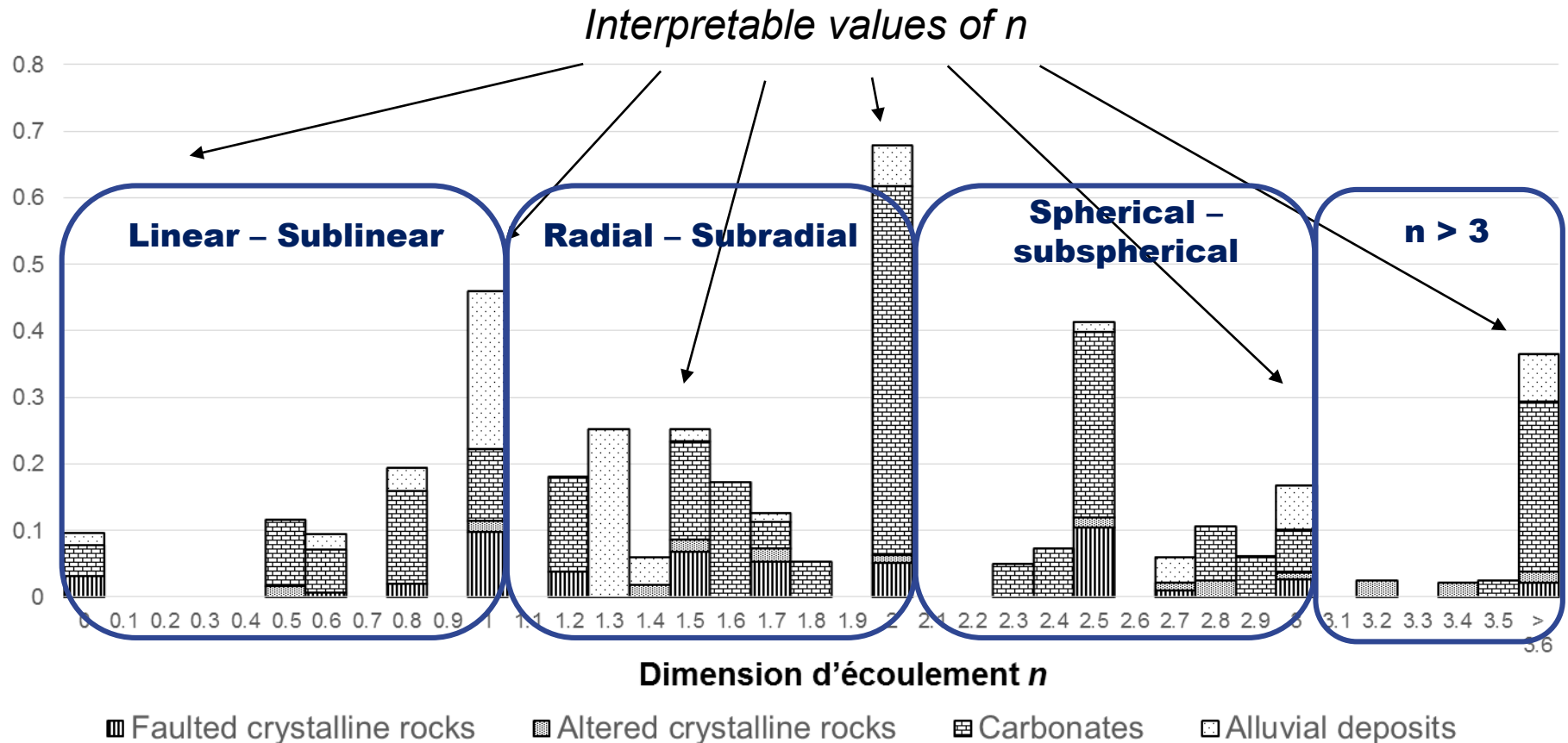
- Several types of flow regimes are recognized, in theory and in nature



- $n = 0$ and 4 : positive or negative unit slopes
- $n = 1, 1.5, 2, 3$: reflects specific hydrodynamic conditions → conceptual models
- Other values of n remain non-interpretable since no consensual conceptual flow model is available

The elementary flow regimes reported in nature

Ferroud ea (2017) database of 69 pumping datasets (121 distinct flow regimes)

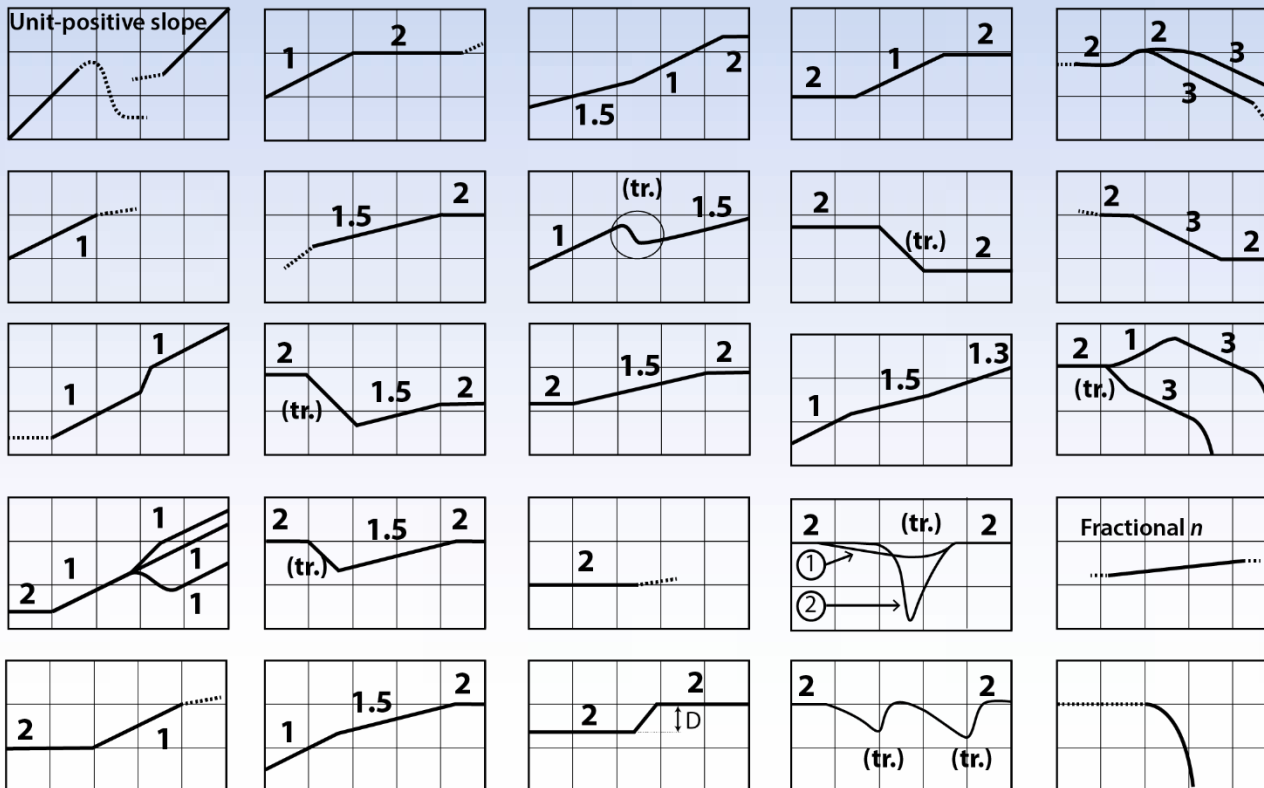


Reminder : $n \neq 2 \rightarrow$ Non-Theis aquifer

This is valid only to 30% of occurrences, and essentially in carbonate aquifers

Catalog of interpretable flow dimension sequences and associated conceptual models

- Comprehensive review of published conceptual flow models from petroleum and hydrogeology literature
- Mostly **analytical models** = various mathematical resolutions of the diffusivity problem with specific assumptions on the flow conditions (hydraulic and geometrical postulates)
- Also **numerical models** = empirical models obtained from experimental simulations, less idealized but criticized for its disputable generalization



Catalog of the interpretable flow dimension sequences, i.e. associated to a conceptual flow model

Flow dimension	Log-derivative response (bilog plots)	Conceptual model	Flow dimension	Log-derivative response (bilog plots)	Conceptual model
0		Early time: wellbore storage ⁽¹⁾ Late time: impermeable boundaries in the four directions ^(1, 2, 9) , referred to as <i>closed reservoir</i> and <i>pseudosteady state</i>	2 - 4		Recharge frontier
2		Theis aquifer (IARF), Cooper-Jacob approximation valid ⁽³⁾	1		<ul style="list-style-type: none"> • 2D structures Early time: <u>submetric</u> fracture ⁽⁵⁾ Medium time: infinitely conductive fault ⁽⁶⁾; finite-conductivity fault with skin effect • 1D conduits (Glacio-)fluvial channel ⁽⁷⁾, elongated reservoir ⁽⁸⁾, fractures or faults intersection ⁽¹⁷⁾
2 - 2		<ul style="list-style-type: none"> • Impermeable vertical frontier $D = \log(2) \rightarrow$ single frontier $D > \log(2) \rightarrow$ dual frontiers ⁽⁴⁾ • Aquifer cross-cut by a weakly-inclined finite-conductivity fault ⁽¹³⁾ • Leaky aquifer with depressurization of the non-pumped aquifer ⁽¹⁶⁾ 	1 - 1		Dual linear ⁽⁹⁾ : two opposite impermeable frontiers (elongated reservoir) reaching an impermeable boundary
2 - (4) - 2		($D < 1$) Presence of a lateral, blind (non-pumped), more <u>transmissive</u> flow domain ⁽¹⁵⁾	2 - 1		(Glacio-)fluvial channel, elongated reservoir
2 - ∞		Leaky aquifer with negligible drawdown into ^(1, 8)	1.5 - (2)		Aquifer cross-cut by a <u>strongly</u> inclined finite-conductivity fault ^(10, 11)
2 - 4 - 1.5 - 2		(Early and late plateau elevations are equal) Aquifer cross-cut by a <u>strongly</u> inclined finite-conductivity fault; the fault is <i>not connected</i> to the wellbore (not intercepted) ^(11, 14)	2 - 4 - 1.5 - 2		($D < 1$) Presence of a lateral, blind (non-pumped), more <u>transmissive</u> flow domain with a conductive fault at the interface ⁽¹⁵⁾
1 - 1.5 - (2)		Aquifer cross-cut by a strongly inclined finite-conductivity fault with occurrence of an early fault-related linear regime ⁽¹⁰⁾ possibly caused by fault's skin effects or internal layering	2 - 3		Aquifer with an inclined substratum or confining layer leading to variable thickness (wedge-like) ^(9, 17)
1.5 - 1 - (2)		Aquifer cross-cut by a strongly inclined finite-conductivity fault with occurrence of a matrix-related linear regime (fault acts as a planar source) ^(1, 10) ; barely realistic	3 - 2		Point source: Screen length shorter than the aquifer thickness ⁽¹⁷⁾ , partially penetrating well ⁽¹⁷⁾ , or pumping between packers
2 - 1.5 - (2)		Aquifer cross-cut by a moderately inclined finite-conductivity fault ⁽¹³⁾	n non-integer		4. Spherical combinations

1. Radial, dual radial

2. Linear, dual linear

3. Bilinear combinations

4. Spherical combinations

(1) Tiab, 2005;
(2) Linear no-flow frontier;
(3) Theis (1935), Cooper et Jacob (1949);
(4) Beauheim and Walker (1998);
(5) Cinco-Ley et al (1978)
(6) Gringarten et Ramey (1974, 1975);
(7) Massonat et al 1993;
(8) Miller (1962; Nutakki and Mattar 1982 ; Escobar et al, 2012; Escobar et al, 2007;
(9) Escobar et al (2004), Escobar and Montealegre (2007) ;
(10) Cinco-Ley et Samaniego (1981);
(11) Rafini et Larocque (2009);
(12) Rafini and Larocque (2012);
(13) Abbazsadeh et Cinco-Ley (1995);
(14) Rafini et al (accepted);
(15) Neuman et Witherspoon (1969);
(16) Ferroud et al (2016);
(17) Hantush (1956), Hanush (1960);
(18) Barker (1988).

Radial and dual radial combinations

$$n = 2$$

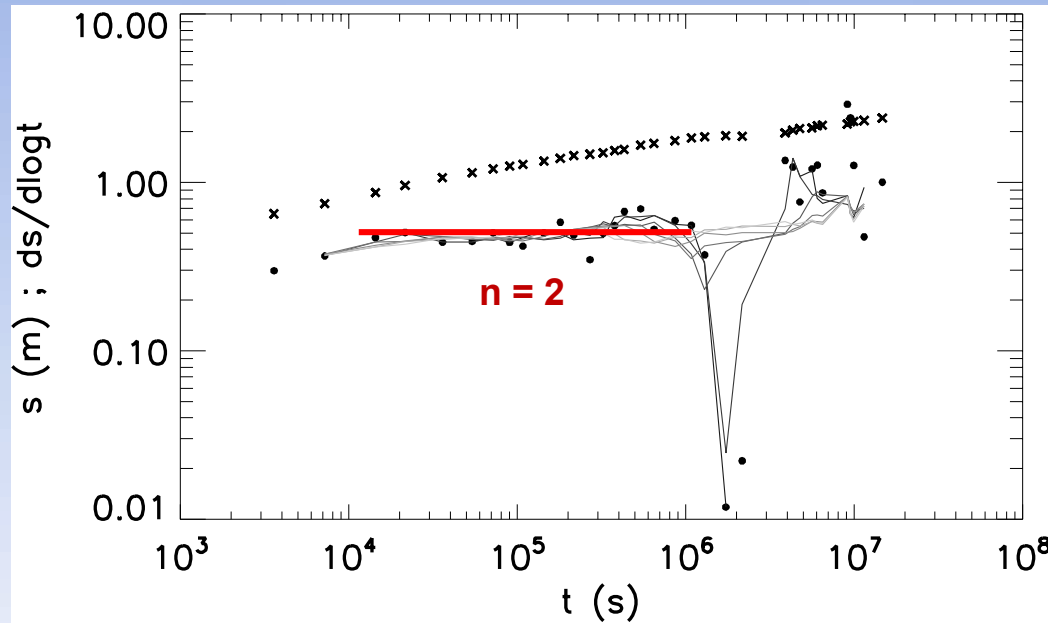
$$n = 2 - 2$$

$$n = 2 - 4 - 2$$

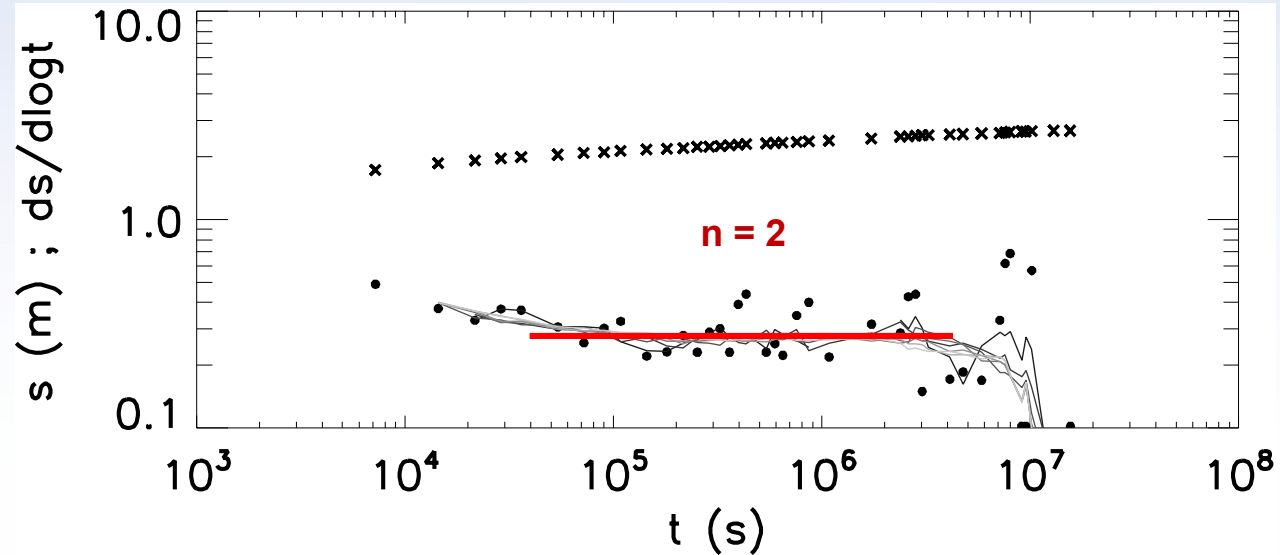
Radial flow regime

- The **most frequently observed** regime in natural media (ca. 20 to 30%)
- Occurs as often in fractured as in granular aquifers
- **3D hydraulic continuum** (e.g., sandstone, dense and conductive fracture network) **or...**
- **Weakly inclined conductive structure**
- In fractured aquifers, occurs **most frequently in carbonateous formations** due to the presence of conductive subhorizontal stratification planes
- Frequently the **last flow regime** (very late pumping time) regardless the aquifer conditions, due to heterogeneities being « diluted » or « averaged » into the high volume of depressurized aquifer → **bulk**, or **large-scale, hydraulic properties**

Radial flow regime : real examples



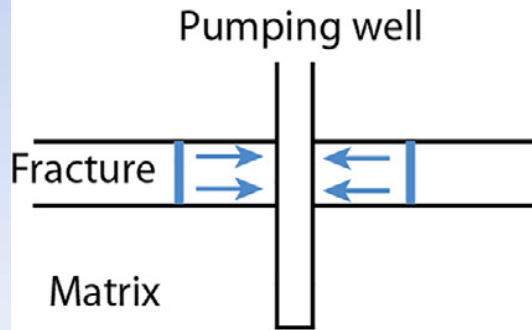
Carbonateous rocks aquifers



Radial flow regime : subhorizontal flow structures

- In stratified carbonateous rocks
- Hard rocks with weakly inclined conductive faults

a) Radial flow regime ($n = 2$) traversing through a horizontal fault

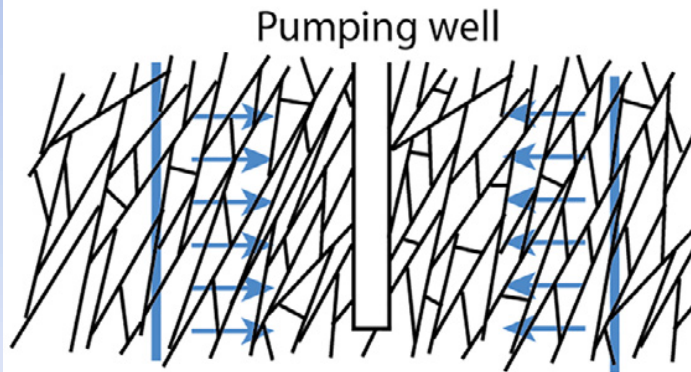


e.g., stratification in carbonateous rocks



Radial flow regime : dense fractured network (continuum-like)

b) Radial flow regime ($n = 2$) in a fractured network



e.g., fractured hard rock aquifer when the fractures network is **dense and connected**

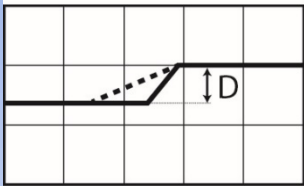
The cross-flow area is large compared to the size of the individual fractures

→ **The fractures network behaves like a continuum**



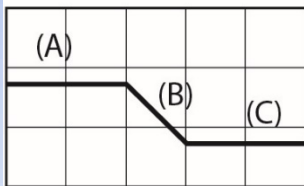
Dual radial sequences: overview

Positive offset



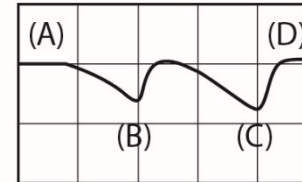
1. No-flow single ($D = 2$) or multiple ($D > 2$) boundaries
2. Weakly inclined conductive fault (low transition slope)

Negative offset

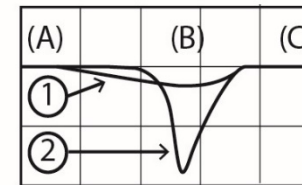


1. Cryptic (non-pumped) contiguous or distal aquifer (highly productive)
2. Aquifer with inclined substratum ($B : n = 3$)

No offset



Triple porosity (fractured rocks)



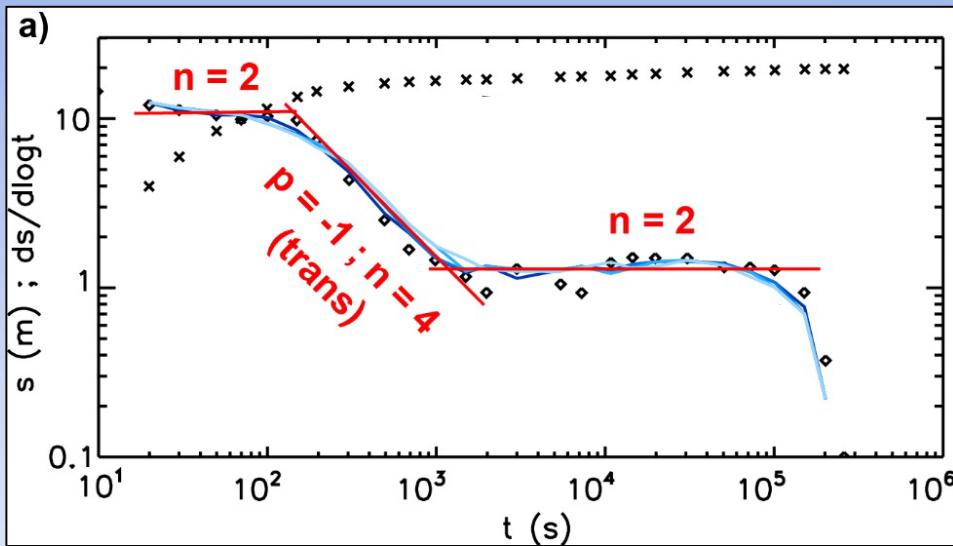
Dual porosity (fractured rocks)

1. Transient interporosity flow
2. Pseudo-steady interporosity flow

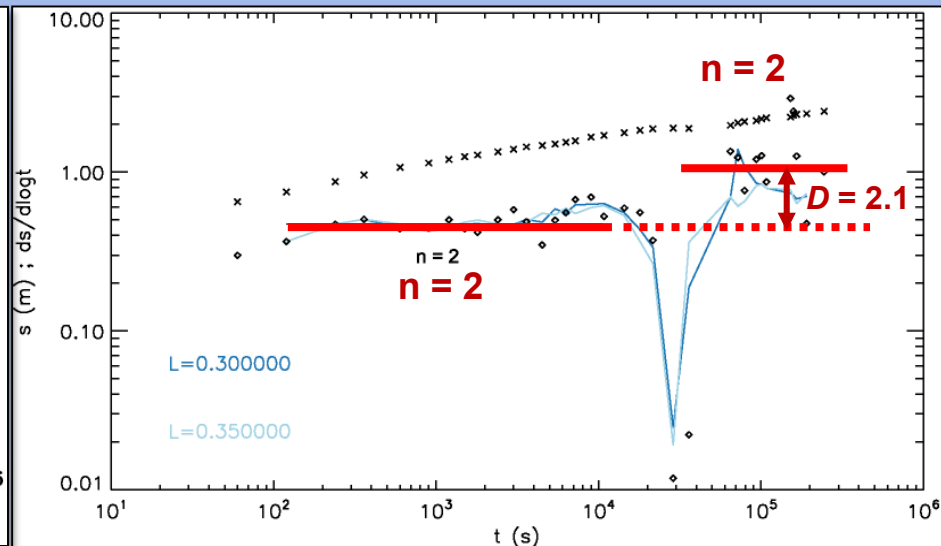
- Indicates the coexistence of several domains, either **juxtaposed** (frontier, contiguous aquifer) or **superposed** (multiple porosity)
- Critical features for interpreting the proper conceptual flow model are :
 - 1) whether the offset is null, negative or positive,
 - 2) the magnitude of the offset (whether it is greater or lower than 2)
 - 3) the shape of the transitional regime

Dual radial sequences: real examples

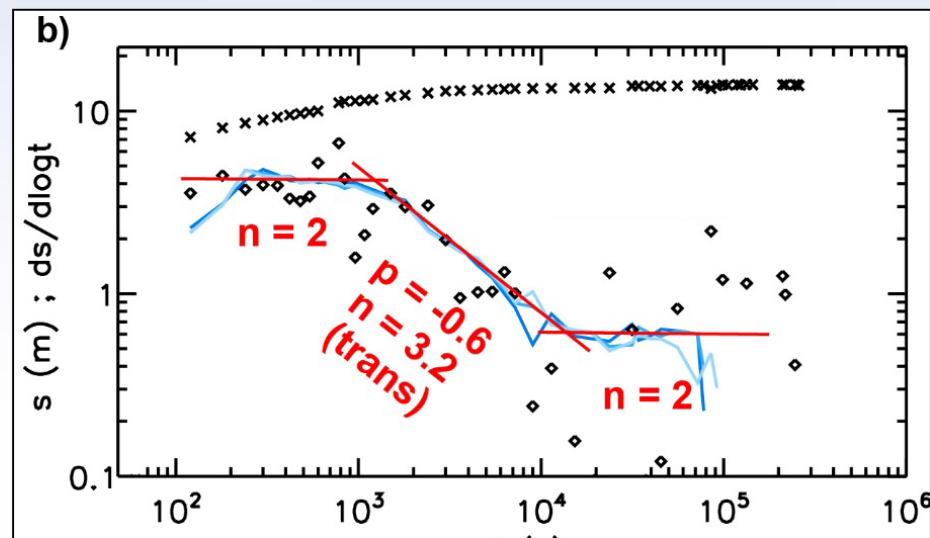
Negative offset



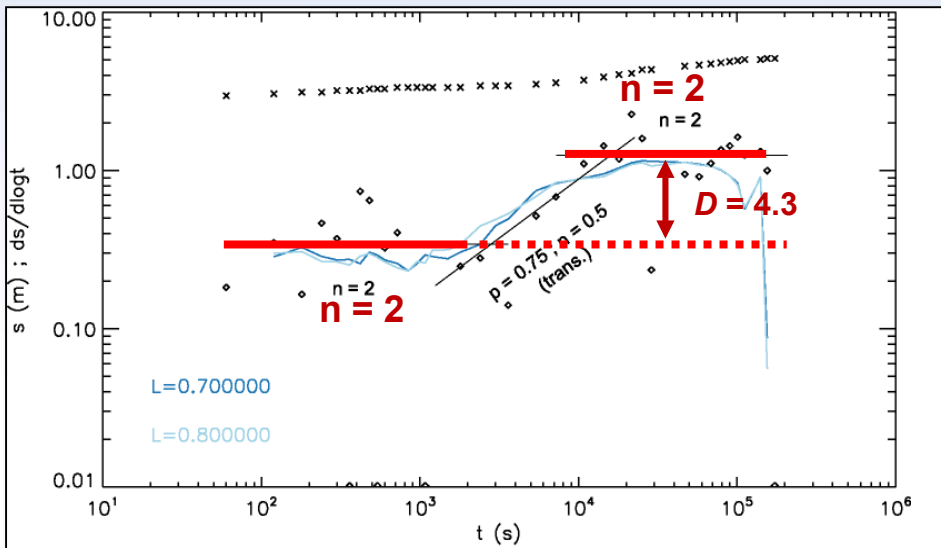
Positive offset ; $D \approx 2$



Negative offset

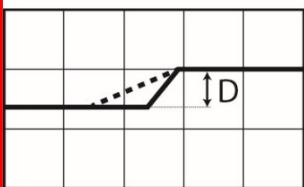


Positive offset ; $D > 2$



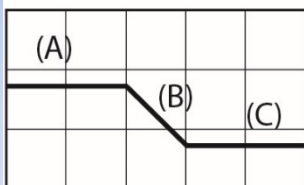
Dual radial sequences: impermeable frontiers

Positive offset



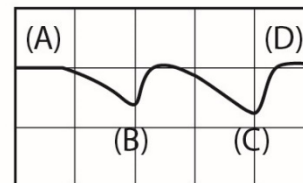
1. No-flow single ($D = 2$) or multiple ($D > 2$) boundaries
2. Weakly inclined conductive fault (low transition slope)

Negative offset

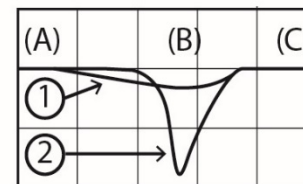


1. Cryptic (non-pumped) contiguous or distal aquifer (highly productive)
2. Aquifer with inclined substratum (B : $n = 3$)

No offset



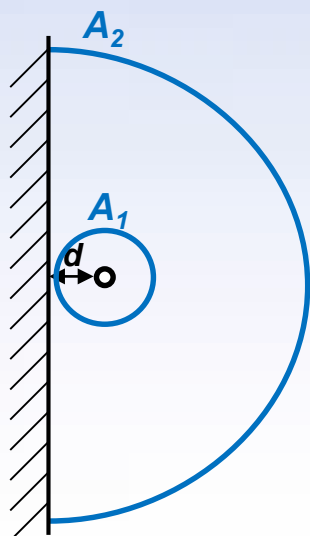
Triple porosity (fractured rocks)



Dual porosity (fractured rocks)

1. Transient interporosity flow
2. Pseudo-steady interporosity flow

CONCEPTUAL MODEL: single impermeable frontier



$$r < d : A_1(r) = 2\pi r b$$

$$r > d : A_2(r) = \frac{1}{2} 2\pi r b = \frac{1}{2} A_1(r)$$

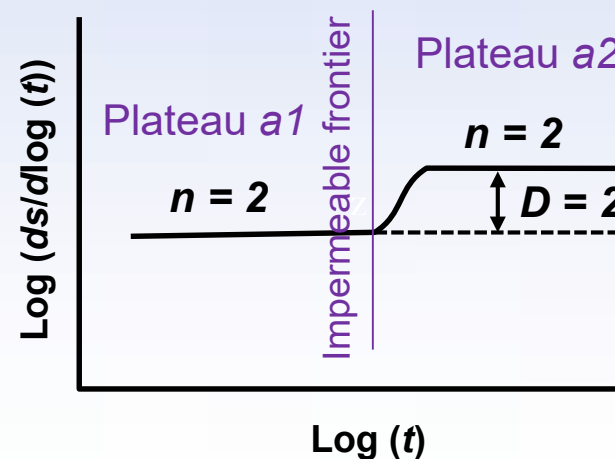
The transmissive surface is halved
 → Drawdown
 log-rate is doubled

$$p = a = 2.3Q/4\pi T$$

$$a_2 = 2 a_1$$

$$(or p_2 = 2 p_1)$$

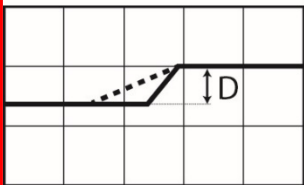
Log-derivative response
 Sequence of $n : 2 - 2$ with $D = 2$



The slope (inversely proportional to transmissivity) is doubled after the no-flow boundary is reached

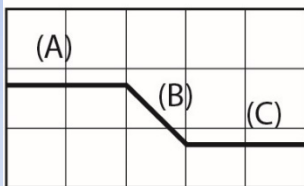
Dual radial sequences: multiple impermeable frontiers

Positive offset



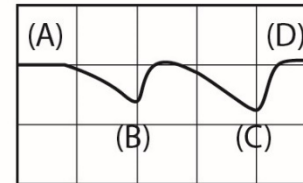
1. No-flow single ($D = 2$) or multiple ($D > 2$) boundaries
2. Weakly inclined conductive fault (low transition slope)

Negative offset

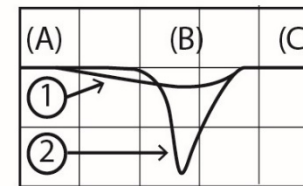


1. Cryptic (non-pumped) contiguous or distal aquifer (highly productive)
2. Aquifer with inclined substratum (B : $n = 3$)

No offset



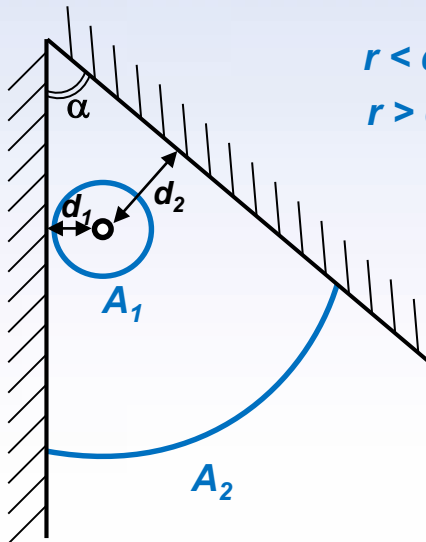
Triple porosity (fractured rocks)



Dual porosity (fractured rocks)

1. Transient interporosity flow
2. Pseudo-steady interporosity flow

CONCEPTUAL MODEL: generalized multiple impermeable frontiers model

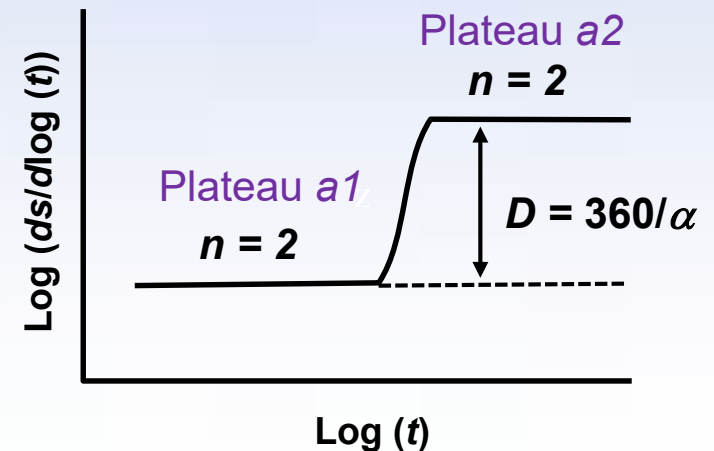


$$r < d_1 : A_1(r) = 2\pi r b$$

$$r > d_2 : A_2(r) = (\alpha/360) A_1(r)$$

The transmissive surface A_2 is decreased by a factor $\alpha/360$
 → Drawdown rate increases by an equal inverse factor

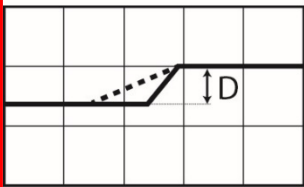
Log-derivative response
 Sequence of $n : 2 - 2$ with $D > 2$



$$T = 2.3Q/4\pi a_1 ; T_{app} = 2.3Q/4\pi a_2 \rightarrow \alpha = 360 T_{app} / T = 360 a_1 / a_2$$

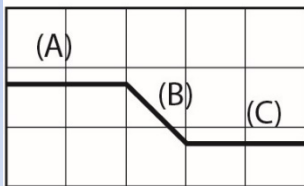
Dual radial sequences: multiple impermeable frontiers

Positive offset



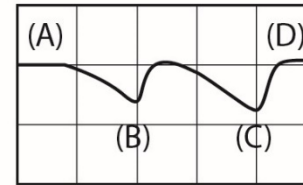
1. No-flow single ($D = 2$) or multiple ($D > 2$) boundaries
2. Weakly inclined conductive fault (low transition slope)

Negative offset

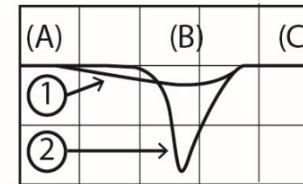


1. Cryptic (non-pumped) contiguous or distal aquifer (highly productive)
2. Aquifer with inclined substratum ($B : n = 3$)

No offset



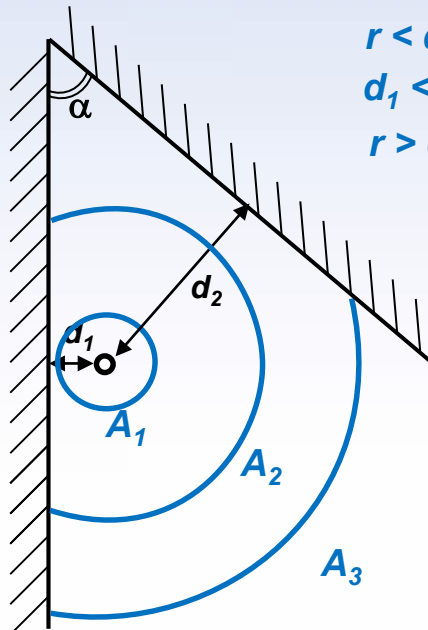
Triple porosity (fractured rocks)



Dual porosity (fractured rocks)

1. Transient interporosity flow
2. Pseudo-steady interporosity flow

CONCEPTUAL MODEL: generalized multiple impermeable frontiers model ; with intermediate stage ($d_2 \gg d_1$)



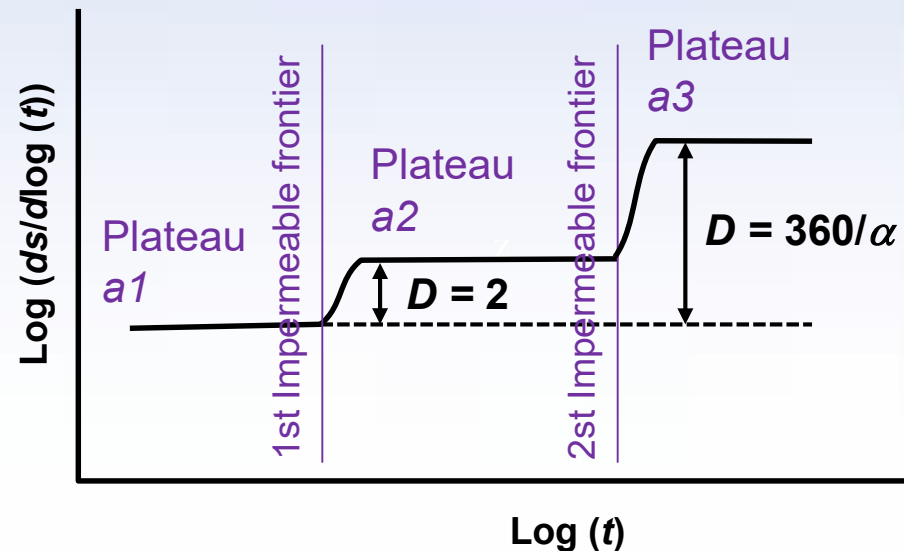
$$r < d_1 : A_1(r) = 2\pi r b$$

$$d_1 < r < d_2 : A_2(r) = \frac{1}{2} A_1(r)$$

$$r > d_2 : A_3(r) = (\alpha/360) A_1(r)$$

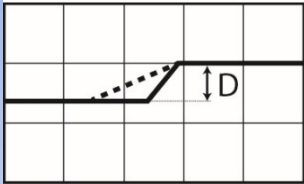
The transmissive surface A_3 is decreased by a factor $360/\alpha$
 → Drawdown rate increases by an equal factor

Log-derivative response
 Sequence of $n : 2 - 2 - 2$
 with successively $D = 2$ and $D > 2$



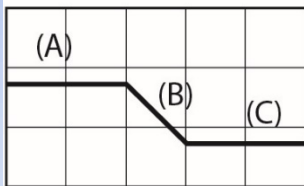
Dual radial sequences: contiguous aquifers

Positive offset



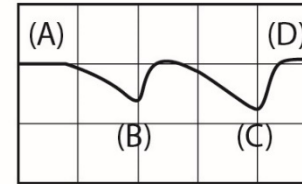
1. No-flow single ($D = 2$) or multiple ($D > 2$) boundaries
2. Weakly inclined conductive fault (low transition slope)

Negative offset

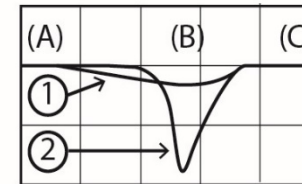


1. Cryptic (non-pumped) contiguous or distal aquifer (highly productive)
2. Aquifer with inclined substratum ($B : n = 3$)

No offset

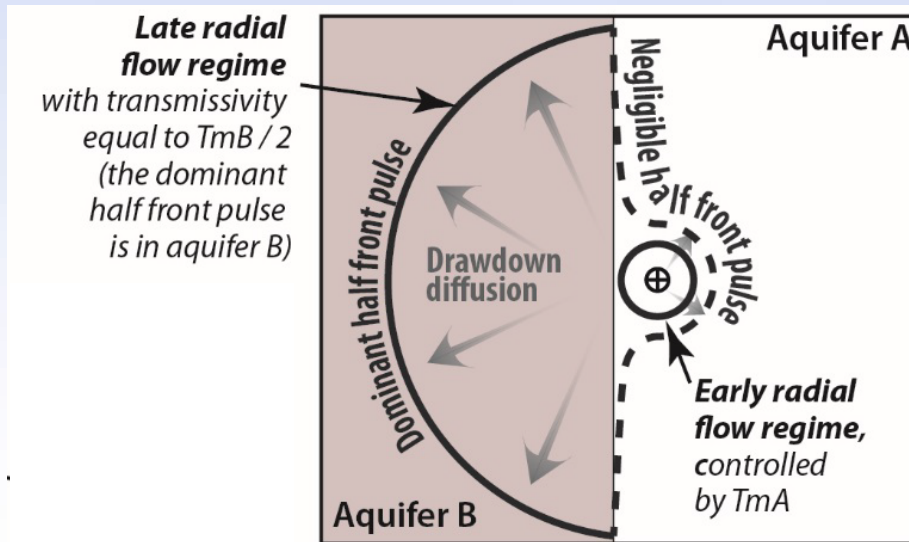


Triple porosity (fractured rocks)

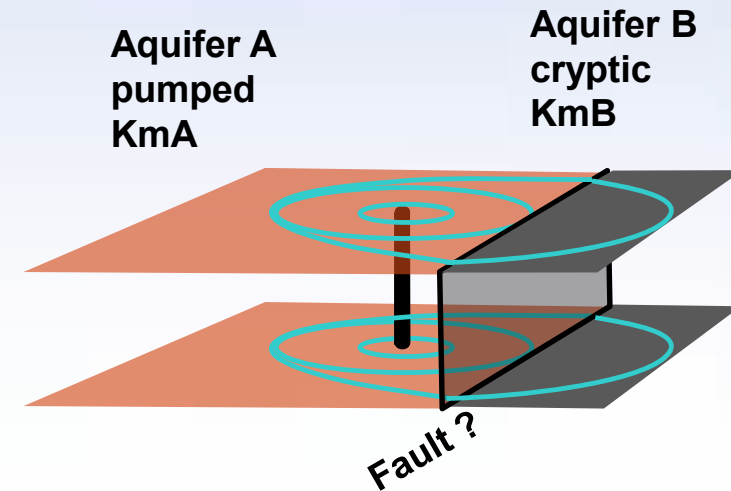


Dual porosity (fractured rocks)

1. Transient interporosity flow
2. Pseudo-steady interporosity flow



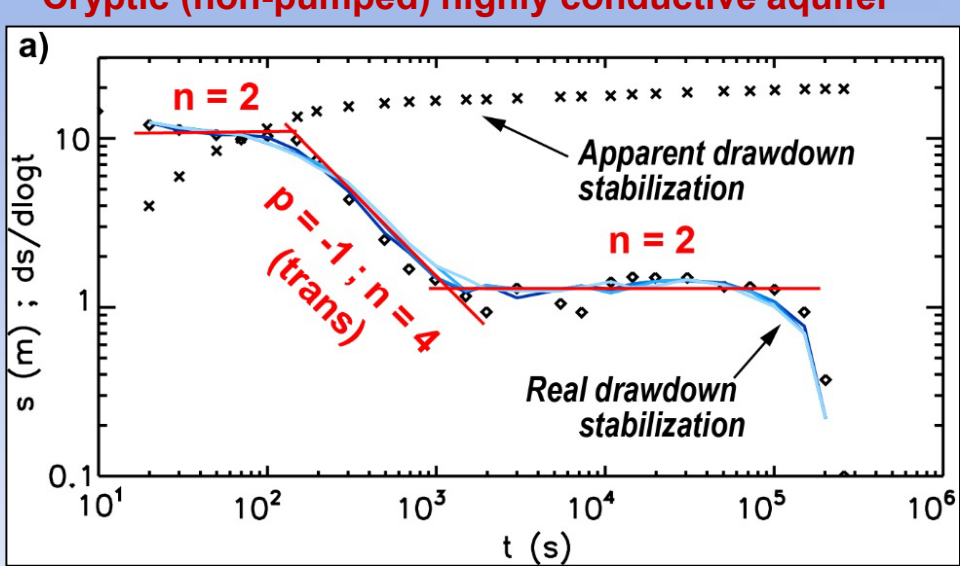
Non-pumped domain is more transmissive:
cryptic aquifer model $\rightarrow K_{app} = K_{mB} / 2$



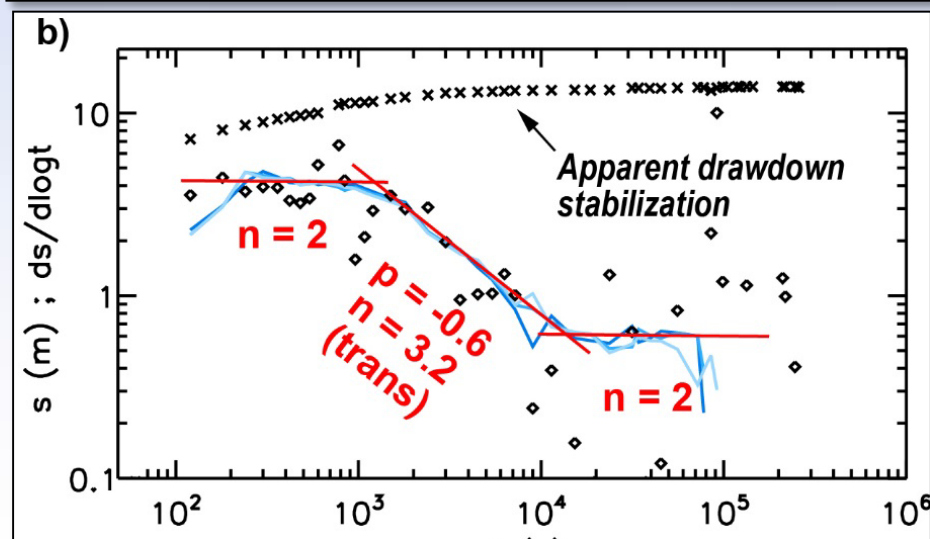
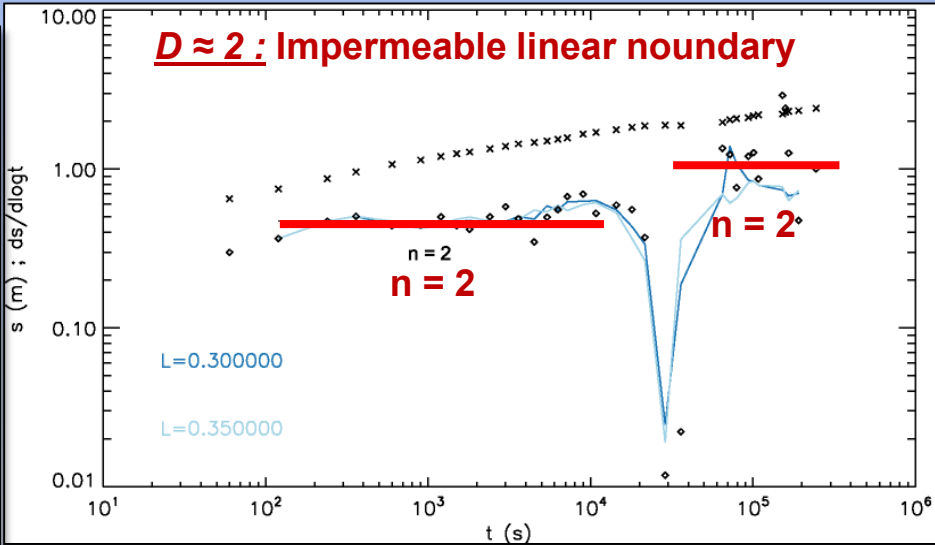
Dual radial sequences: real examples

Negative offset :

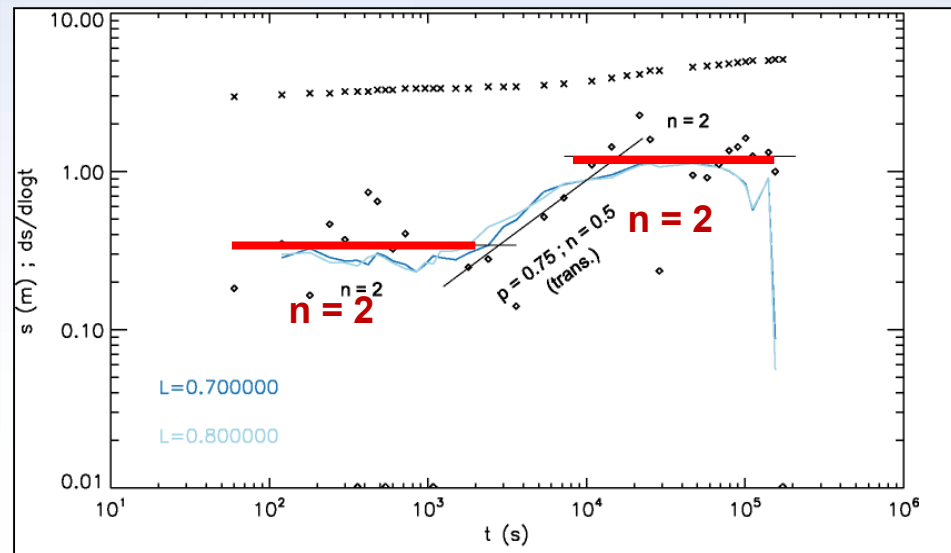
Cryptic (non-pumped) highly conductive aquifer



Positive offset



$D > 2$: Multiple linear boundaries or a horizontal conductive fault

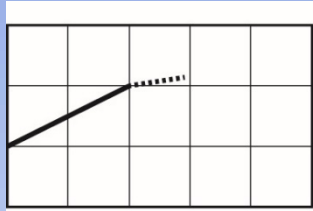


Linear, dual linear combinations

$$**n = 2 - 1**$$

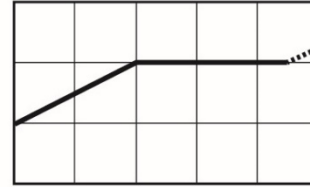
$$**n = 1 - 1**$$

Linear combinations: overview



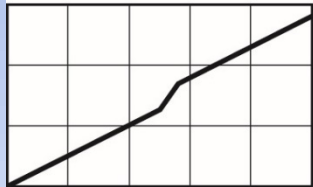
1

1. Fluvial channels (delta fans, glacial deposits, etc.)
2. Conductive fault in an aquitard
3. Infinitely conductive fault
4. Elongated aquifer



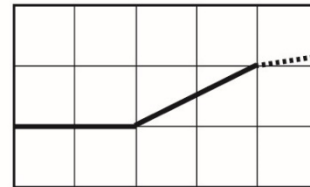
1 - 2

Dense and connective fractured network (continuum-like)



1 - 1

1. T shaped fluvial channels
2. Impermeable boundary in an elongated aquifer



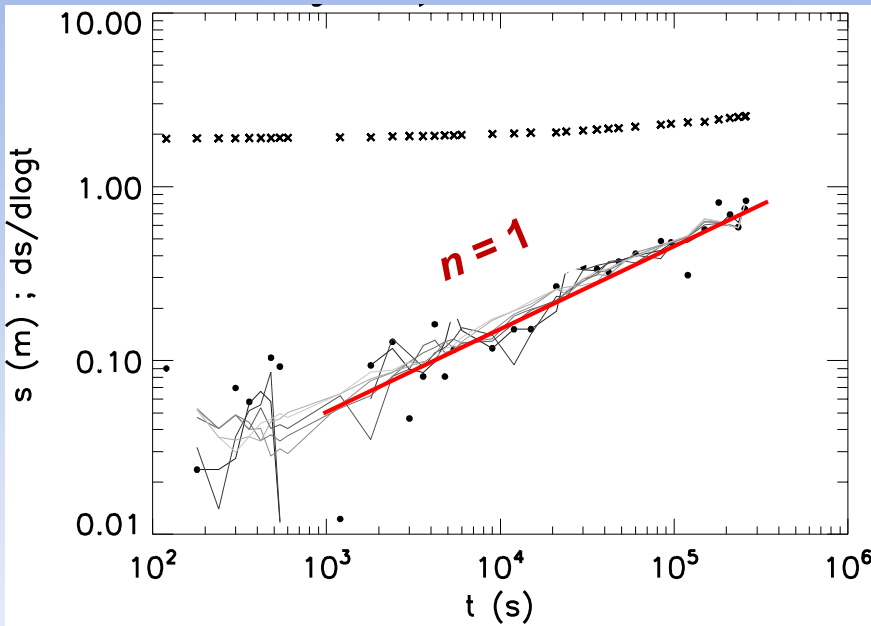
2 - 1

Elongated aquifer (large channel)

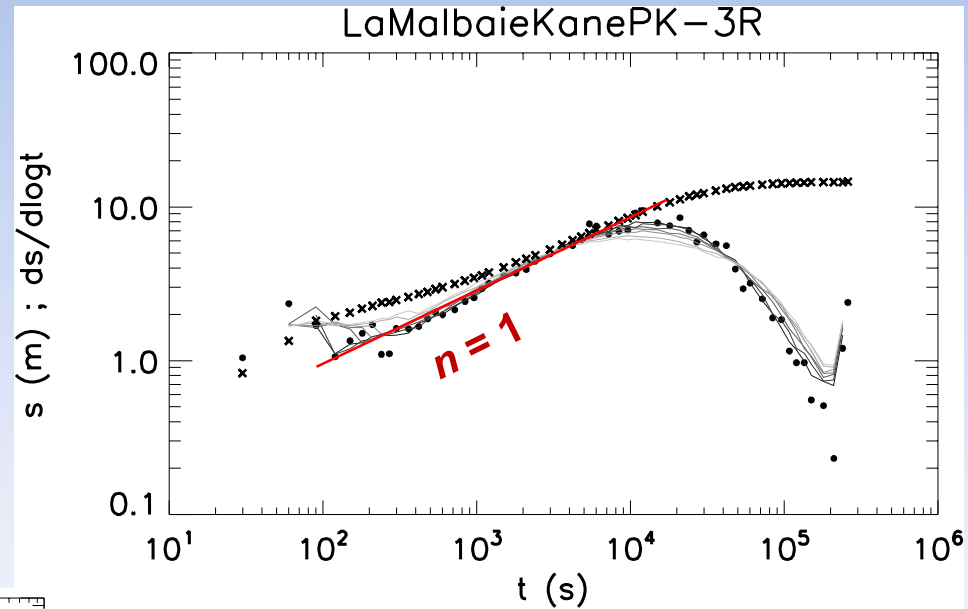
- Linear response = **laterally restricted flow**
- Lateral restriction may be caused by
 - Impermeable boundaries: **fluvial channel, elongated aquifer** models
 - Hardrocks aquifers :
 - **High diffusivity ratios** between a conductive fracture or fault and the **surrounding aquifer or aquitard**
 - **Delayed pressure transfers** to the matrix due to either skin effects on the fractures walls or impermeable material in a layered fault's core zone
- Features allowing to decipher between these various models are: **the time of occurrence** (early, medium or late time pumping time) and the **associations to prior or late radial stage**

Linear combinations: real examples

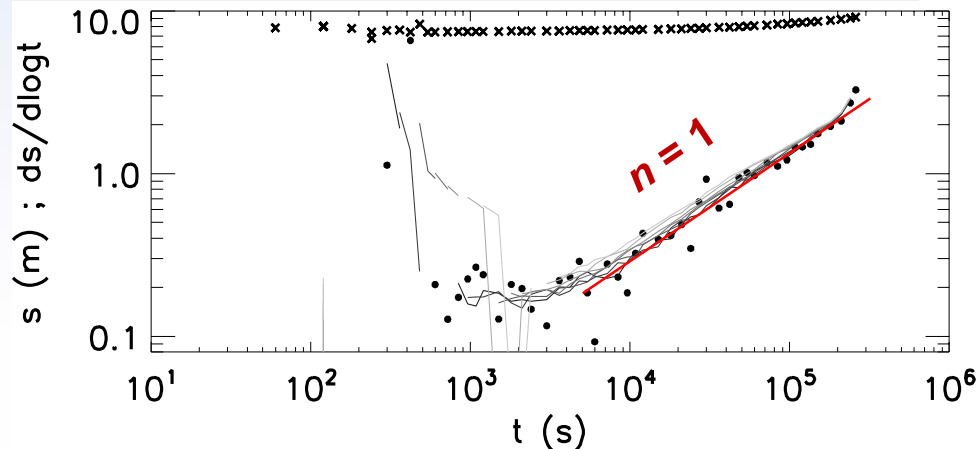
Deltaic deposits aquifer



Faulted hardrock aquifer

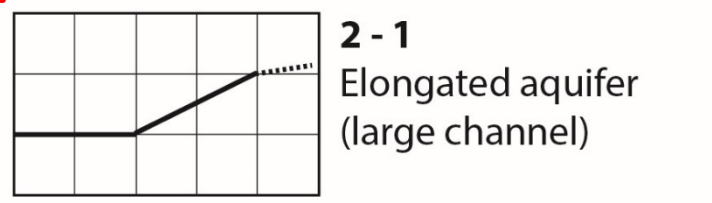
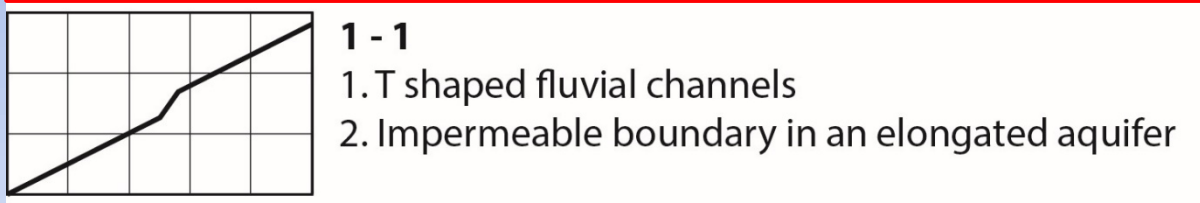
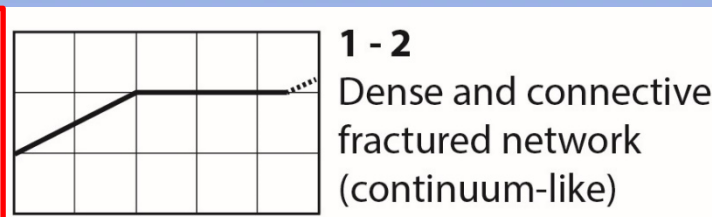
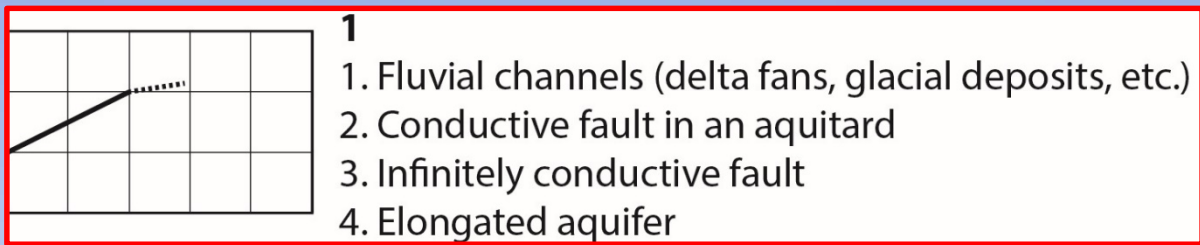


Fluvio-glacial deposits aquifer



- Occurs in short or long time ?
- Infinite-acting ?
- Association to other regimes ($n = 2, n = 1.5$) ?

Linear regime: elongated (channel) aquifer

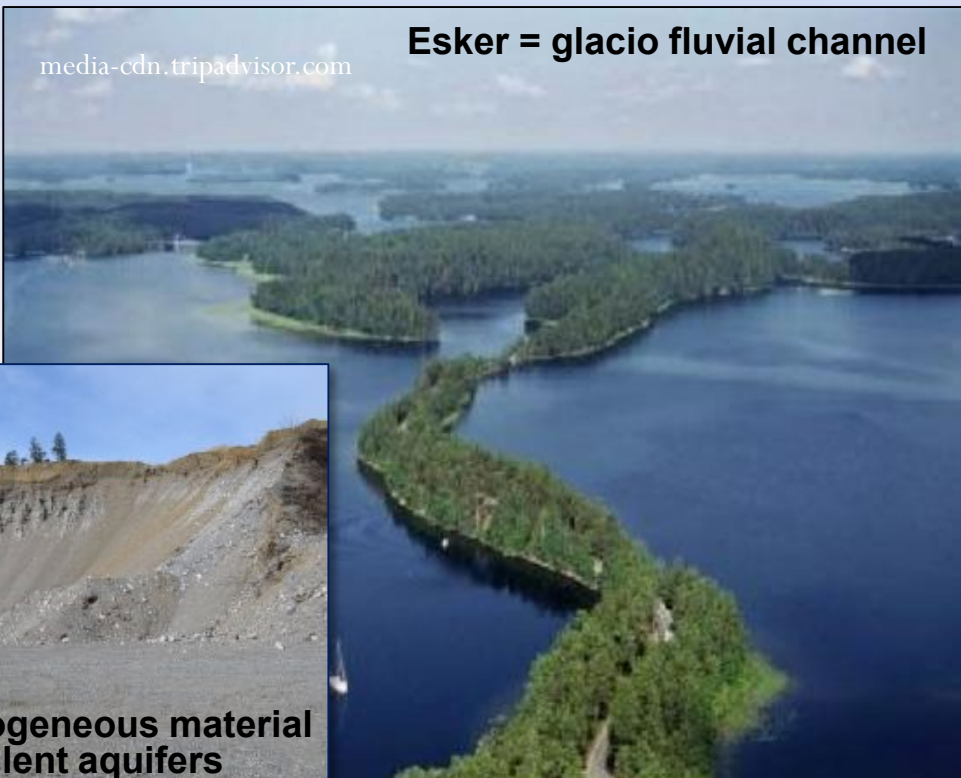
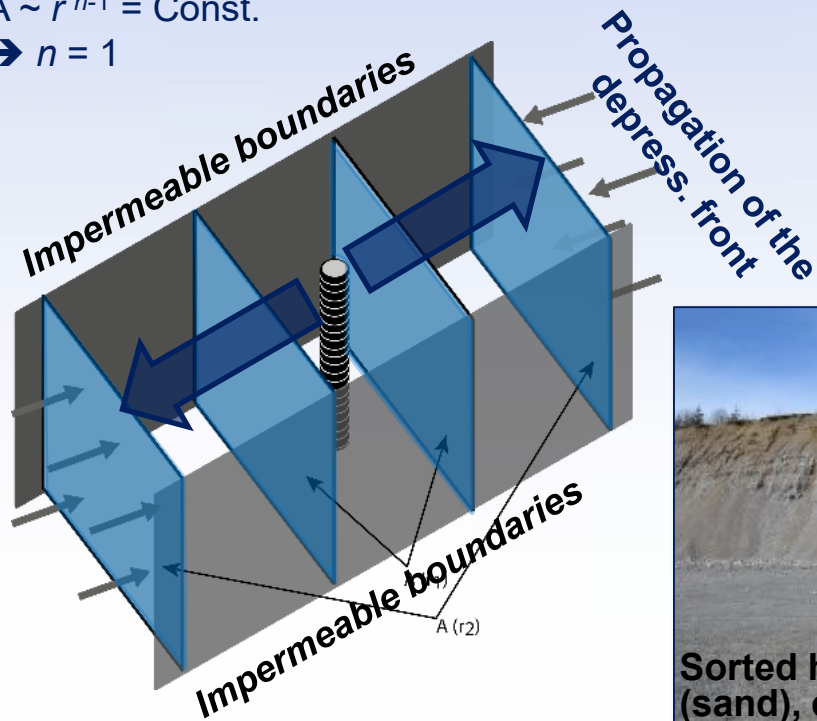


Lateral flow restriction due to two impermeable opposite boundaries

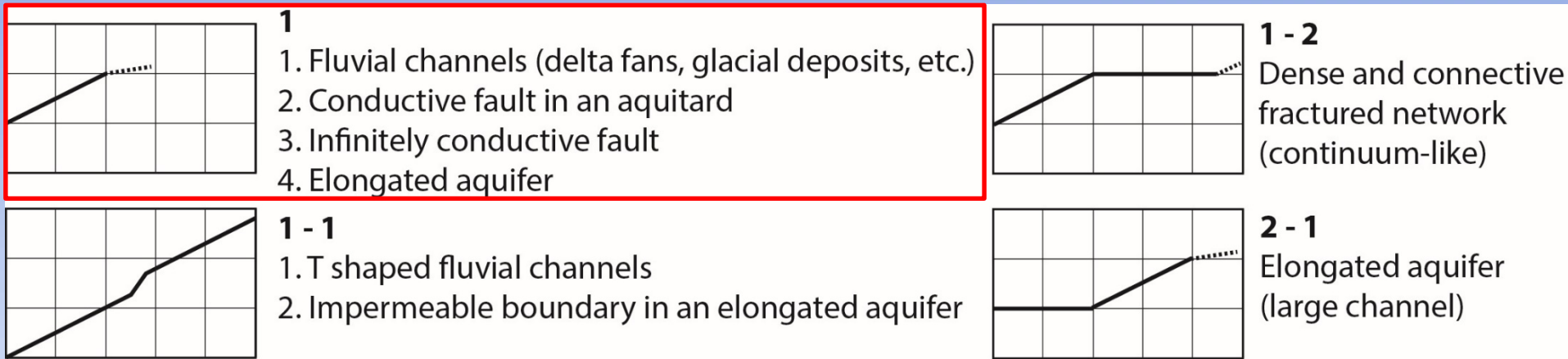
Constant cross flow surface

$$A \sim r^{n-1} = \text{Const.}$$

$$\rightarrow n = 1$$



Linear regime: faulted/fractured aquifers

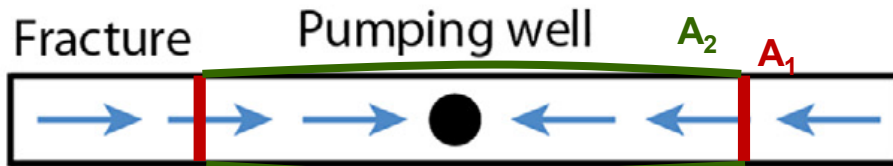


Lateral flow restriction due to **high fracture/matrix diffusivity contrast** or **delayed pressure transfers (skin)** (or confinement into a fault's internal impermeable core zone)

Fault-related linear flow regime
 Introduced in Cinco-Ley and Samaniego (1981)
 finite conductivity fracture model

The highest diffusivity contrasts, the longest the linear regime

a) Fault-related linear flow regime ($n = 1$)
 Matrix depressurization is negligible

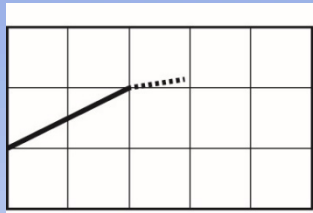


The constant cross flow surface = the fracture's section A_1
 → Linear stage remains as long as $A_1 > A_2$

→ In hard rocks aquifers, the linear regime is typically a **short-duration, early stage regime**

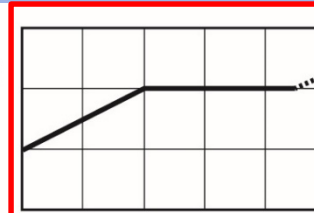
→ After $A_2 > A_1$: radial flow regime ;
 sequence 1 – 2 for densely fractured bedrock aquifers

Linear regime: faulted/fractured aquifers



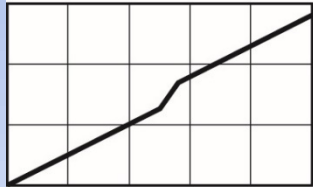
1

1. Fluvial channels (delta fans, glacial deposits, etc.)
2. Conductive fault in an aquitard
3. Infinitely conductive fault
4. Elongated aquifer



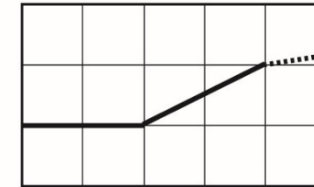
1 - 2

- Dense and connective fractured network (continuum-like)



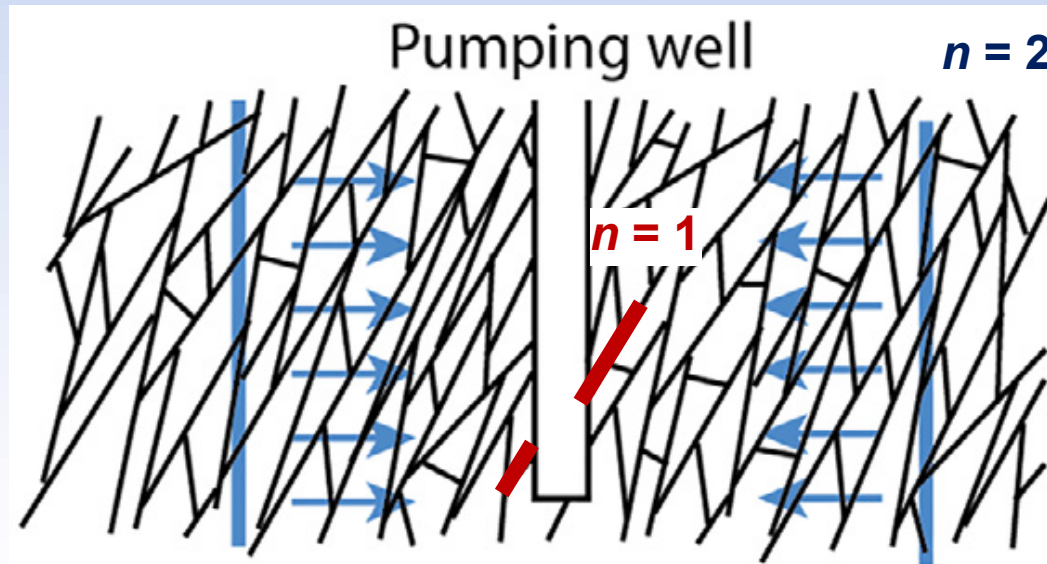
1 - 1

1. T shaped fluvial channels
2. Impermeable boundary in an elongated aquifer



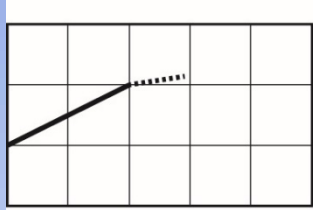
2 - 1

- Elongated aquifer (large channel)

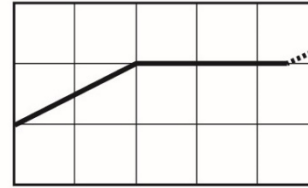


Sequence 1 – 2 with increasing observation scale (front pulse diffusion)

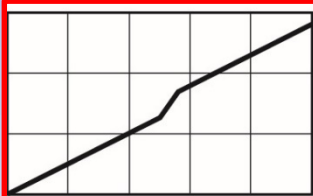
Dual linear regime: fluvial channel



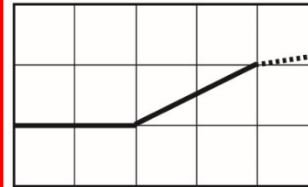
- 1**
1. Fluvial channels (delta fans, glacial deposits, etc.)
 2. Conductive fault in an aquitard
 3. Infinitely conductive fault
 4. Elongated aquifer



- 1 - 2**
- Dense and connective fractured network (continuum-like)



- 1 - 1**
1. T shaped fluvial channels
 2. Impermeable boundary in an elongated aquifer



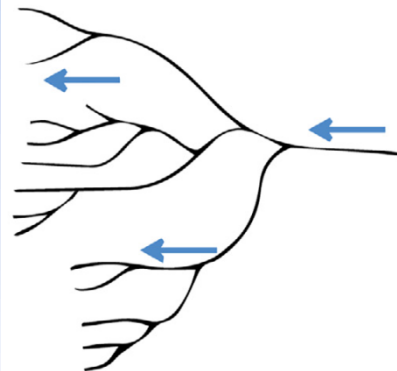
- 2 - 1**
- Elongated aquifer (large channel)

Two successive linear regimes

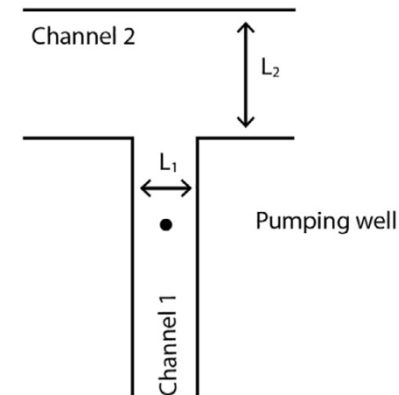
T shaped channel model:
channel enlargement or junction of two channels

Channel network in deltaic fans

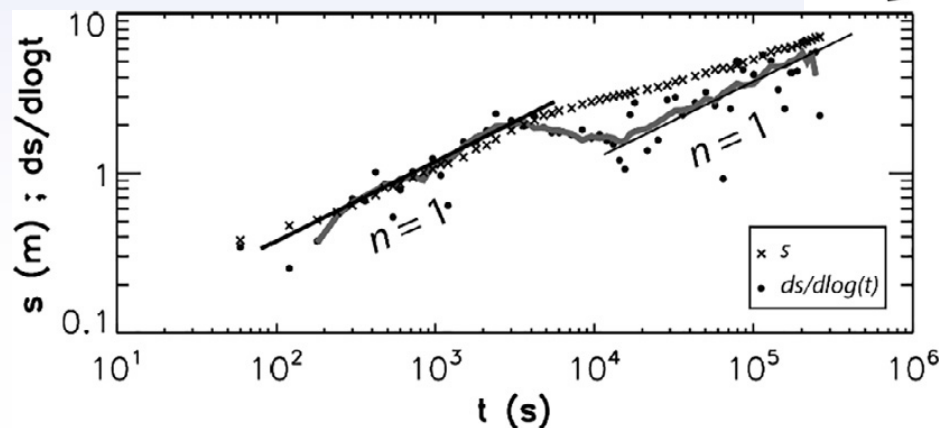
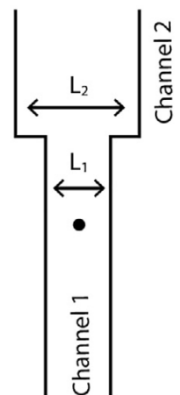
a) Distributary channels



b) T-shaped channels (perpendicular channels)



c) T-shaped channels (successive channels)



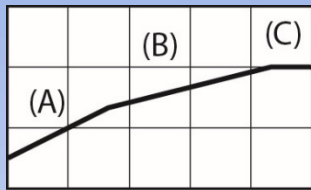
$n = 1.5$ « bilinear » combinations

$$n = 1 - 1.5 - 2$$

$$n = 2 - 1.5$$

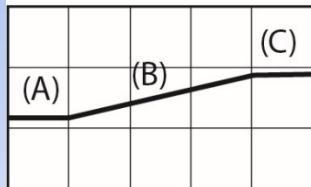
$$n = 2 - 4 - 1.5 - 2$$

$n = 1.5$ « bilinear » combinations: overview



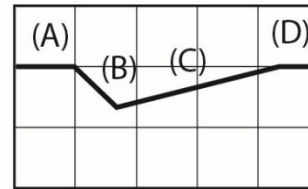
(1) - 1.5 - 2

Strongly inclined conductive fault
connected to the well



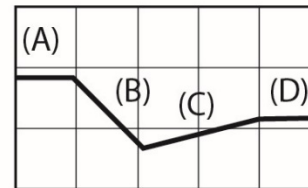
2 - 1.5 - 2

Weakly inclined conductive fault



2 - 4 - 1.5 - 2

Strongly inclined conductive fault
non-connected to the well

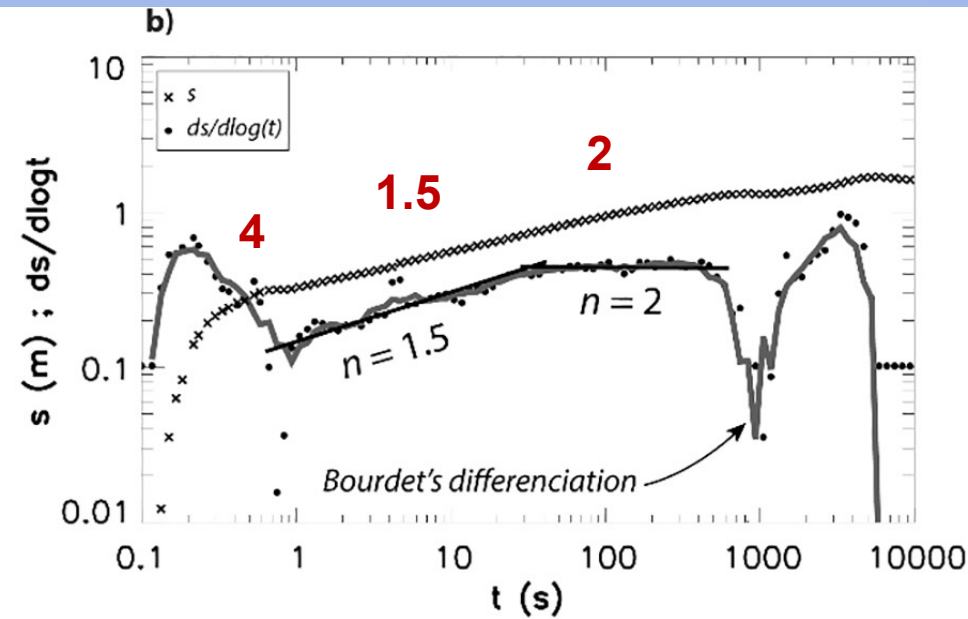
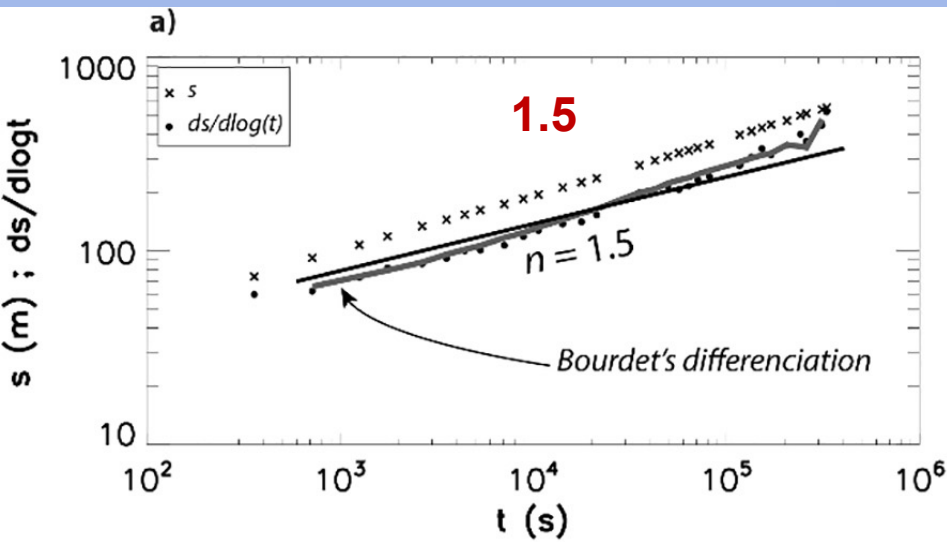


2 - 4 - 1.5 - 2 with offset between
the two radial plateaus :

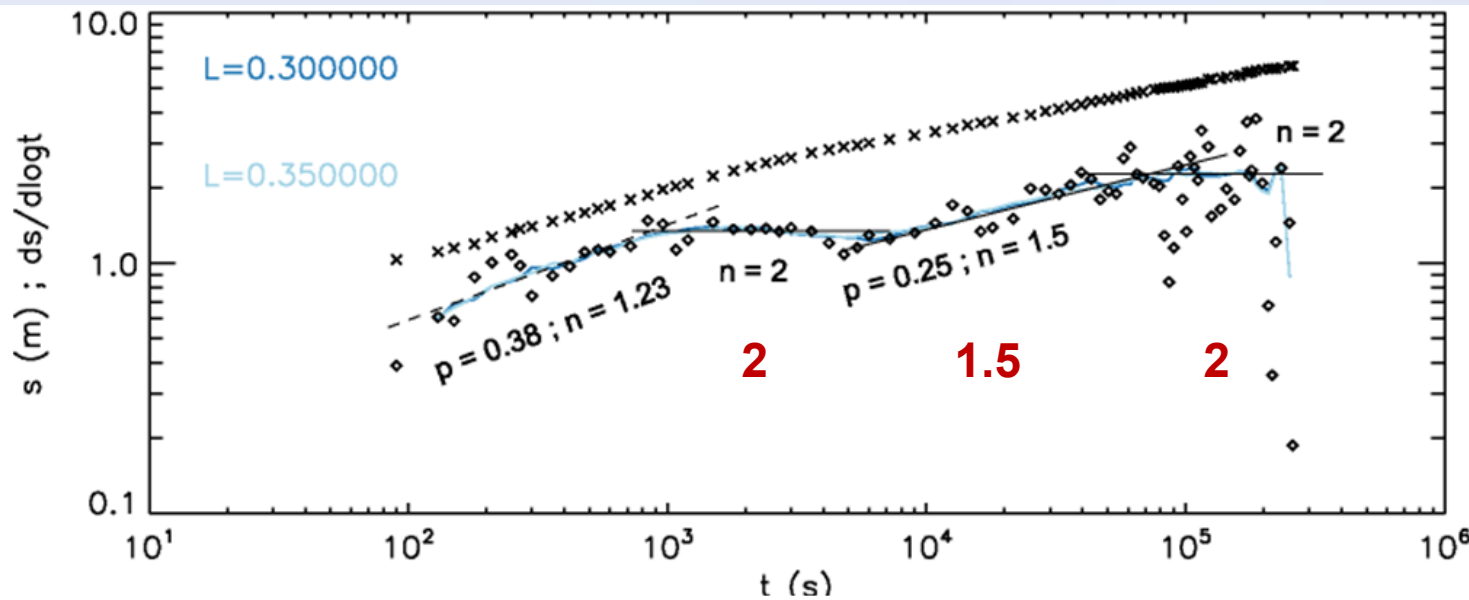
Strongly inclined conductive fault
non-connected to the well separating
two aquifers with distinct properties

- The $n = 1.5$ flow regime, early referred to as bilinear flow regime, has been long recognized as produced by a **vertical conductive fault embedded into a non-impermeable aquifer** (Cinco-Ley and Samaniego, 1981; Rafini and Larocque, 2009, 2012)
- The associated flow regimes with $n = 1, 2, \text{ or } 4$ indicate whether the fault is distal or directly connected to the pumping well and, in a lower extent, the attitude of the fault into the aquifer

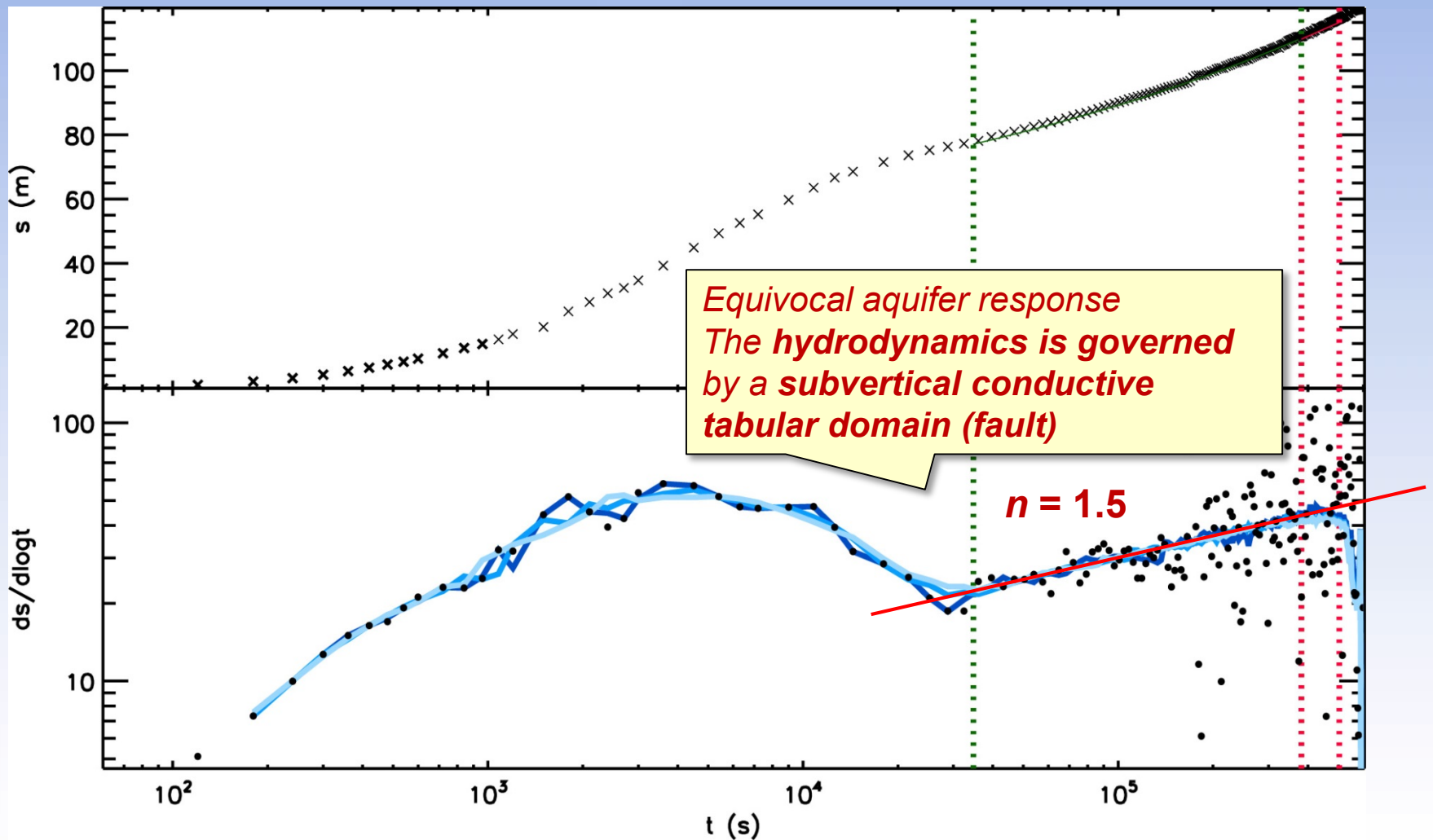
$n = 1.5$ « bilinear » combinations: real examples



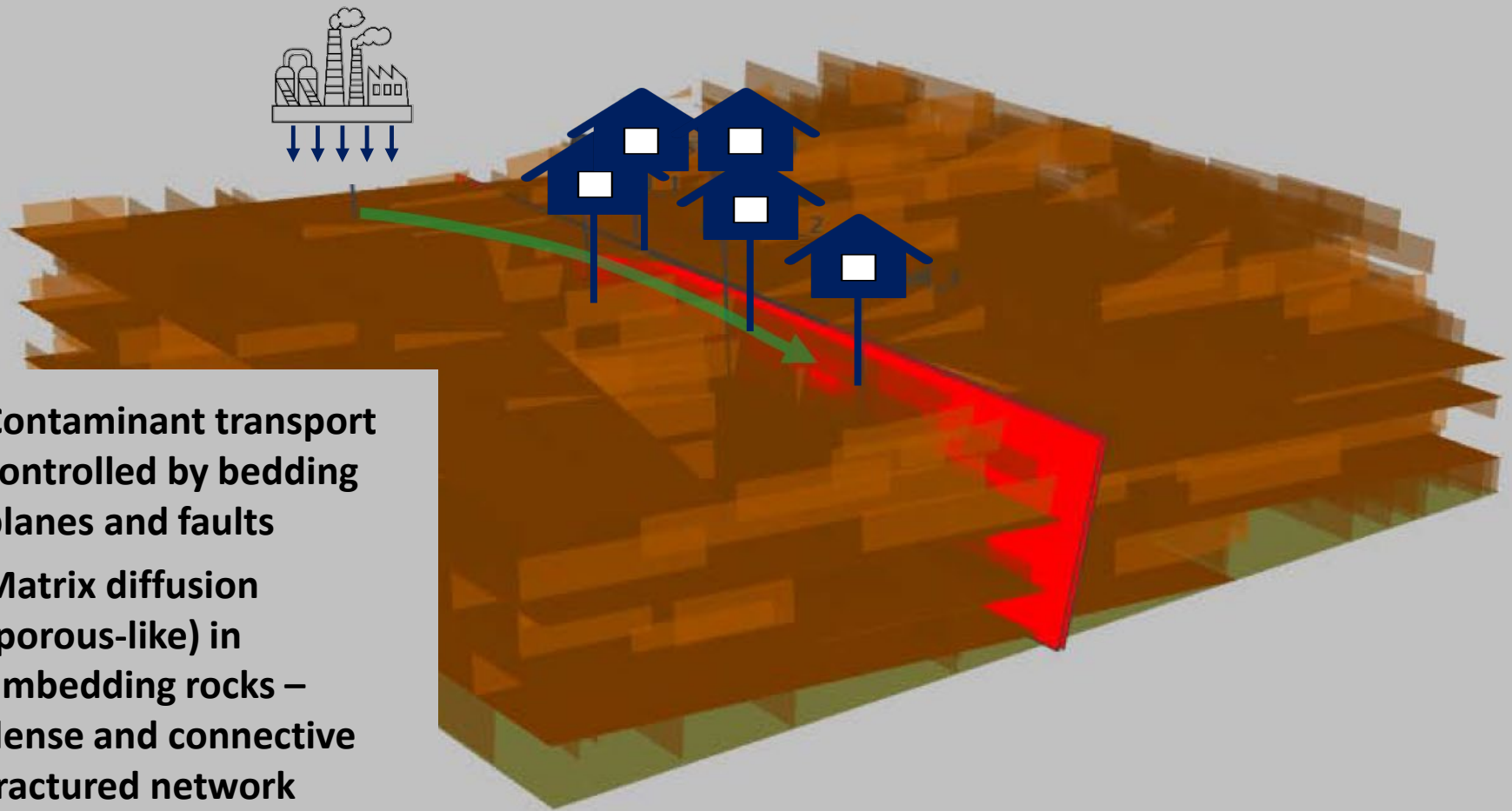
Hard rocks
aquifers



$n = 1.5$ « bilinear » combinations: real examples

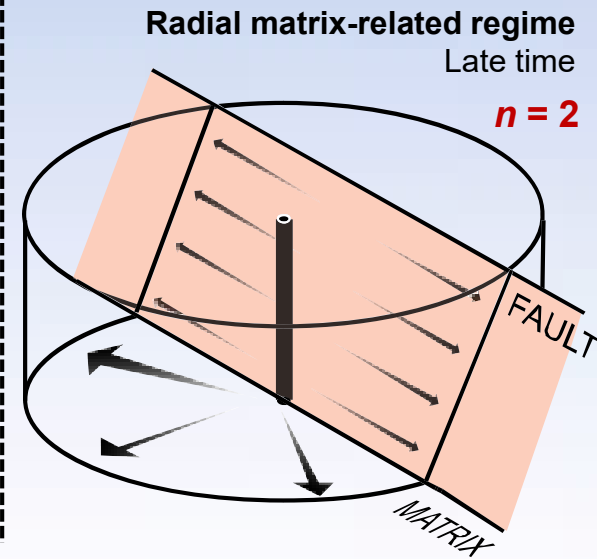
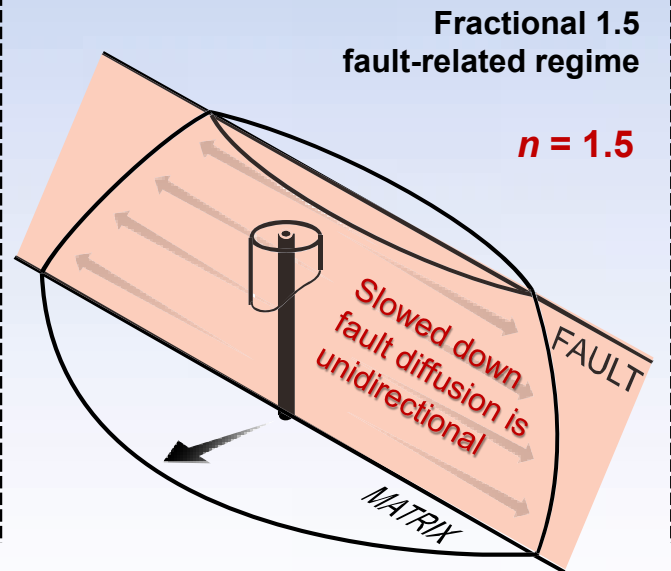
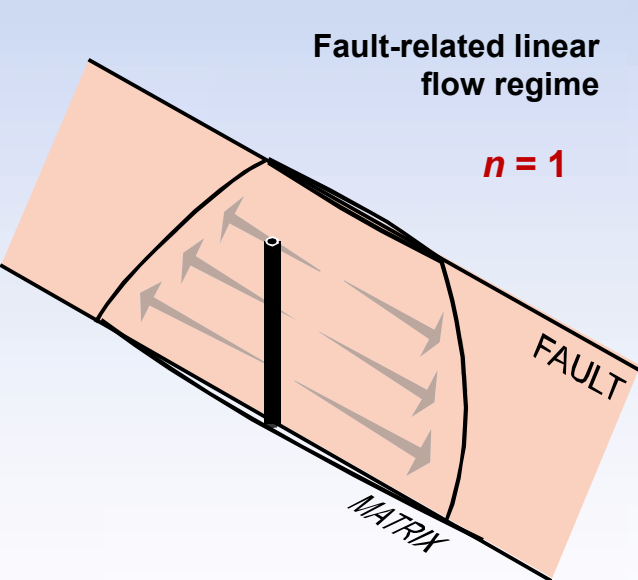
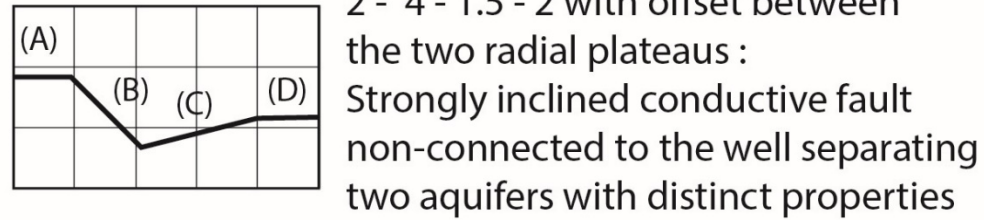
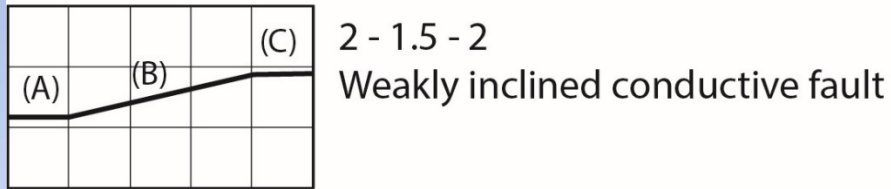
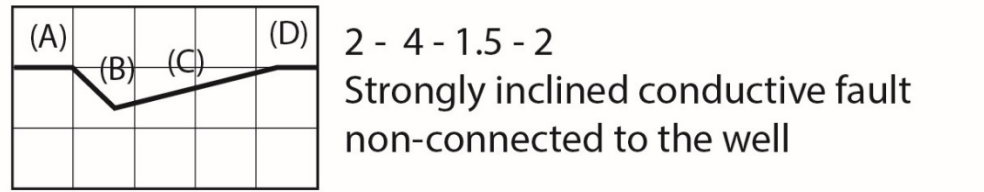
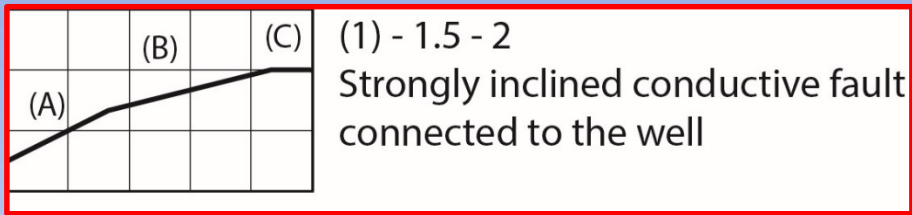


Differentiation (Bourdet, 1983) of dataset measured with levellogger



- Contaminant transport controlled by bedding planes and faults
- Matrix diffusion (porous-like) in embedding rocks – dense and connective fractured network

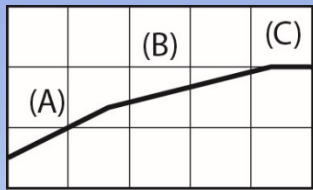
$n = 1.5$ « bilinear » combinations



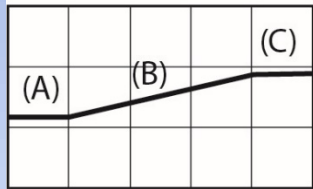
- **Early stage $n = 1$** : matrix depressurization is negligible
- **Mid-stage $n = 1.5$** : the response of the system is governed by fault properties, fault diffusion slow-down
- **Late stage $n = 2$** : the fault does not exert any influence on the hydrodynamics, the response is governed by matrix properties

Rafini and Larocque, 2009, 2012

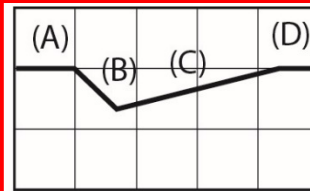
$n = 1.5$ « bilinear » combinations



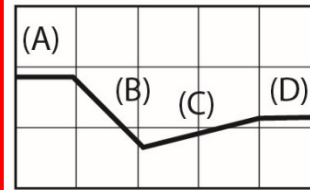
(1) - 1.5 - 2
Strongly inclined conductive fault
connected to the well



2 - 1.5 - 2
Weakly inclined conductive fault



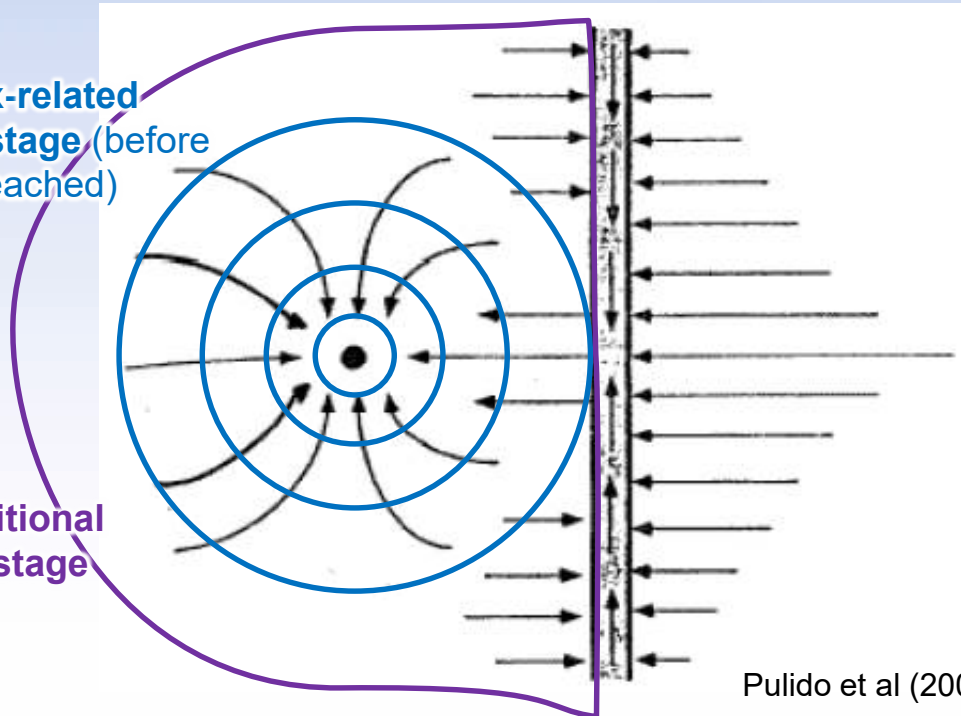
2 - 4 - 1.5 - 2
Strongly inclined conductive fault
non-connected to the well



2 - 4 - 1.5 - 2 with offset between
the two radial plateaus :
Strongly inclined conductive fault
non-connected to the well separating
two aquifers with distinct properties

Early matrix-related
radial flow stage (before
the fault is reached)

Transitional
 $n = 4$ stage



Pulido et al (2003)

**Theoretical model of a not-
connected vertical finite-
conductivity fault:**

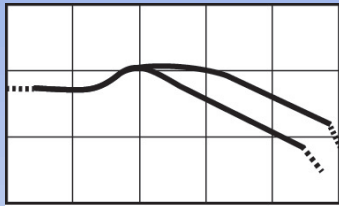
Abbaszadeh and Cinco-Ley
(1995); Rafini and Larocque
(2009)

Spherical combinations

$$n = 3 - 2$$

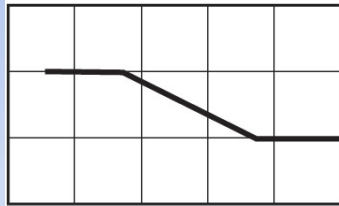
$$n = 2 - 3 - 2$$

Spherical combinations: overview and real examples



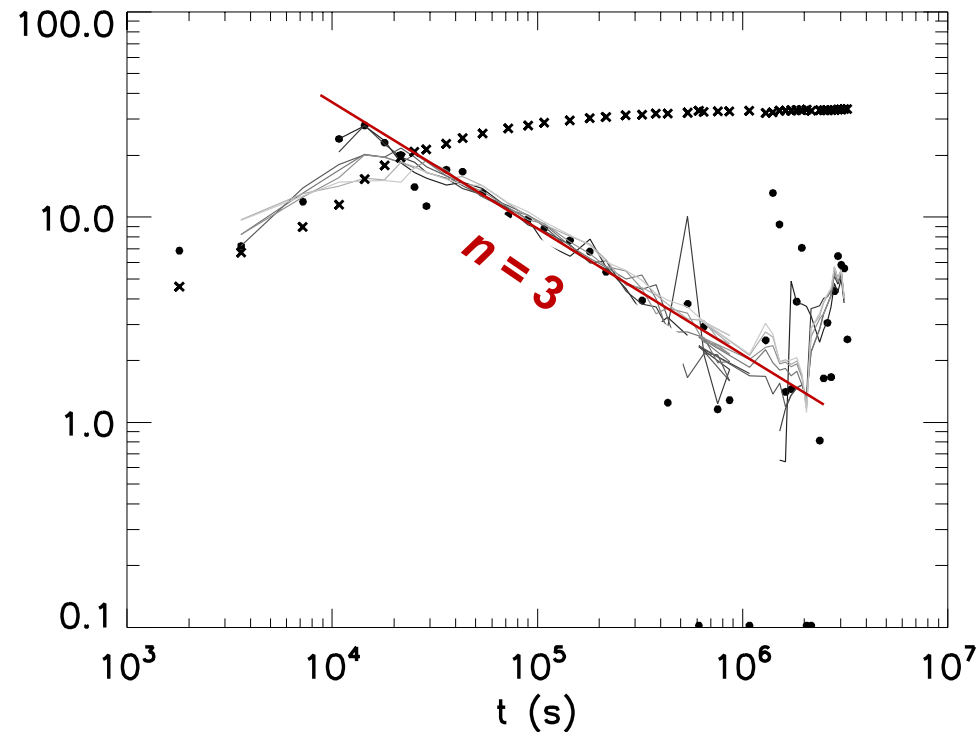
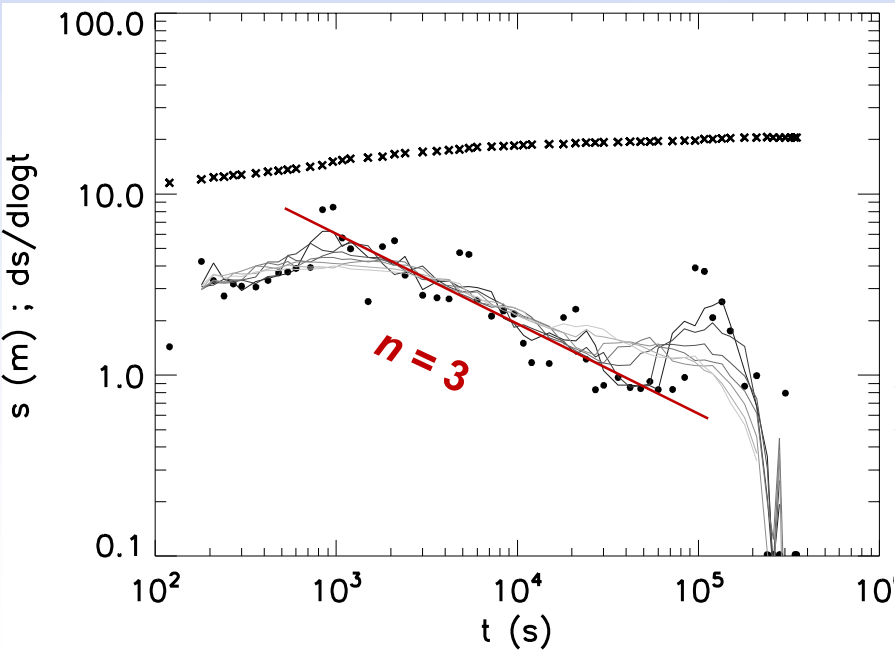
2 - (2) - 3

1. Inclined substratum aquifer with wedge



(2) - 3 - 2

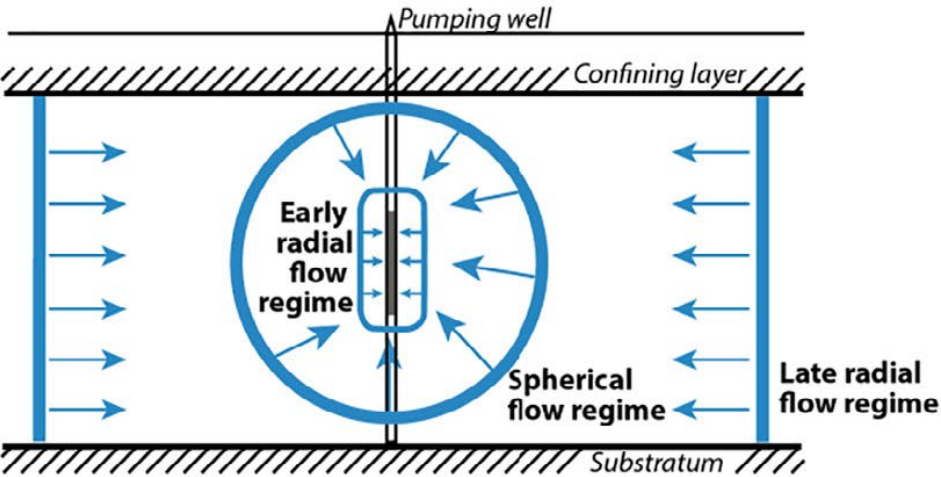
1. (Early radial stage is absent or very short-duration) partially penetrated aquifer or screen length very shorter than the aquifer thickness (or packer tests)
2. Punctual connection to an outer homogeneous aquifer (confinement gap)



Spherical combinations: conceptual models

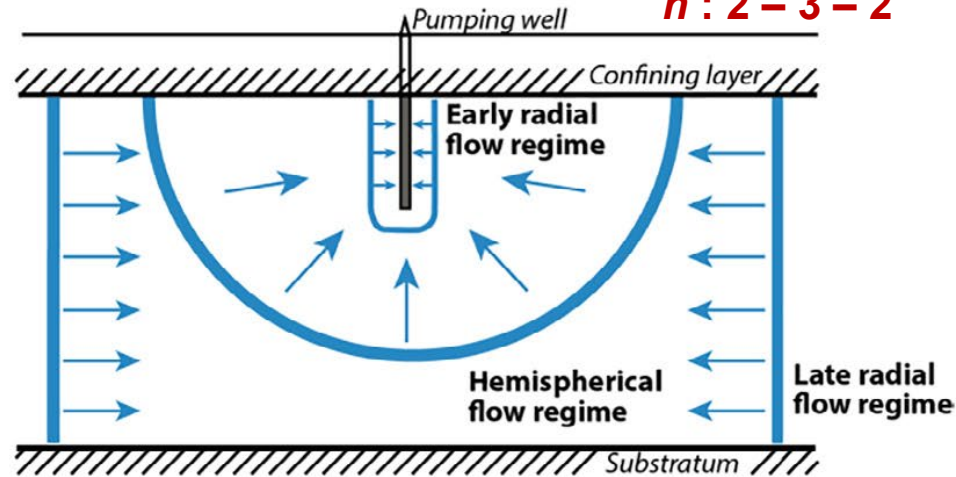
a) Packer test

$$n : 2 - 3 - 2$$



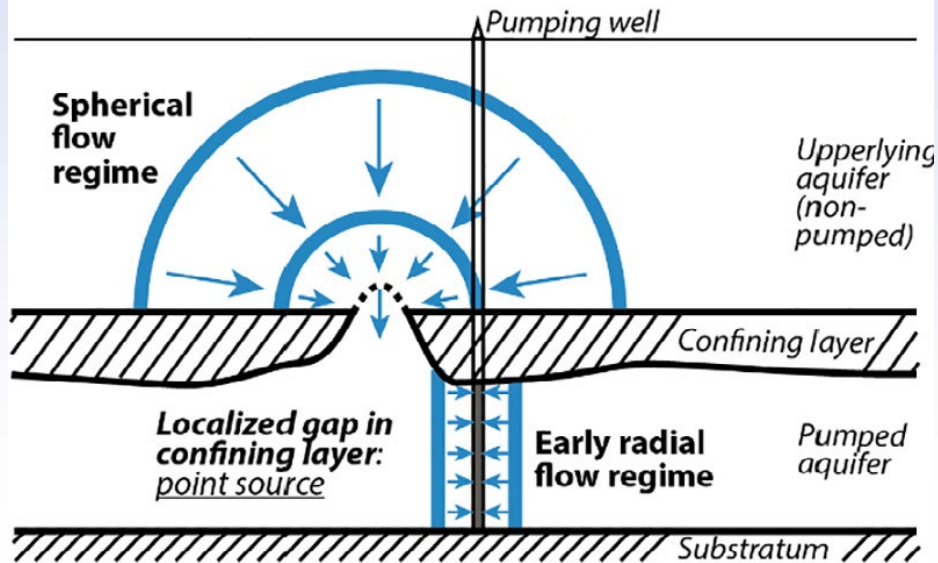
b) Partially penetrating/completing well

$$n : 2 - 3 - 2$$



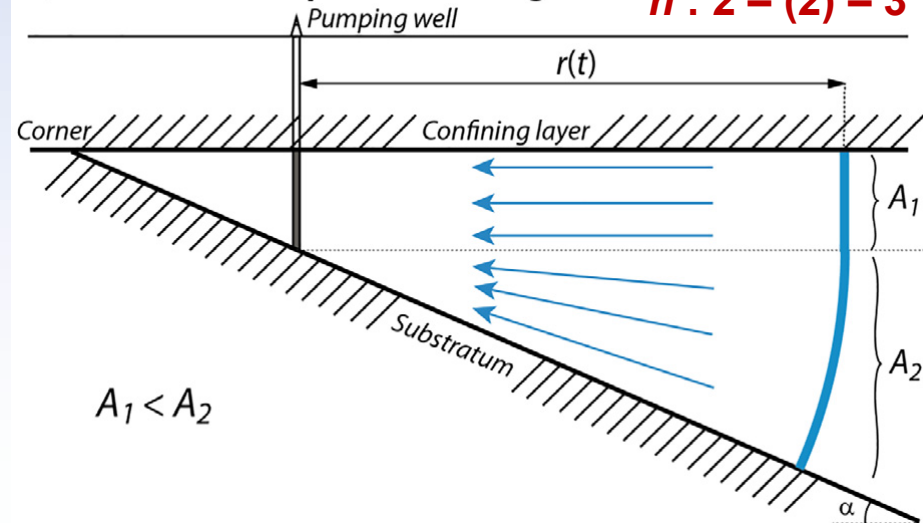
a) Confinement gap

$$n : 2 - 3$$



c) Truncated hemispherical flow regime

$$n : 2 - (2) - 3$$



Discussions

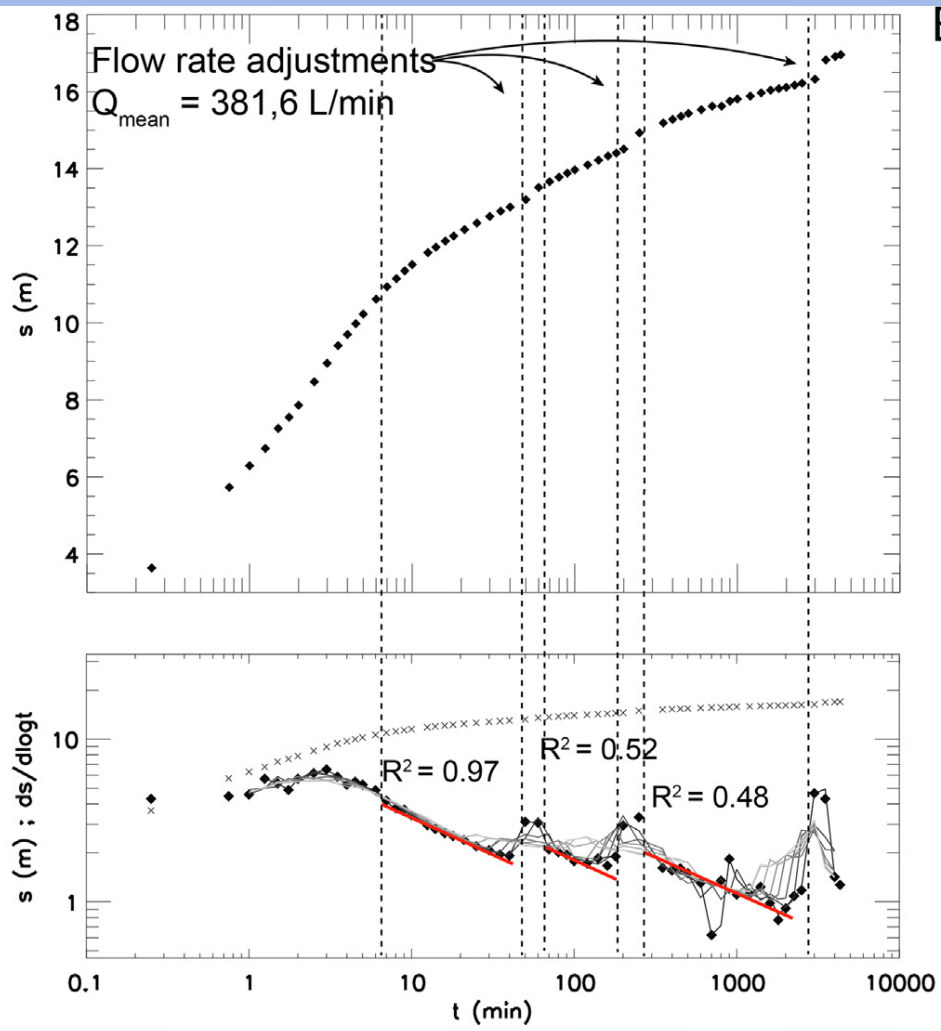
- Global methodology
- Limitations of the approach

Limitations of the submitted methodology

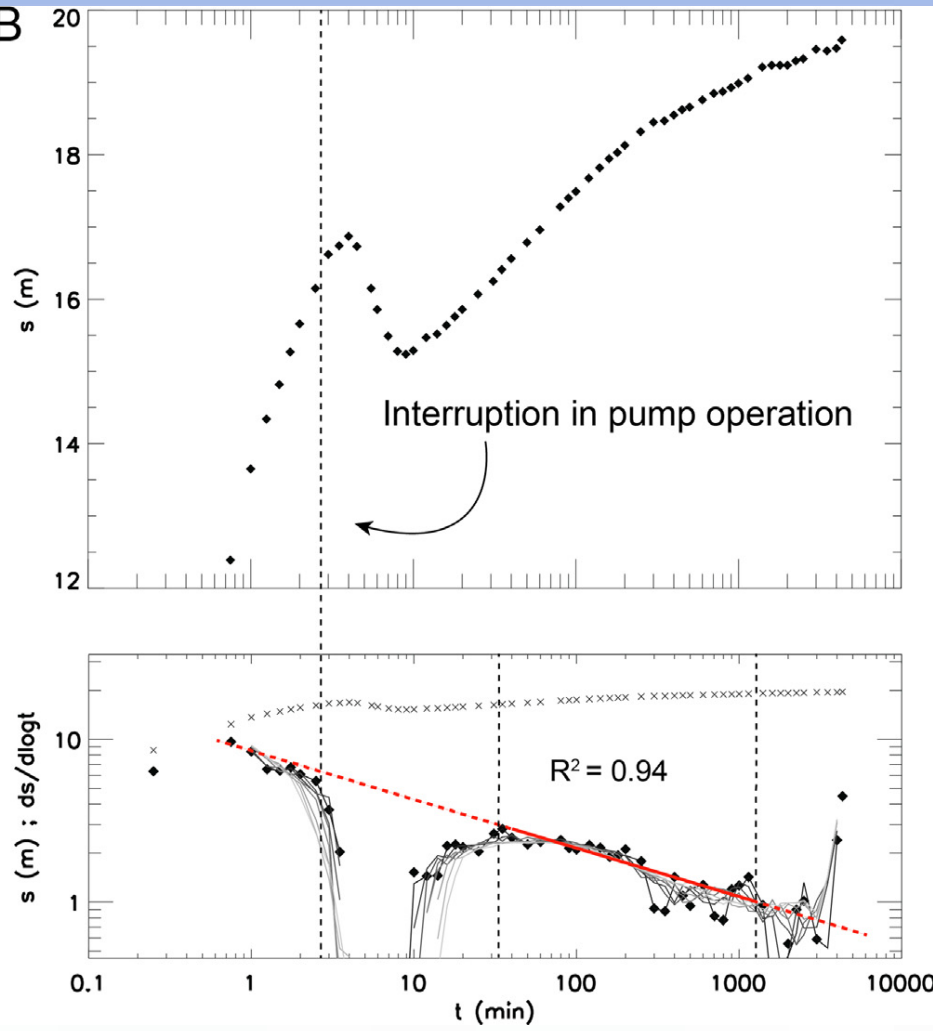
- Non-unique interpretation of the conceptual flow model: different conceptual models predict similar flow regime sequences
 - Look for other data: geological environment, probes data
- Noisy derivative datasets lead to uncertain graphical interpretations
 - Data preprocessing helps...
- Truncated sequences = partial diagnostic

Flow rate unstabilities

A



B



- **Only qualitative diagnostic were presented in this talk**
- **Every conceptual model provides with sets of equations for the estimation of specific hydraulic parameters (oil & gas, hydrogeology)**
- **Integration of observation wells drawdowns series: higher complexity, spatial interperatation of the hydraulic objects**
- **Ongoing PhD - Daouda Meite**

SIREN: *in progress* interface dedicated to sequential analysis

Settings

Flow regimes description

Conceptual model interpretation

The screenshot displays the SIREN software interface. On the left, there are settings panels for 'Correction de Bourdet' (with Lmin, Linc, Lmax values) and 'Affichage plusieurs puits' (with n = 0.00). The main area contains two plots: a 'Drawdown semilog plot' (top) and a 'Drawdown log-derivative bilog plot' (bottom). The top plot shows a curve with a linear segment and a curved segment, with time markers t1, tp1, t2, tp2, tp3, t4, and tp4. The bottom plot shows a similar curve with a linear segment and a curved segment, with time markers t1, tp1, t2, tp2, tp3, t4, and tp4. The bottom plot is annotated with 'GRF normalization' and 'Simultaneous manual fit on bilog plot and semilog plot'. The right side of the interface shows four 'Segment' panels (Segment #1 to #4) with parameters like 'Début', 'Fin', 'Pente', 'ds/dlog(t)=1 sur bilog', and 's(t)=0 sur semilog'. The 'Interprétation' panels on the right show conceptual model options and hydraulic properties.

Drawdown semilog plot

Bourdet's correction for random noise

Simultaneous manual fit on bilog plot and semilog plot

GRF normalization Drawdown log-derivative bilog plot

Adjust the slope

Flow dimension value

Adjust beginning and ending time

Adjust the vertical offsets on bilog plot

Adjust the vertical offsets on semilog plot

n = 4 n = 1.5 n = 2 n = 0

Interpretation: this aquifer is governed by a transmissive steep fault $T_f = 5.86 \times 10^{-3} \text{ m}^2/\text{s}$ in the vicinity of the wellbore (not intercepted), embedded into a transmissive matrix $T_m = 9.49 \times 10^{-4} \text{ m}^2/\text{s}$

+

Late time: closed reservoir = impermeable barriers... (in all directions?)

Thank you

**Provide pump test dataset for our
compilation**

silvain.rafini@gmail.com

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**Natural Sciences and Engineering
Research Council of Canada**

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