



Alpine Hydrogeology: Linking Field Observations to Basin-Scale Hydrology

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Acknowledgements

Students, Technicians, Post-docs, Colleagues

Jaime Hood, Greg Langston, Danika Muir, Andrius Paznekas, James Roy, Alastair McClymont, Larry Bentley, Nathan Green, Kate Forbes,

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Biogeoscience Institute (Univ. of Calgary)

Alberta Ingenuity Centre for Water Research

Canadian Foundation for Climate and Atmospheric
Science (IP3 Network)

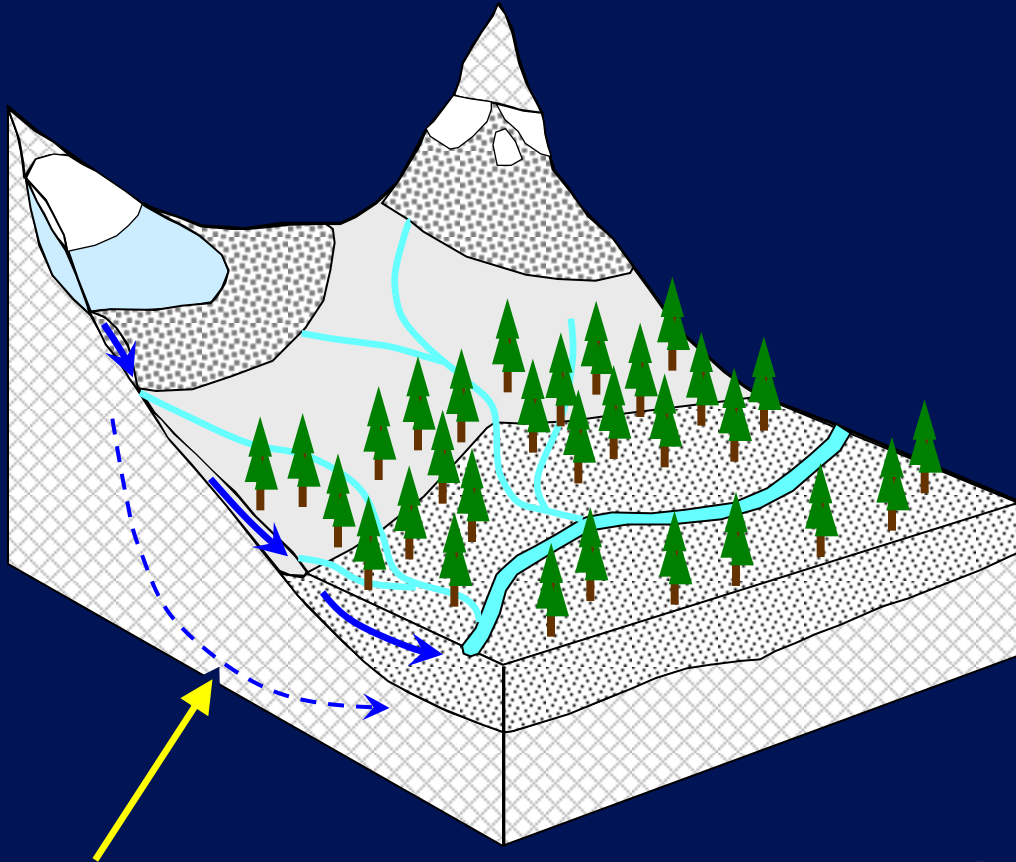
Environment Canada

Logistical Support

Lake O'Hara Lodge

Parks Canada

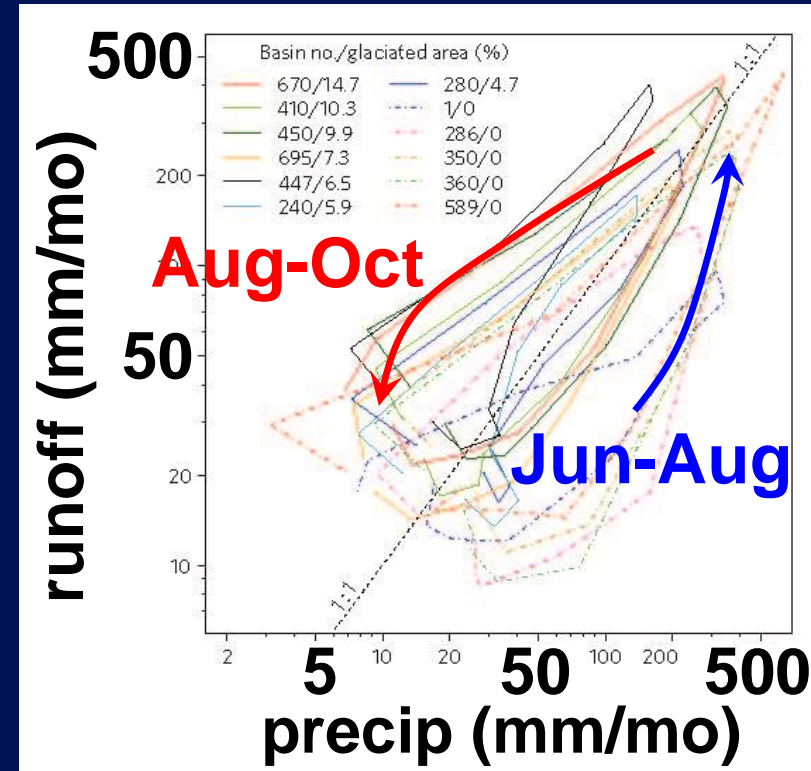
Mountain Groundwater



Mountain block recharge

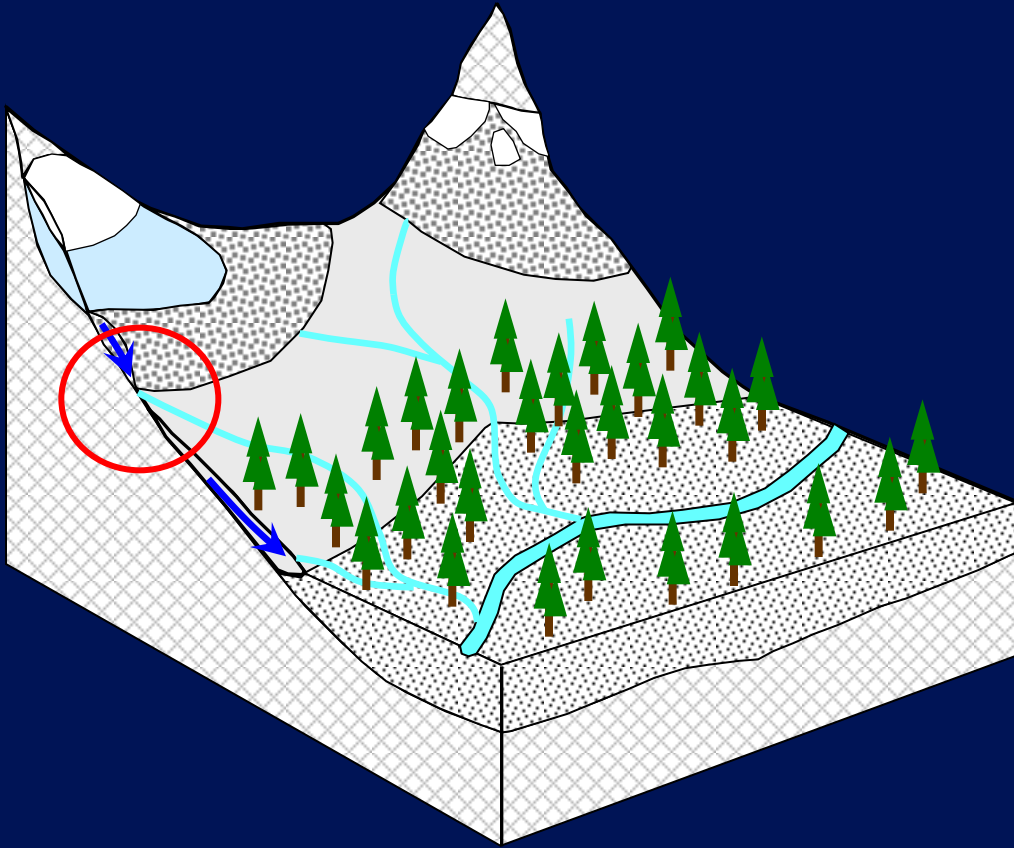
- Manning and Solomon (2005, *Water Resour. Res.*)
- Ajami et al. (2011, *WRR*)
- Welch and Allen (2012, *WRR*)

Himalayan River Basins



Andermann et al. (2013, *Nature GS*)

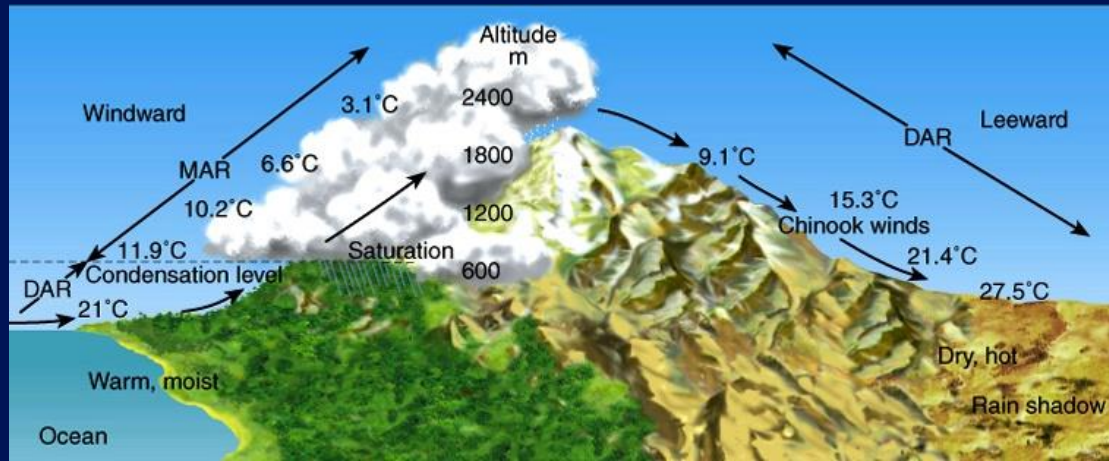
Mountain Groundwater



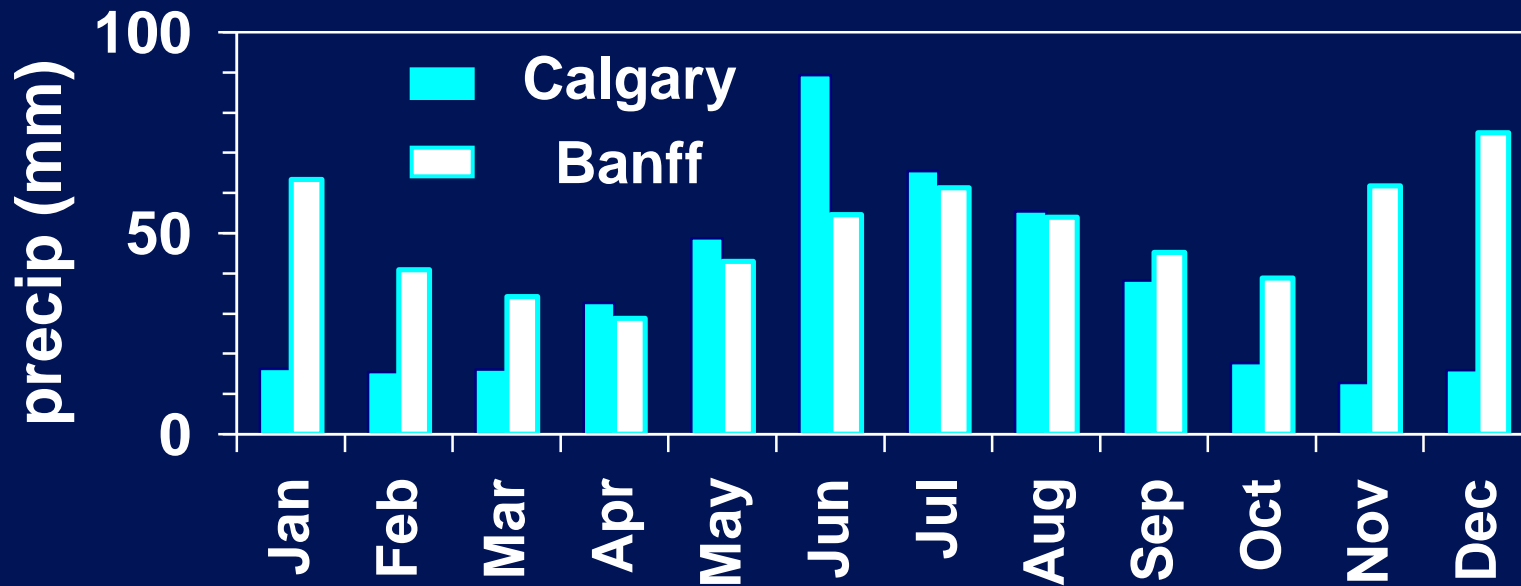
Alpine groundwater sustains critical habitats.

- Lowry et al. (2011, *WRR*)
- Brown et al. (2009, *Freshwater Biol*)

Mountains as 'Water Tower'



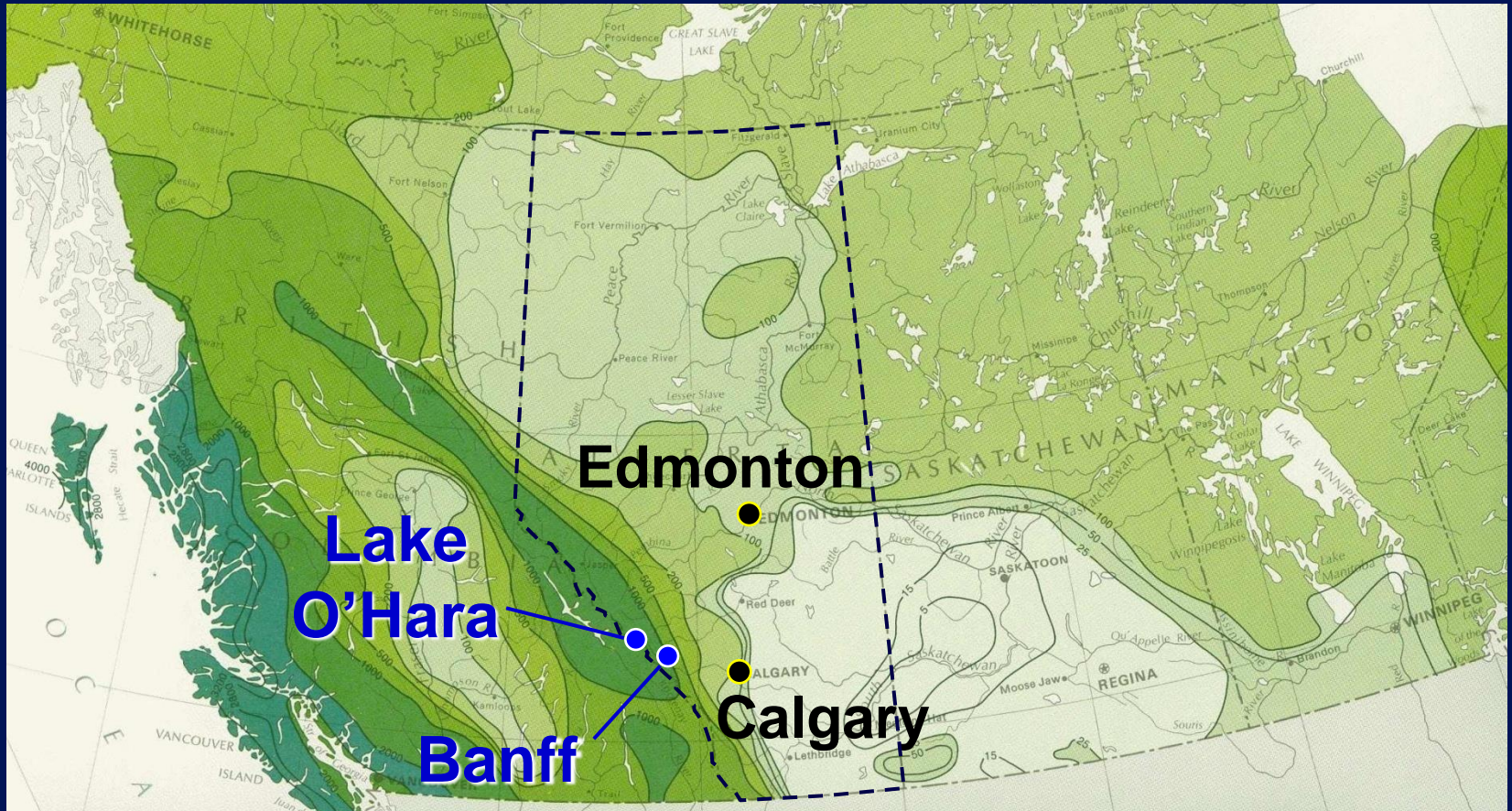
Christopherson (2000)



Mountain Rivers Provide Water Supply



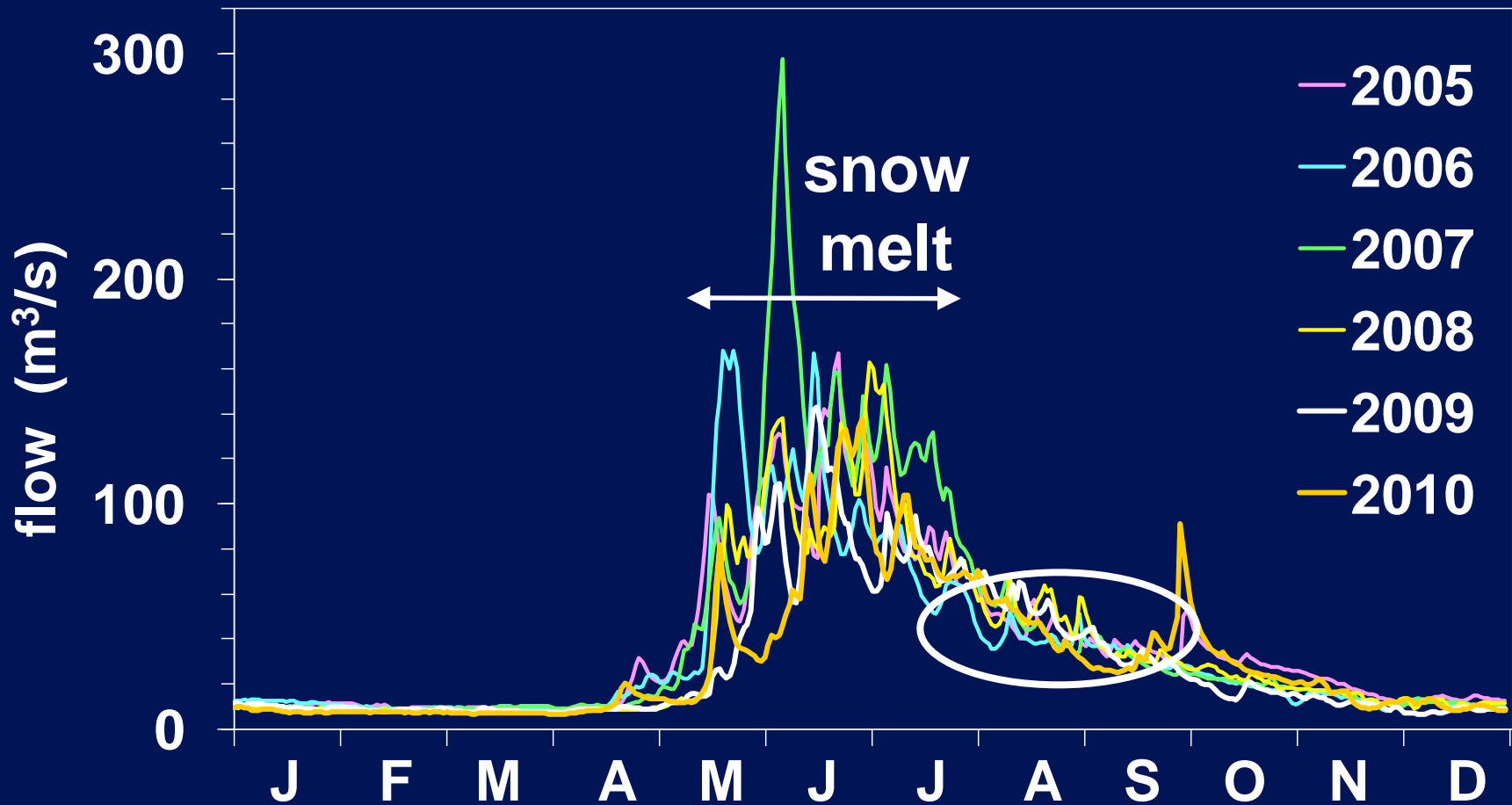
Annual Average Runoff



Hydrological Atlas of Canada (1978)

Bow River in the Canadian Rockies

Discharge at Banff (2200 km², unregulated)



Data: Water Survey of Canada

Lake O'Hara Hydrological Observatory (14 km²)

Elevation: 2000-3500 m

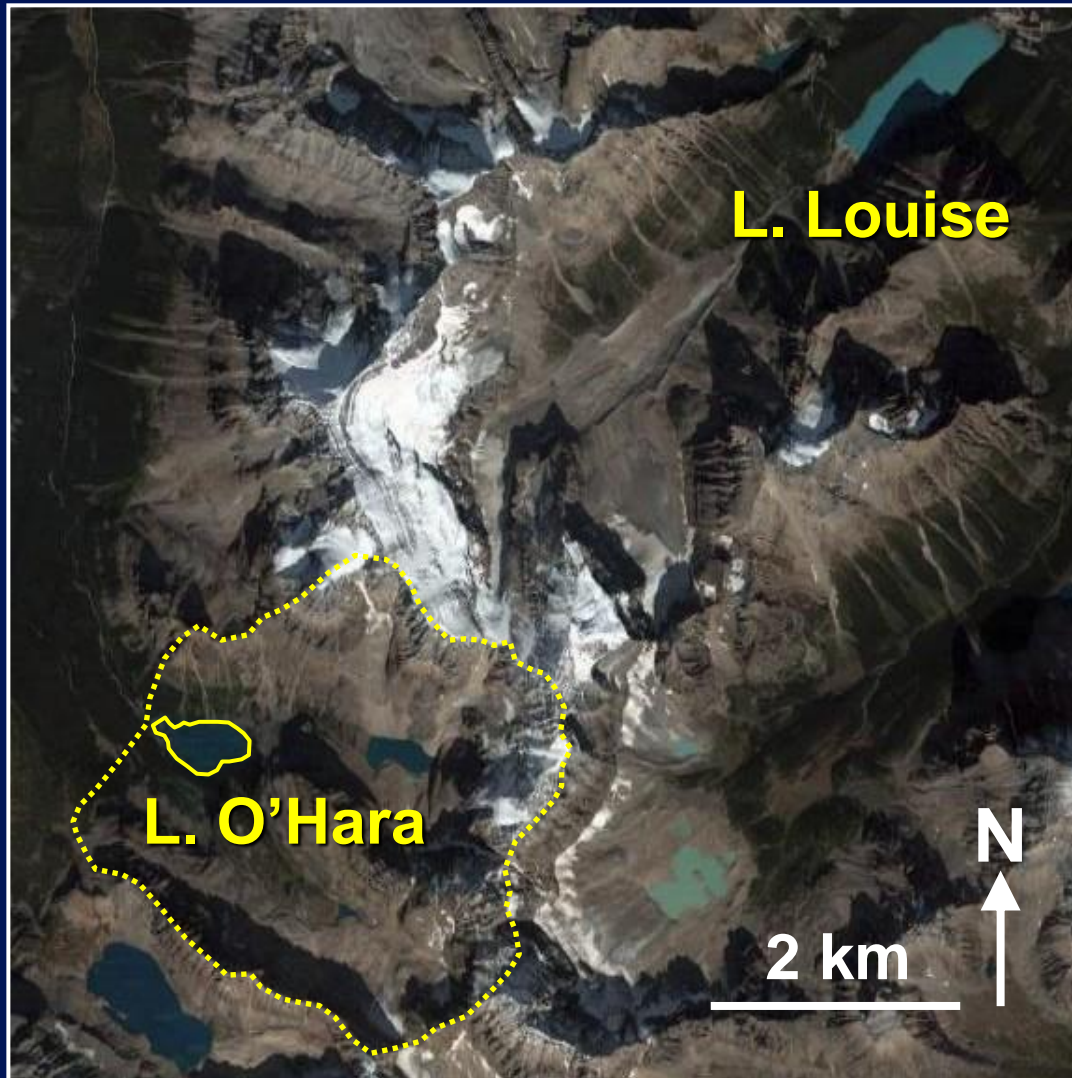
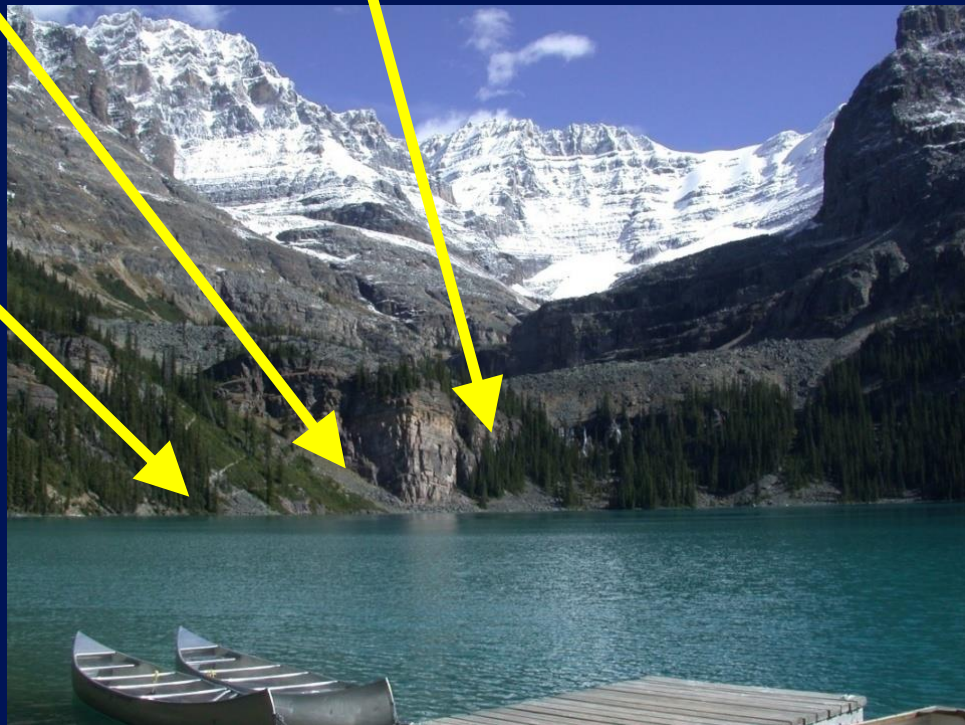


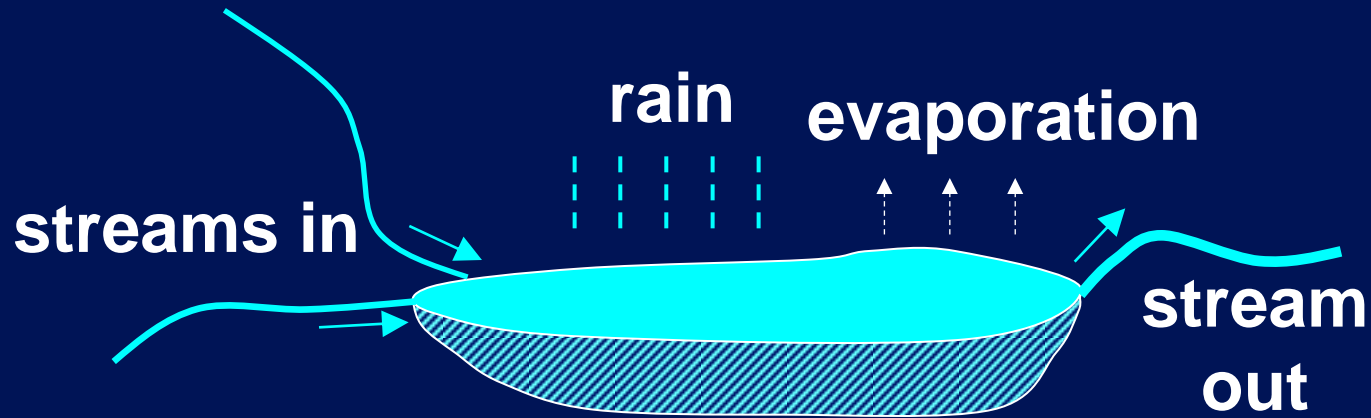
Image from Google Earth

- Weather stations
- Water level gauges
- Stream flow gauges
- Other instruments



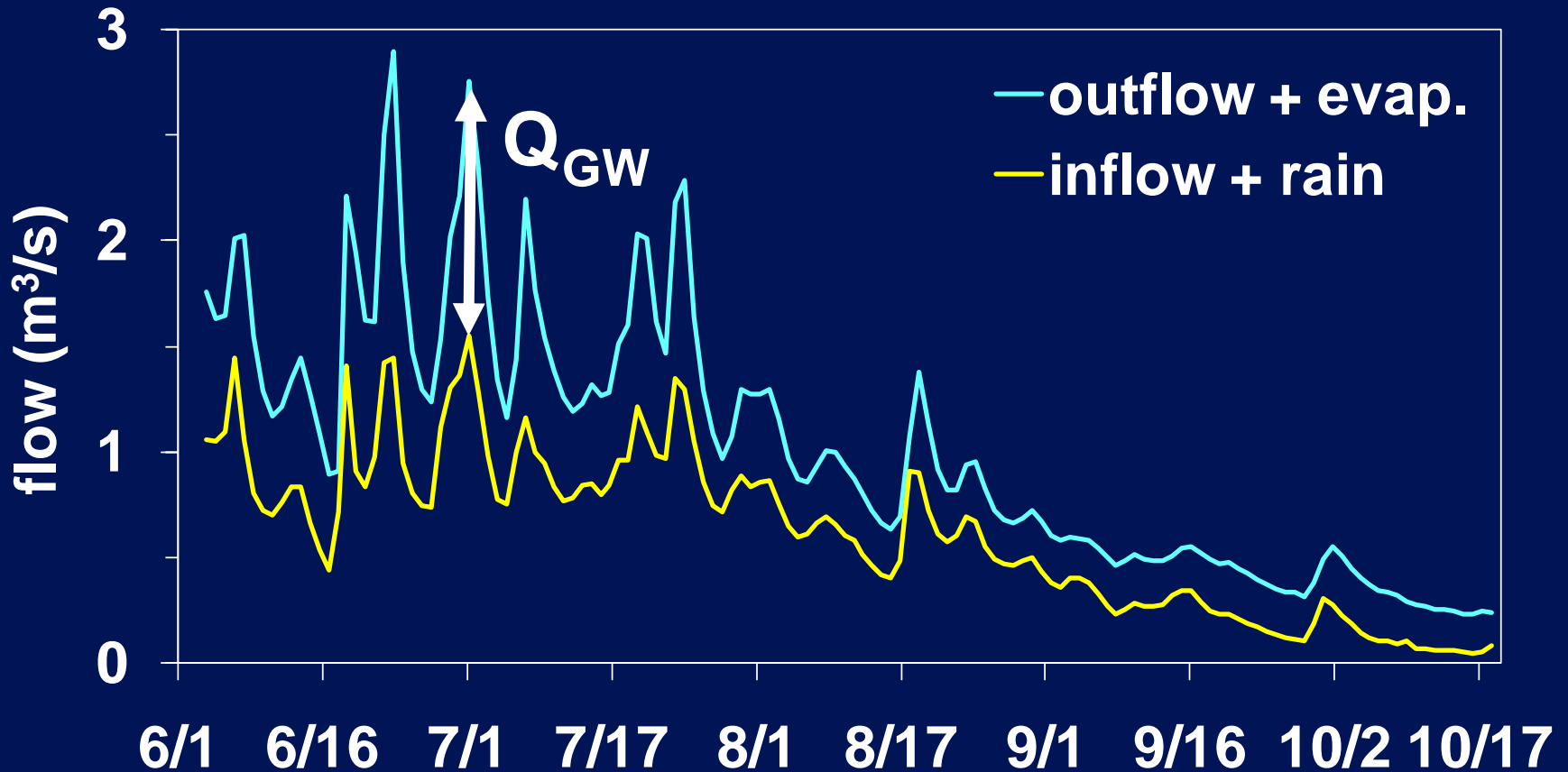


Lake Water Balance



If there is no groundwater flowing in or out,
streams in + rain = stream out + evaporation

The lake has more outflow than inflow.



Q_{GW} = groundwater inflow – outflow

30-70 % of water input is by groundwater.

Hood et al. (2006, *Geophys. Res. Lett.*, 33, L13405)

Opabin Creek Sub-Watershed (5 km²)

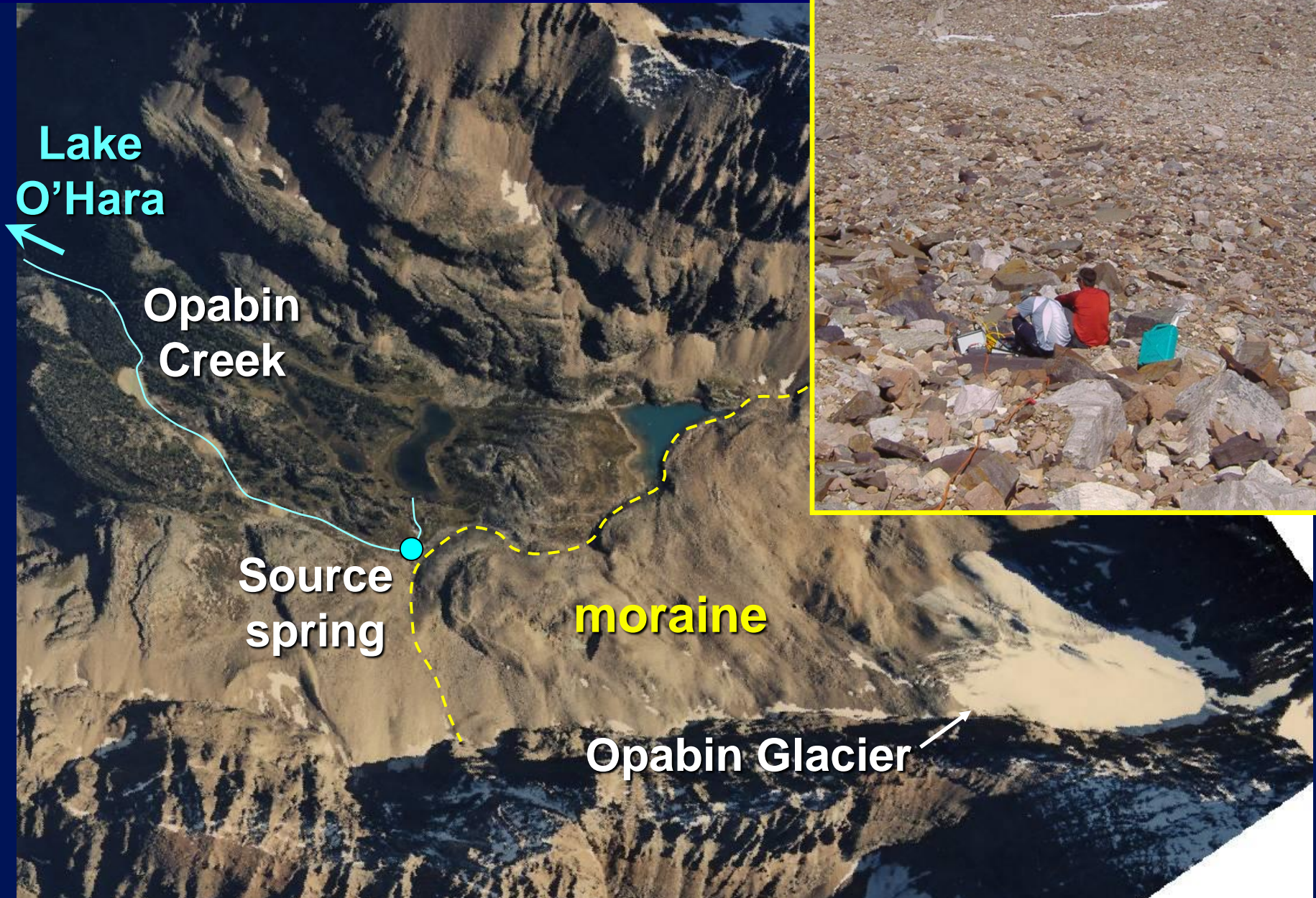
Opabin Creek
entering Lake O'Hara



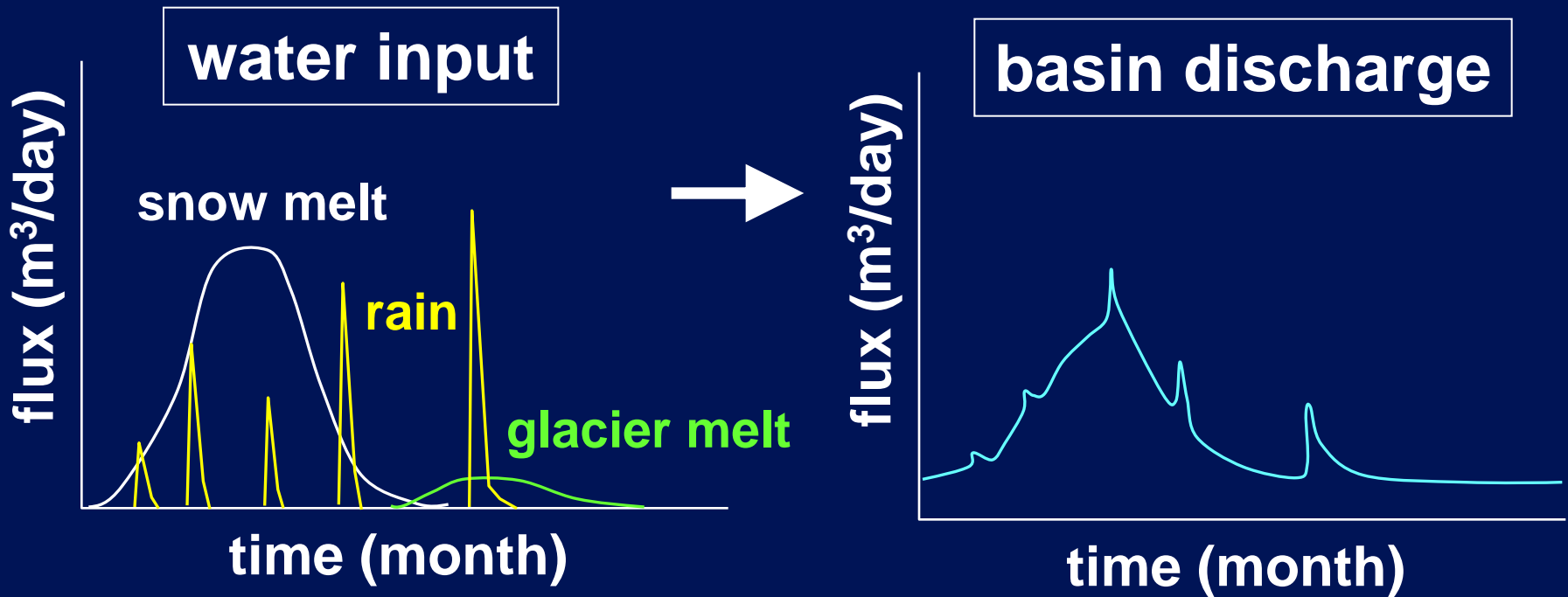
Opabin Creek source:
Groundwater spring



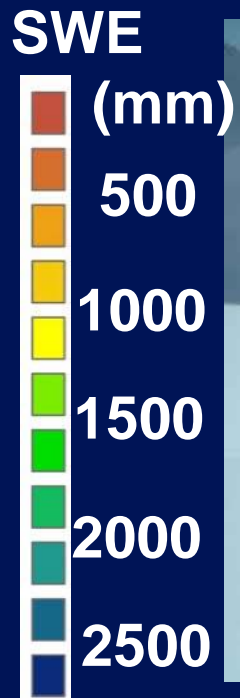
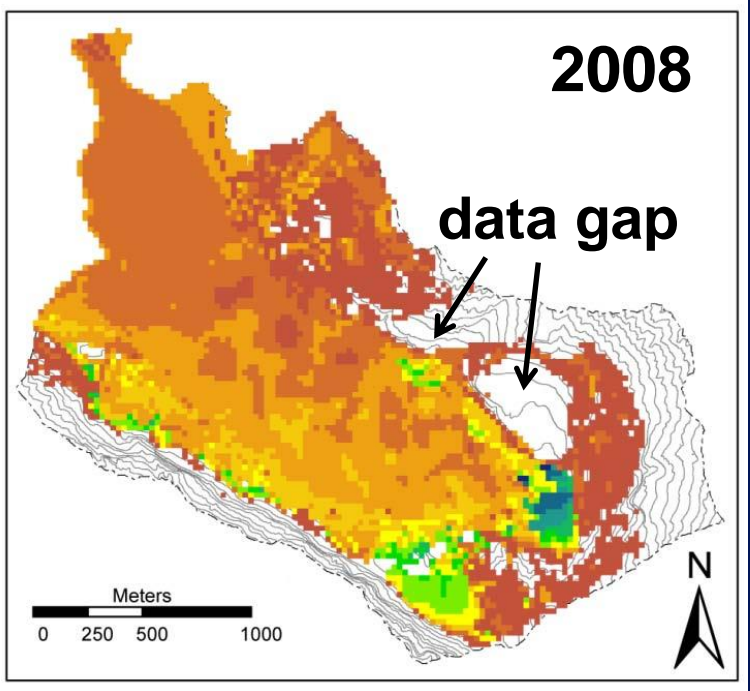
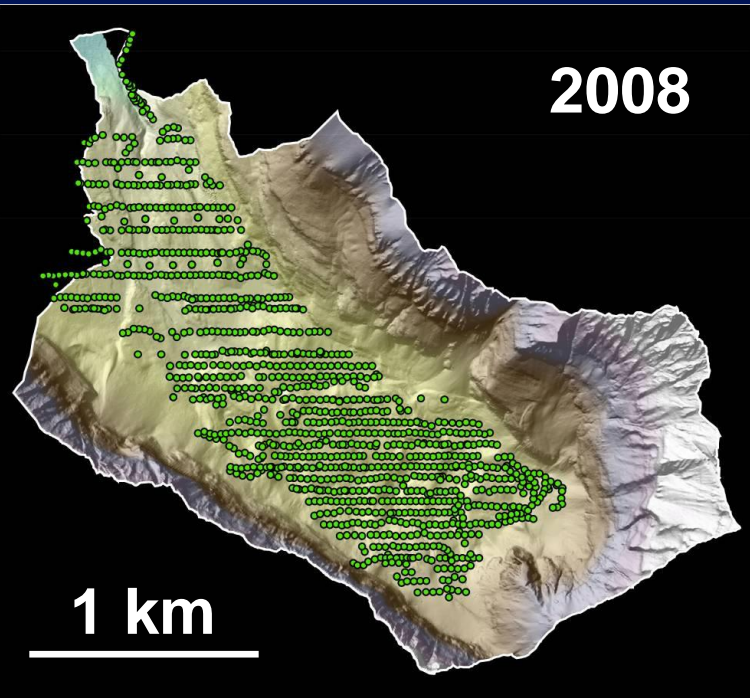
Opabin Plateau



Water Input and Output in the Opabin Watershed



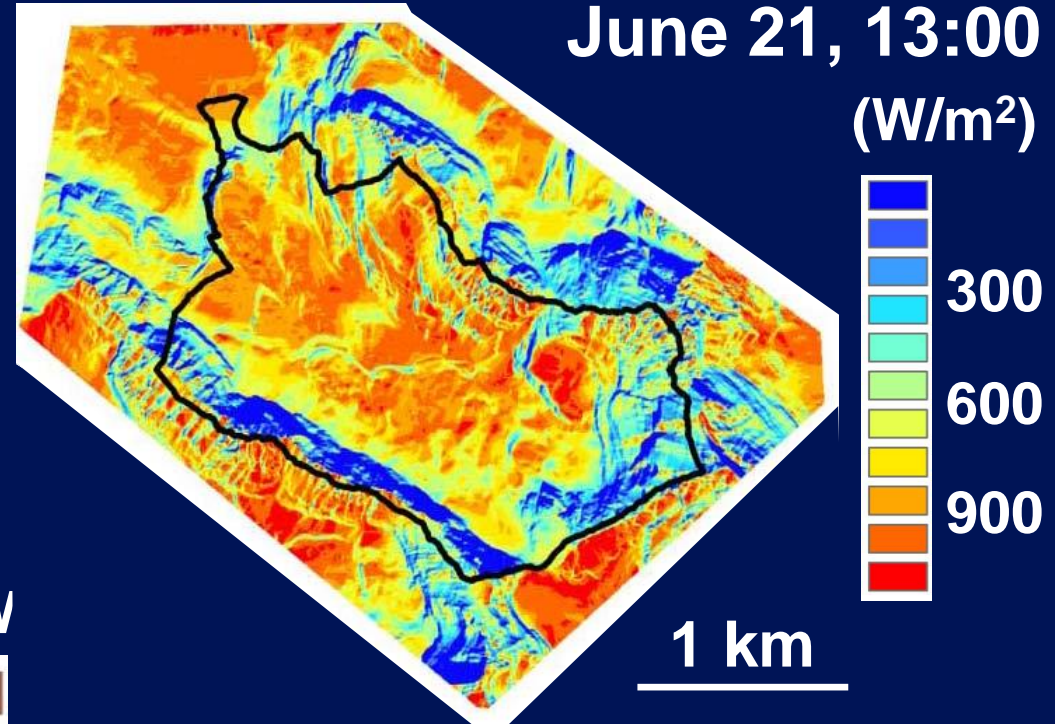
Groundwater provides temporary storage.





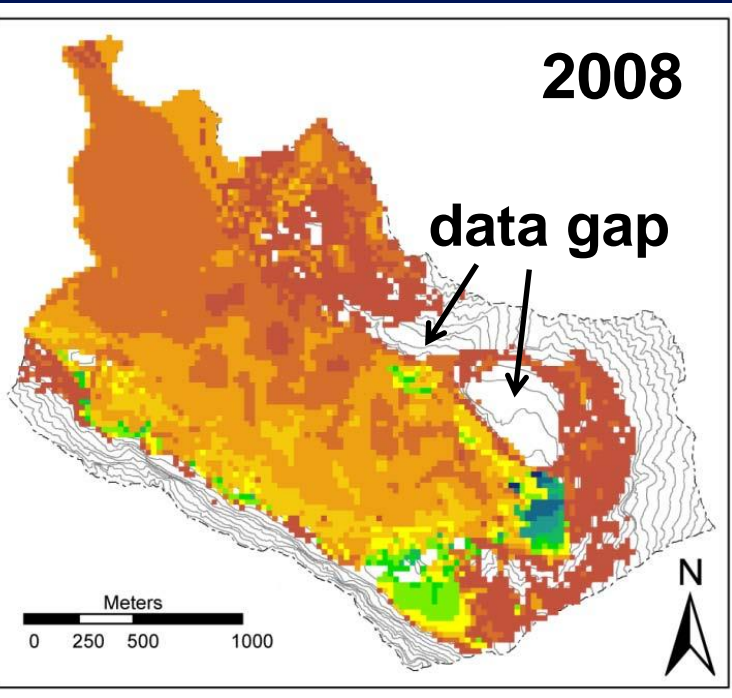
Shortwave radiation

June 21, 13:00

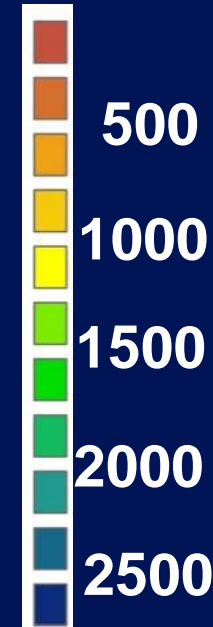


2008

data gap



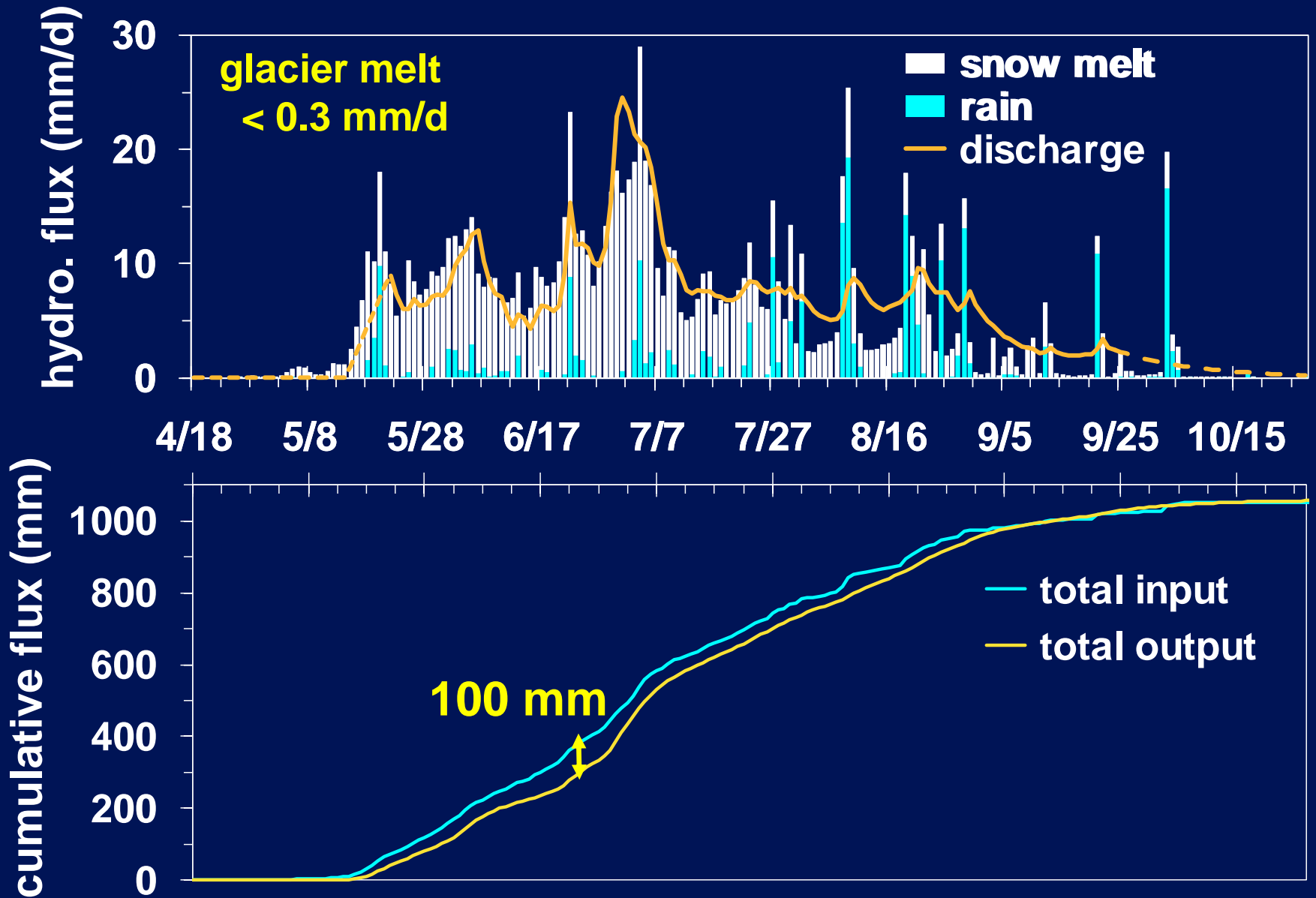
SV



Snowmelt model

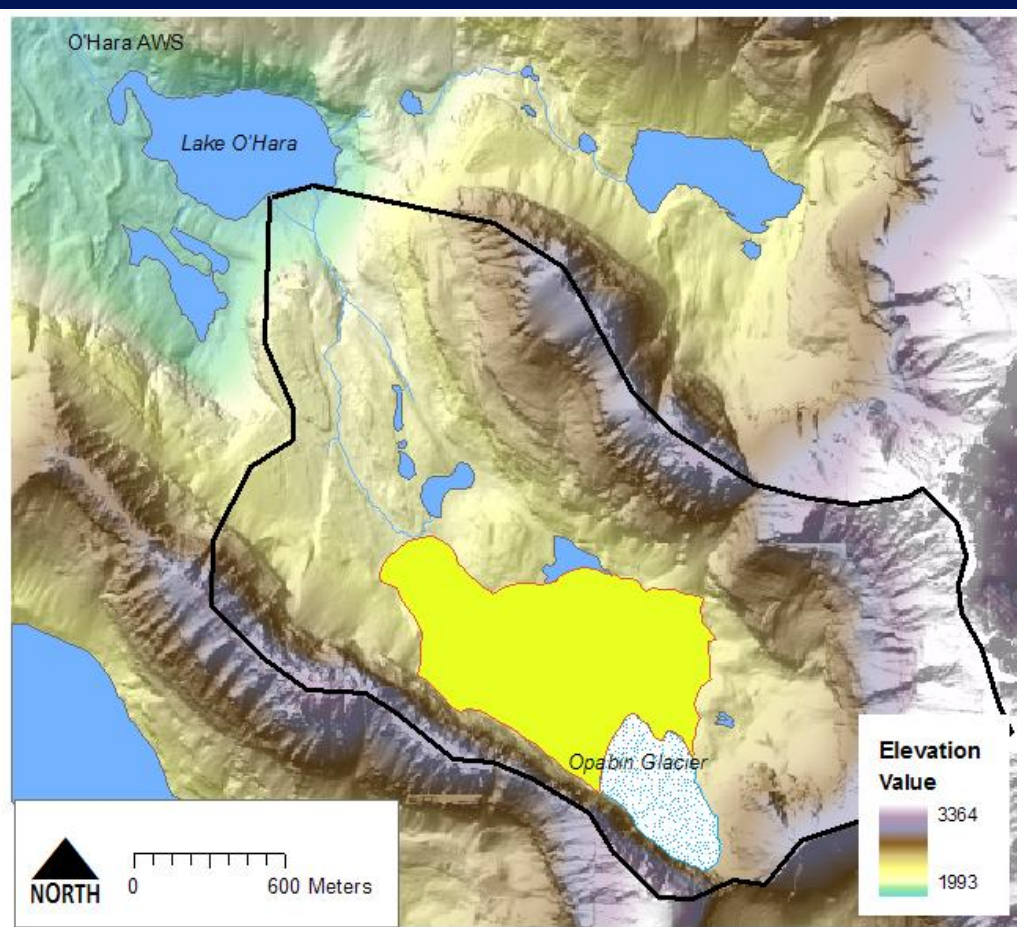
- Utah Energy Balance
- Hourly time step
- 50 m grid cells

Opabin Basin Water Balance (2008)



Hydrogeological Response Units

- Bedrock (hard quartzite)
- Proglacial moraine

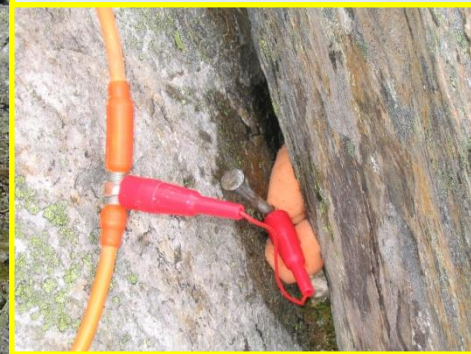




seismic refraction

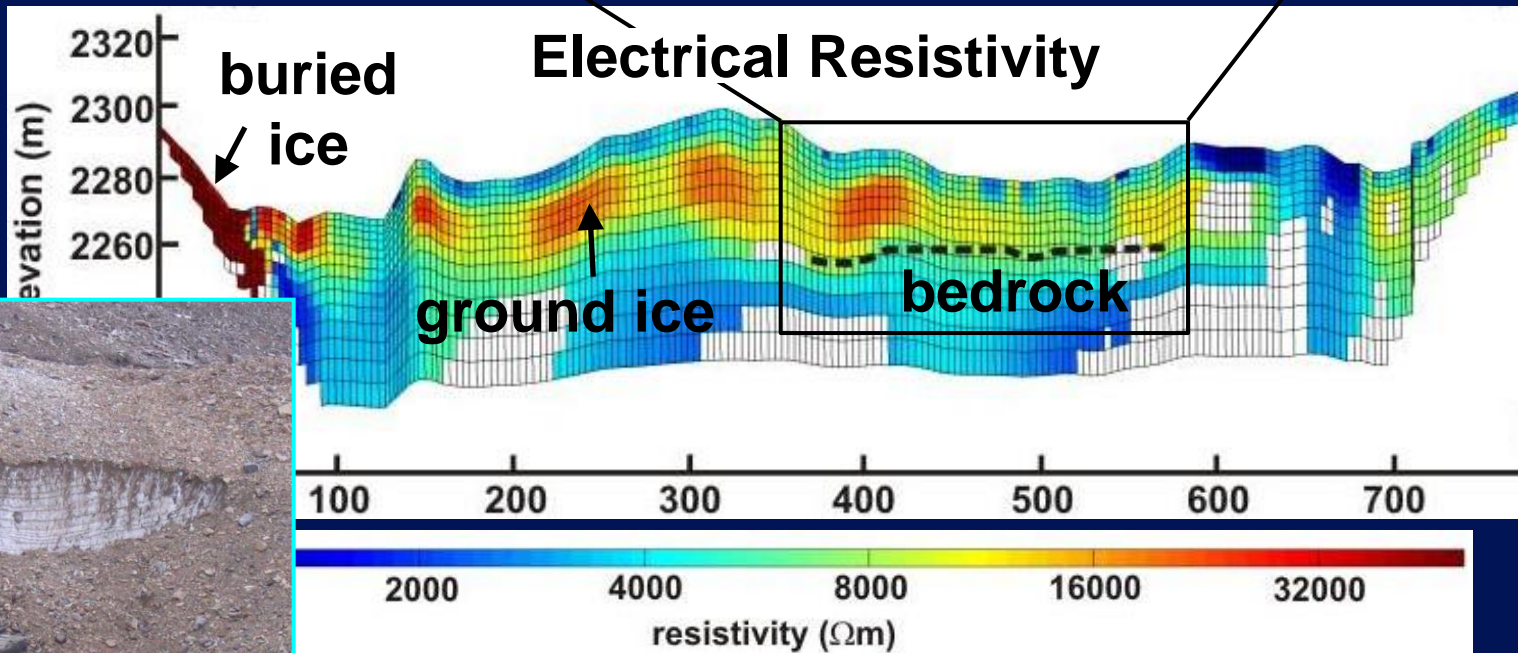
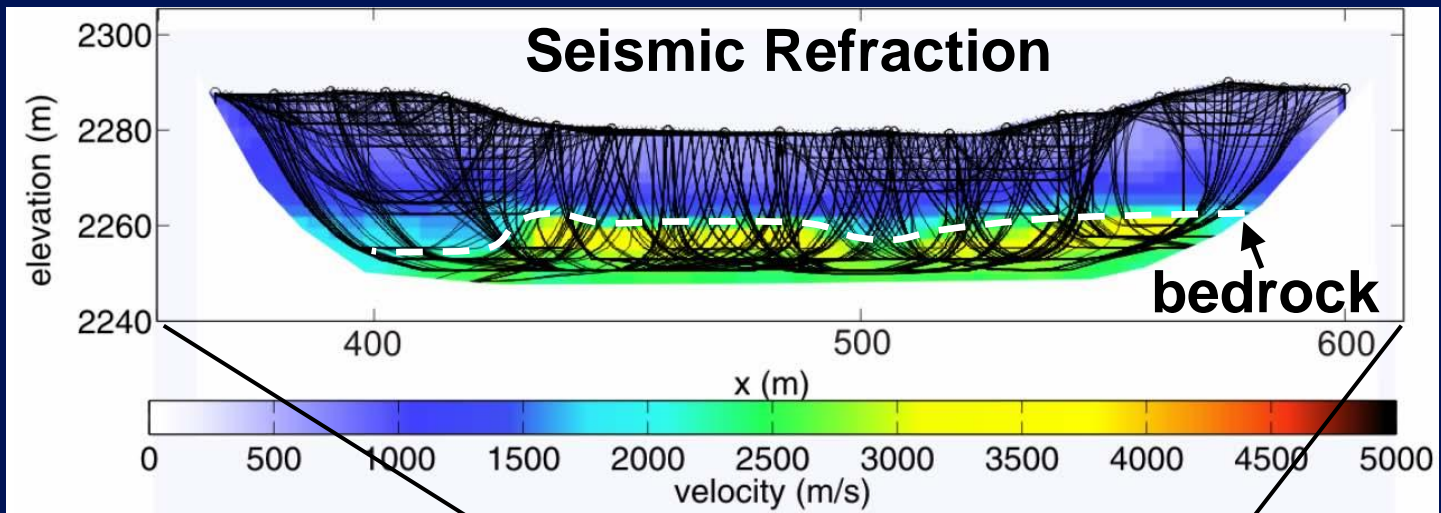


electrical resistivity



**Opabin
Glacier**

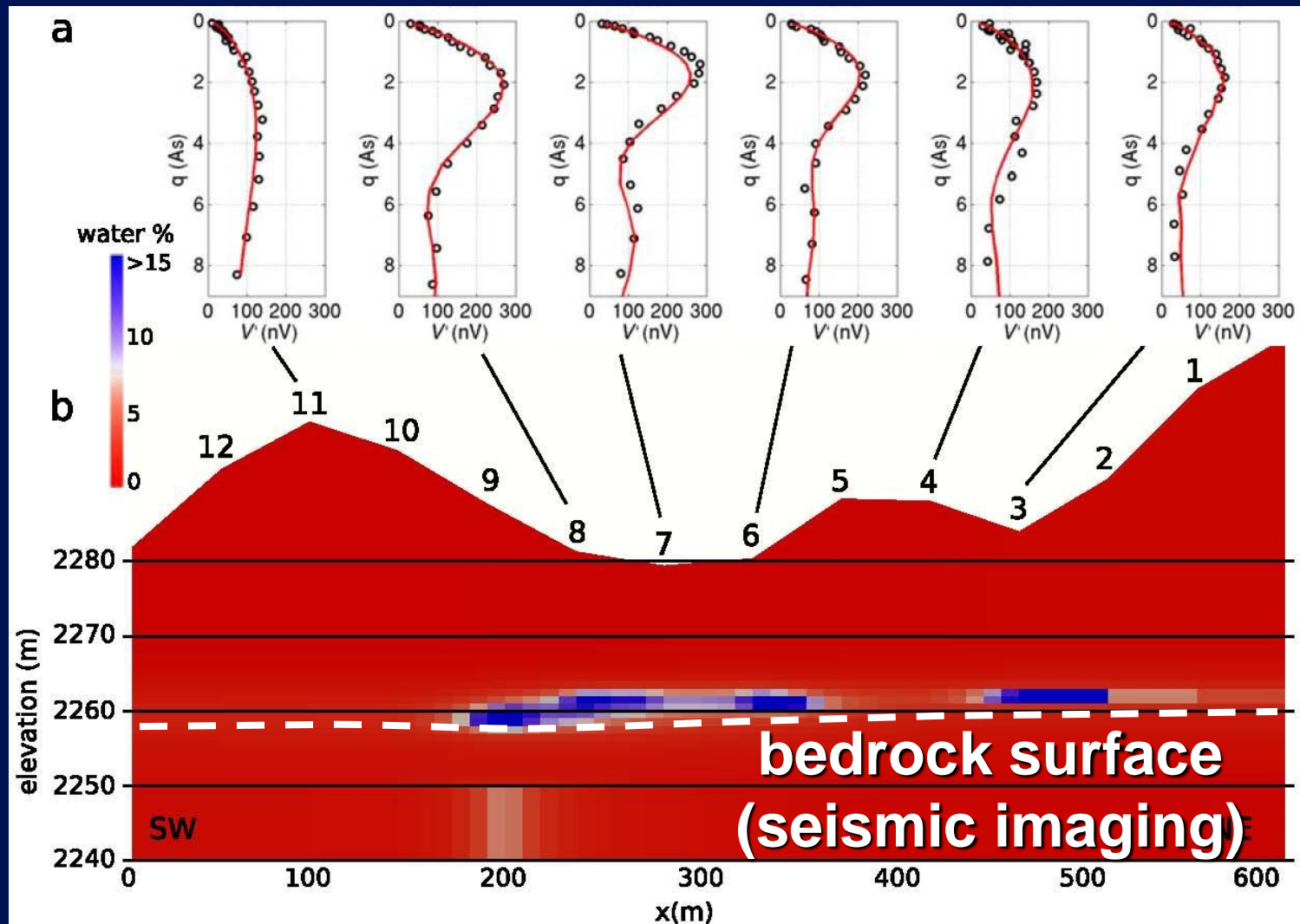




Langston et al. (2011, *Hydrol. Process.* 25: 2967)

Nuclear Magnetic Resonance Imaging

Blue colour indicates water molecules



Lehmann et al. (2012. *Geophysics*, 76: B165)

Ground-Penetrating Radar



**Opabin
Glacier**

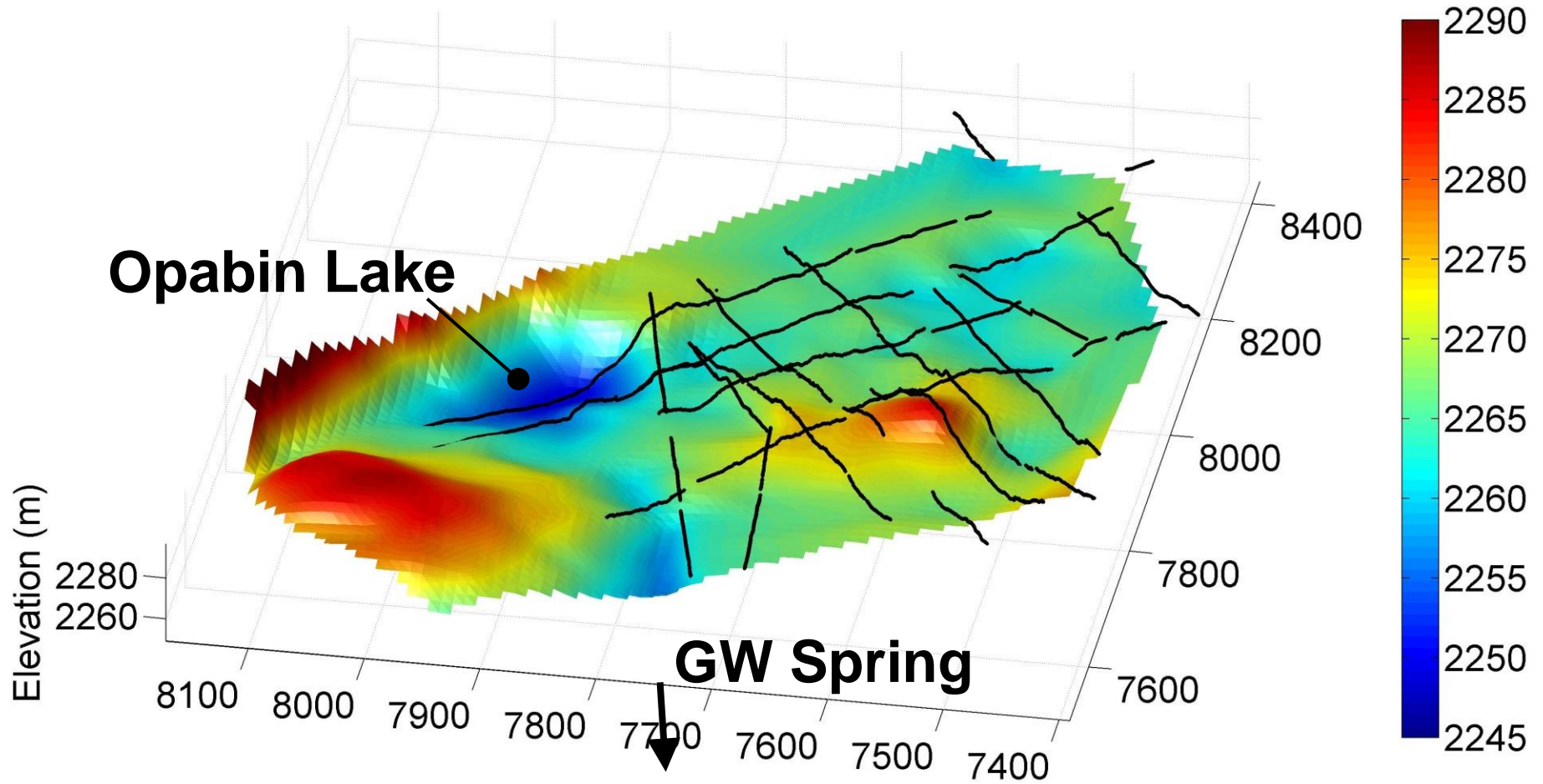
Opabin Lake



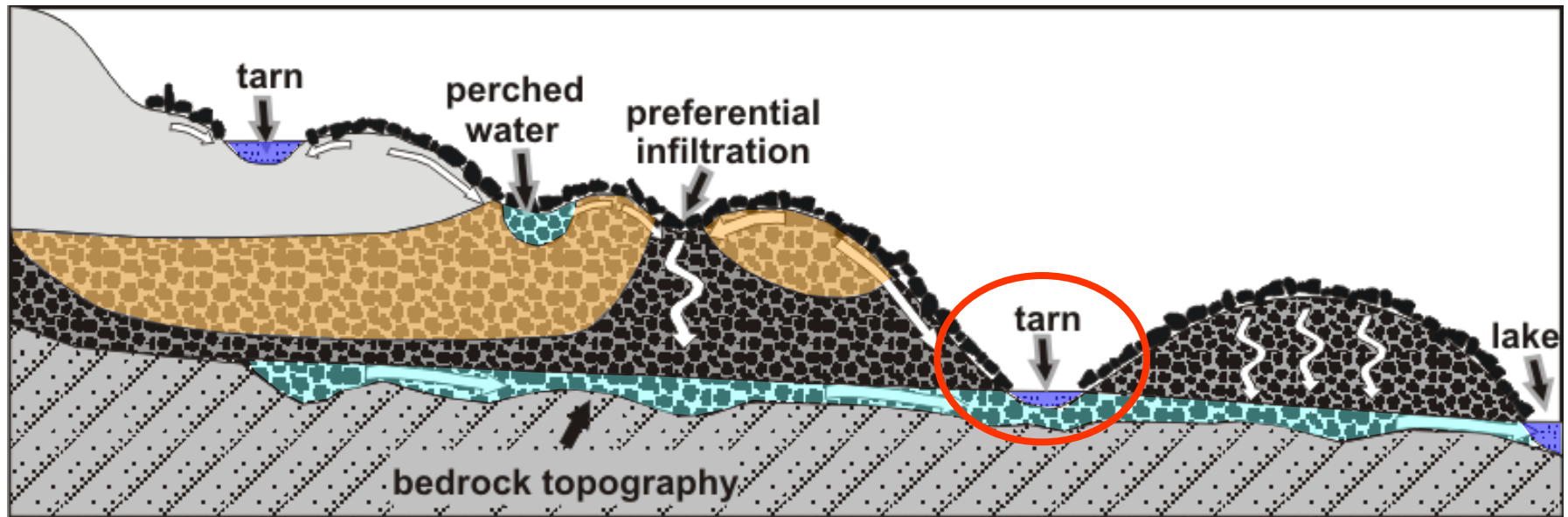
Spring









Bedrock Surface Map from Radar Data



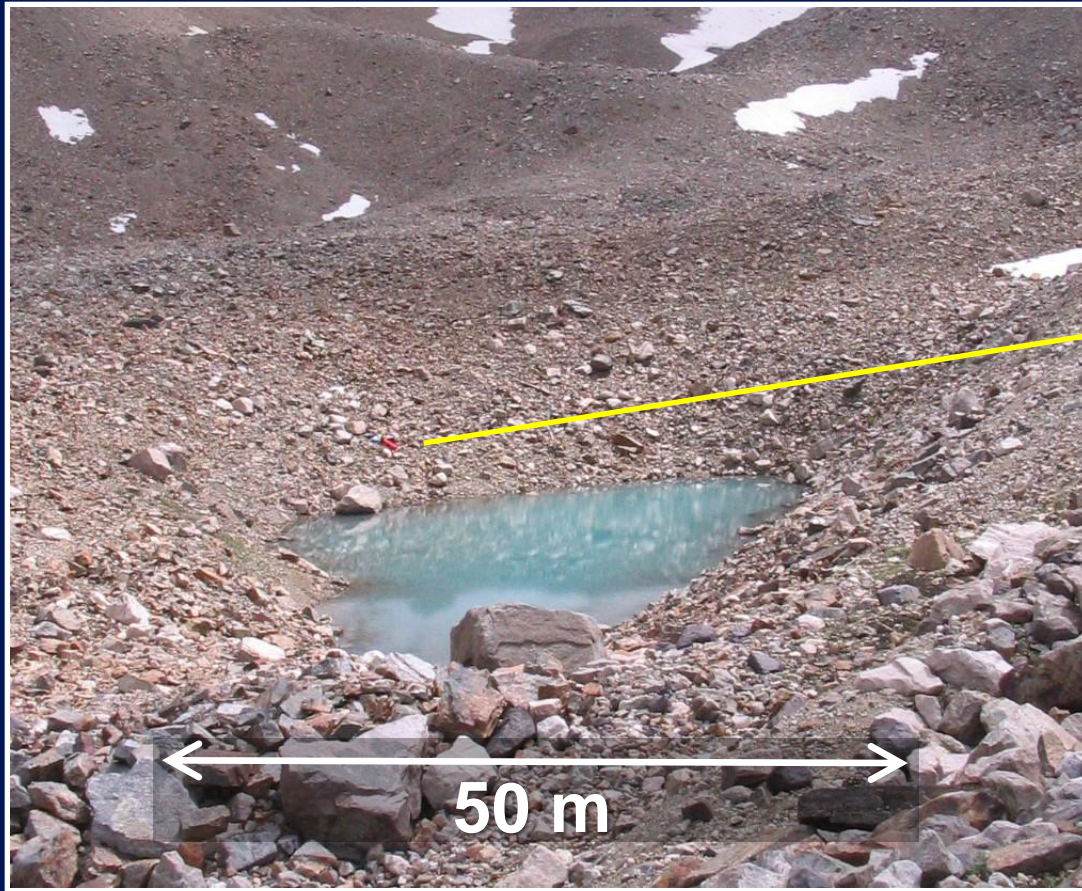
Emerging Conceptual Model



- | | | | |
|--|----------------------|---|----------------------------|
|  | Dry Moraine Material |  | Saturated Moraine Material |
|  | Massive Ice |  | Bedrock |
|  | Permafrost |  | Tarn or Lake |

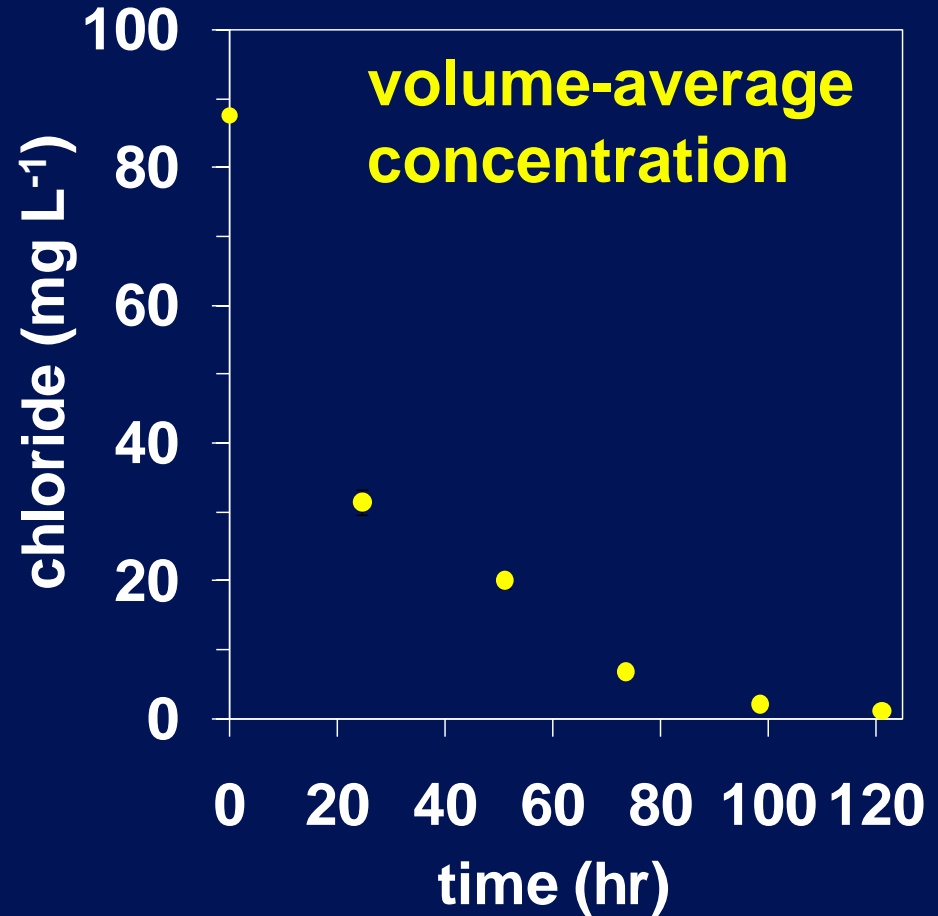
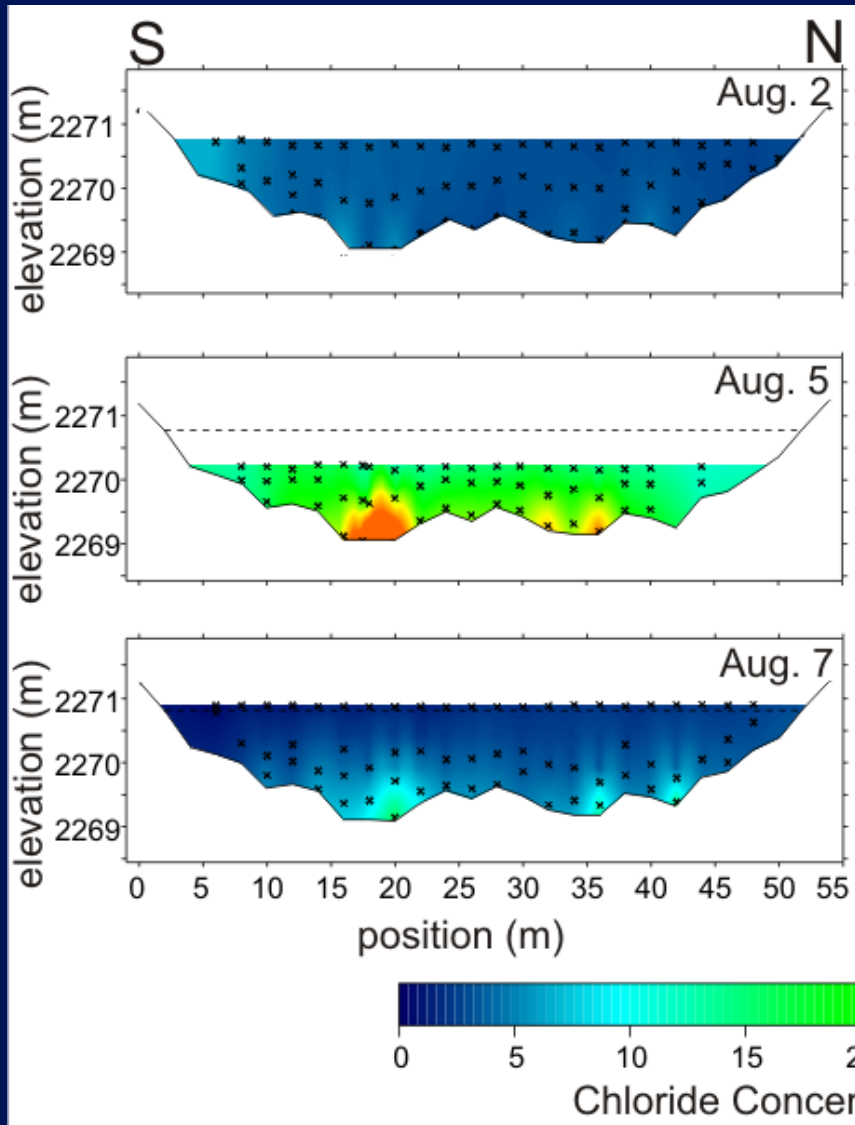
Langston et al. (2011, *Hydrol. Process.* 25: 2967)

Estimating Hydraulic Conductivity of the Moraine



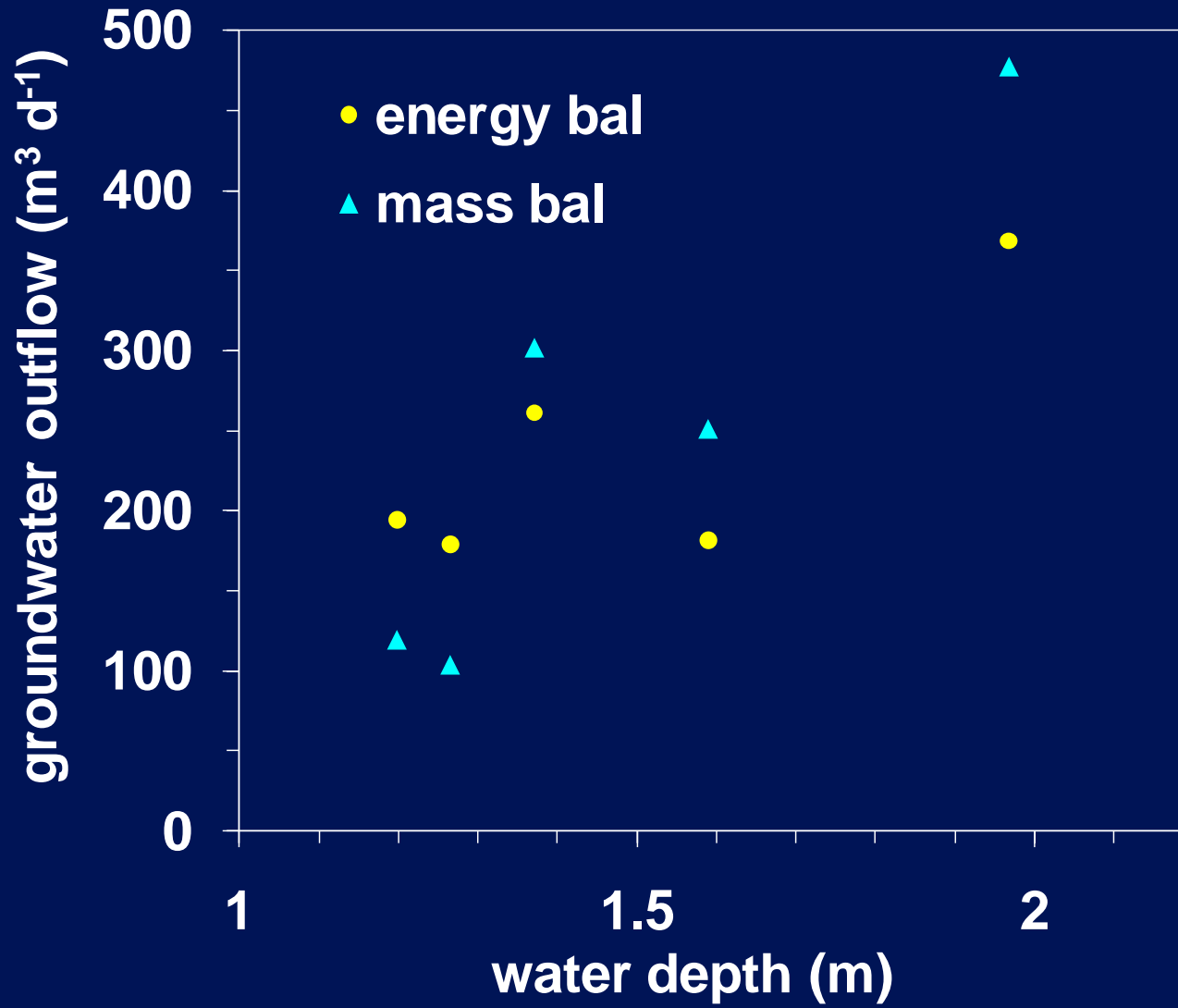
- Water balance equation
- Energy balance equation
- Tracer mass balance equation

Chloride in the Tarn



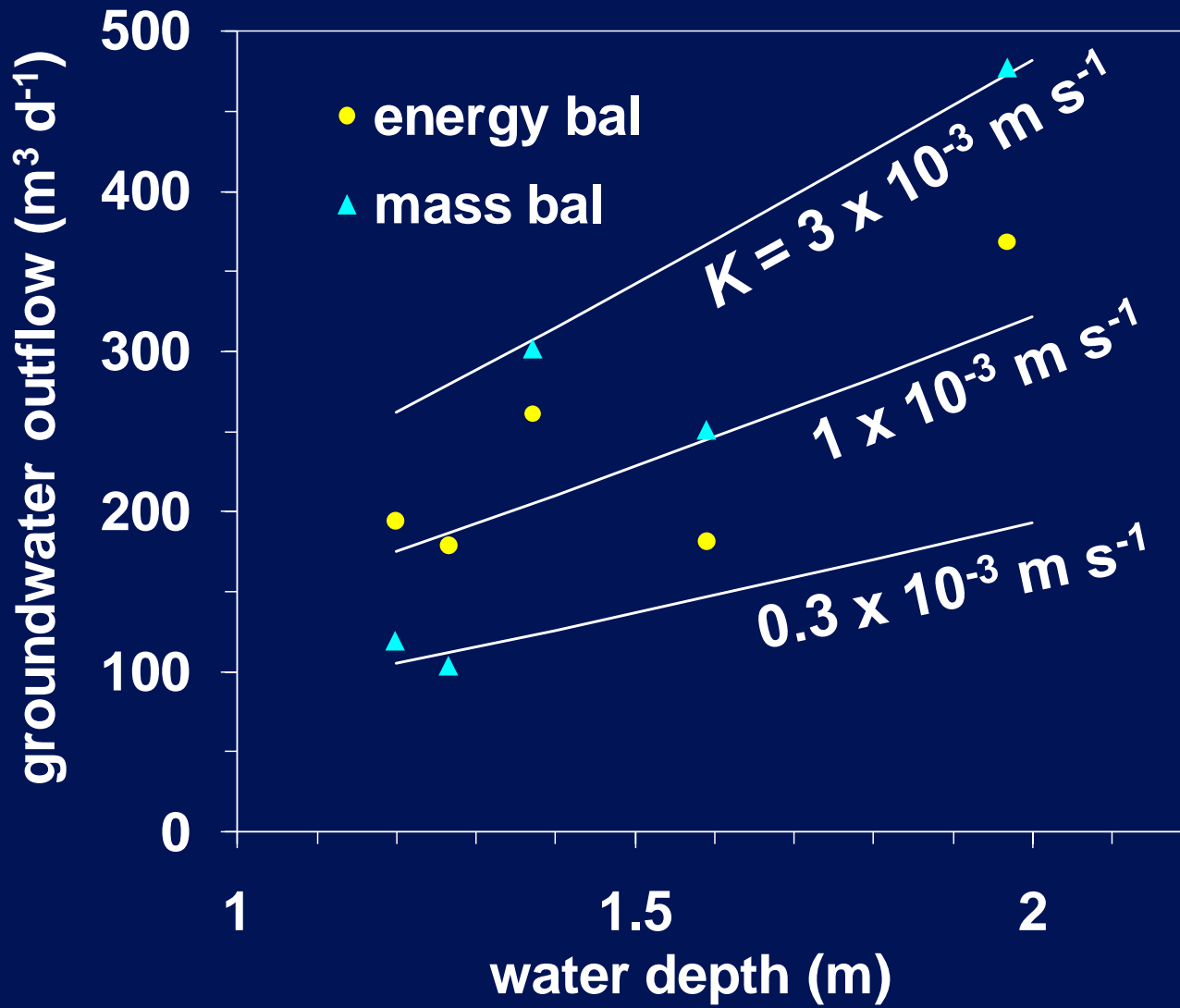
Langston et al. (2013, *Water Resour. Res.* 49: 5411)

Groundwater outflow from the tarn



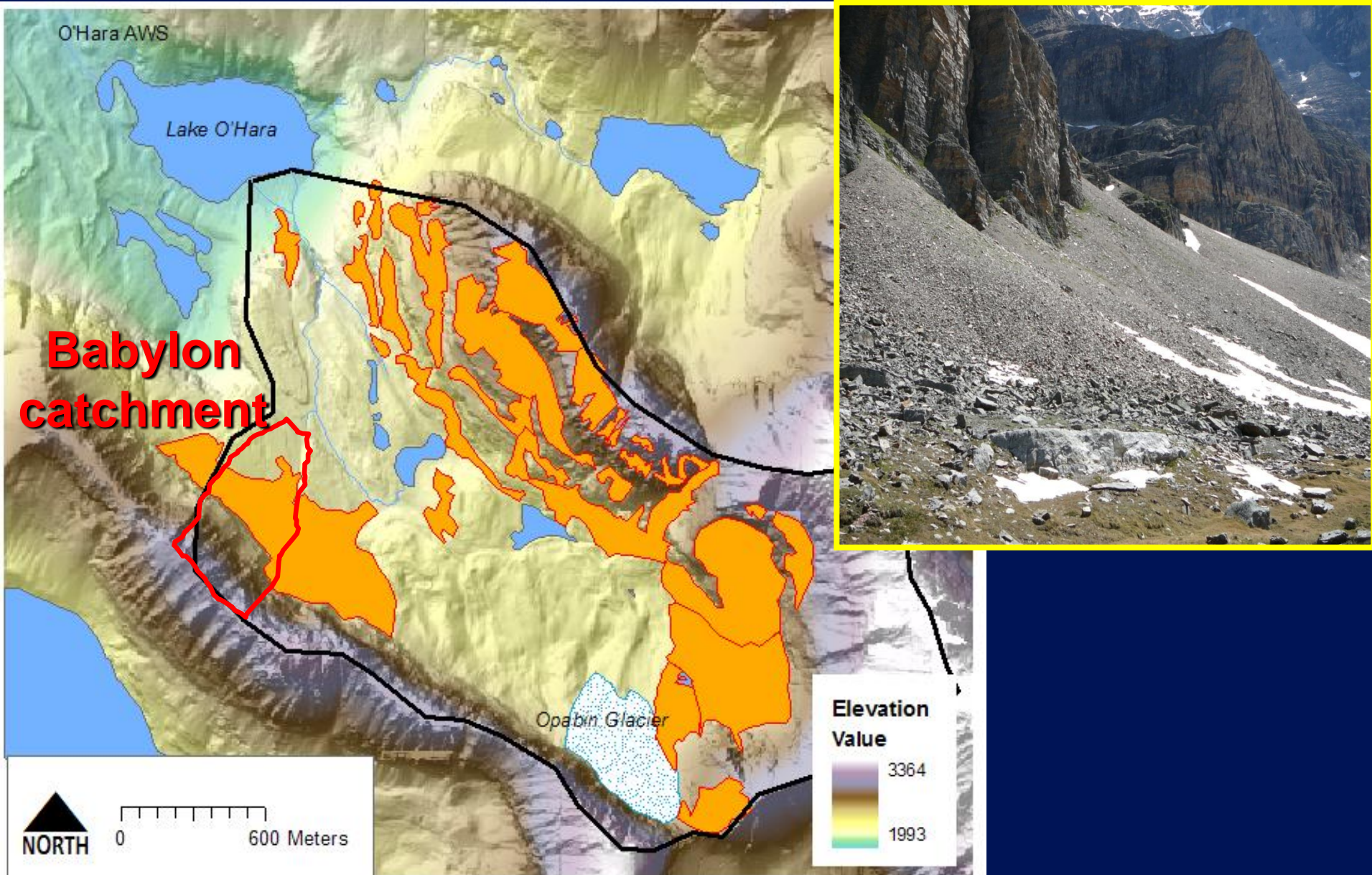
Langston et al. (2013, *Water Resour. Res.* 49: 5411)

Groundwater outflow from the tarn



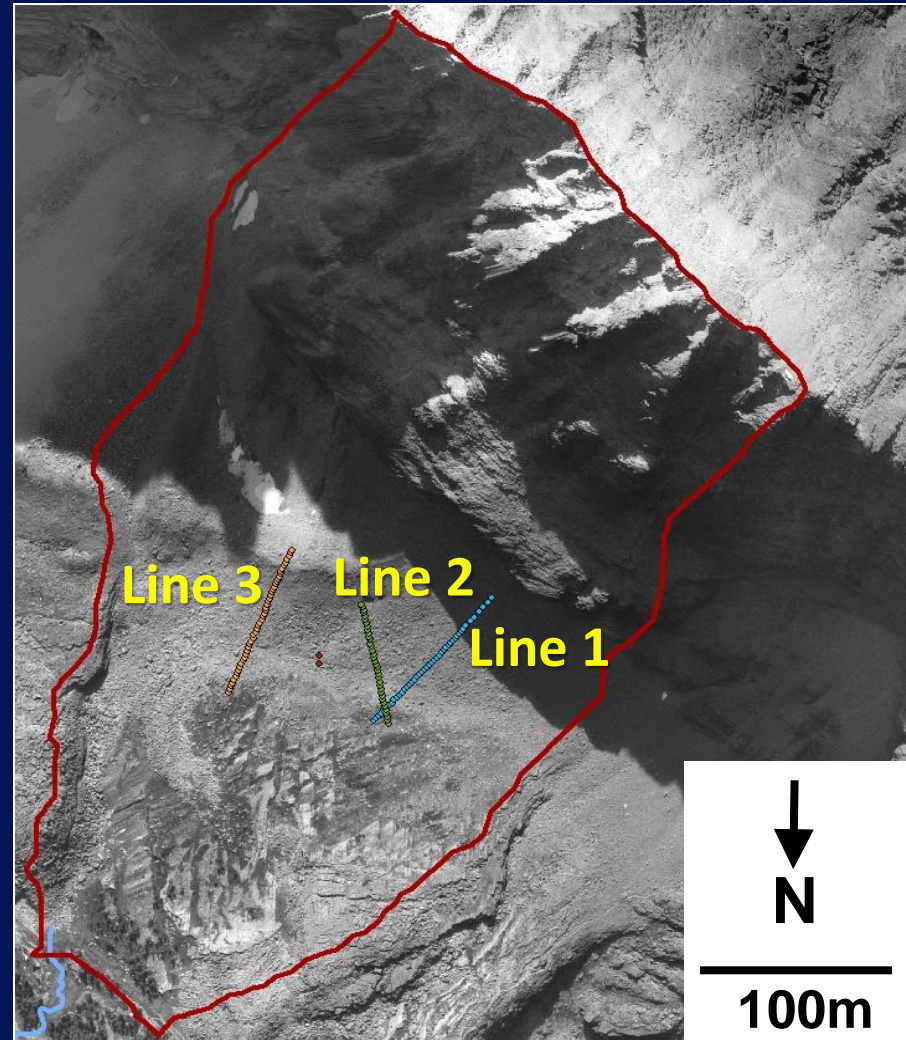
Langston et al. (2013, *Water Resour. Res.* 49: 5411)

Hydrogeological Response Unit: Talus

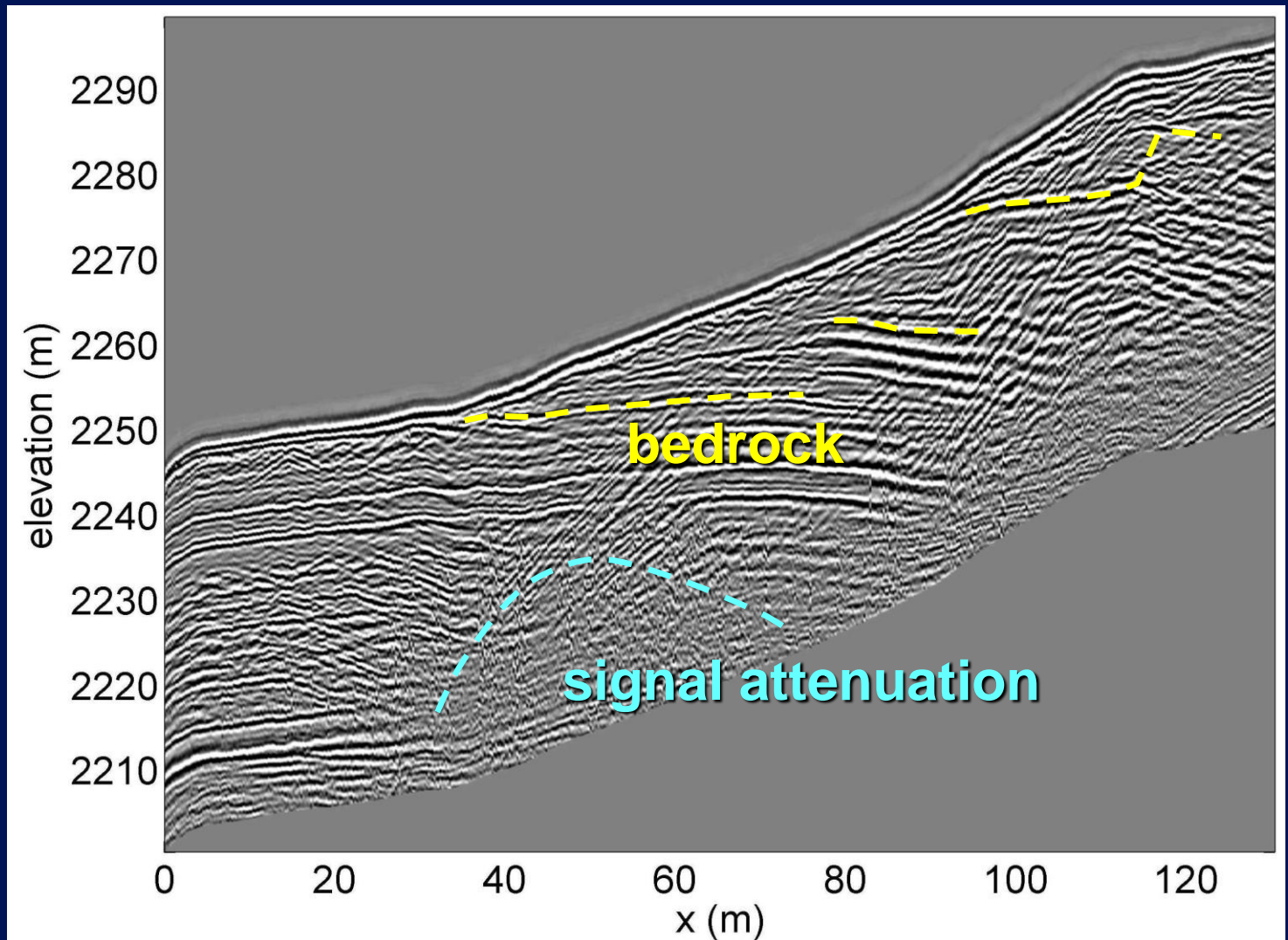


Geophysical Survey of Talus

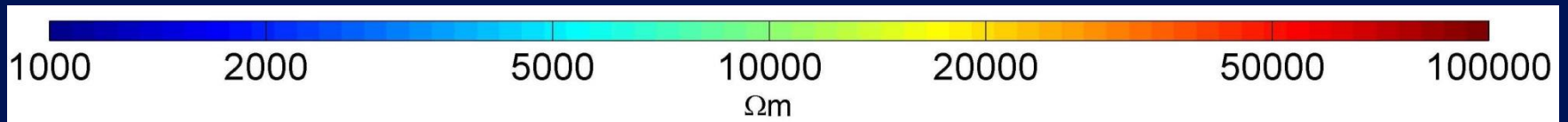
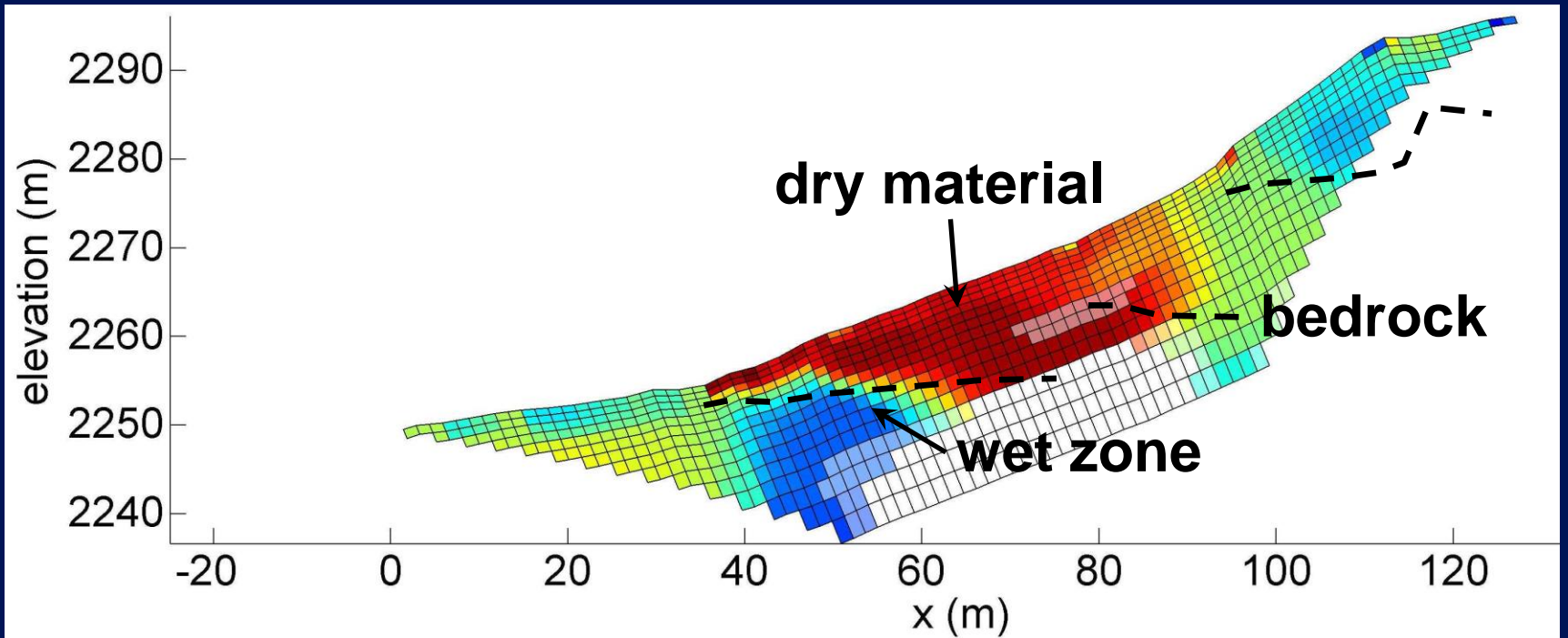
Resistivity Imaging and Ground Penetrating Radar (GPR)



Line 3: Ground Penetrating Radar

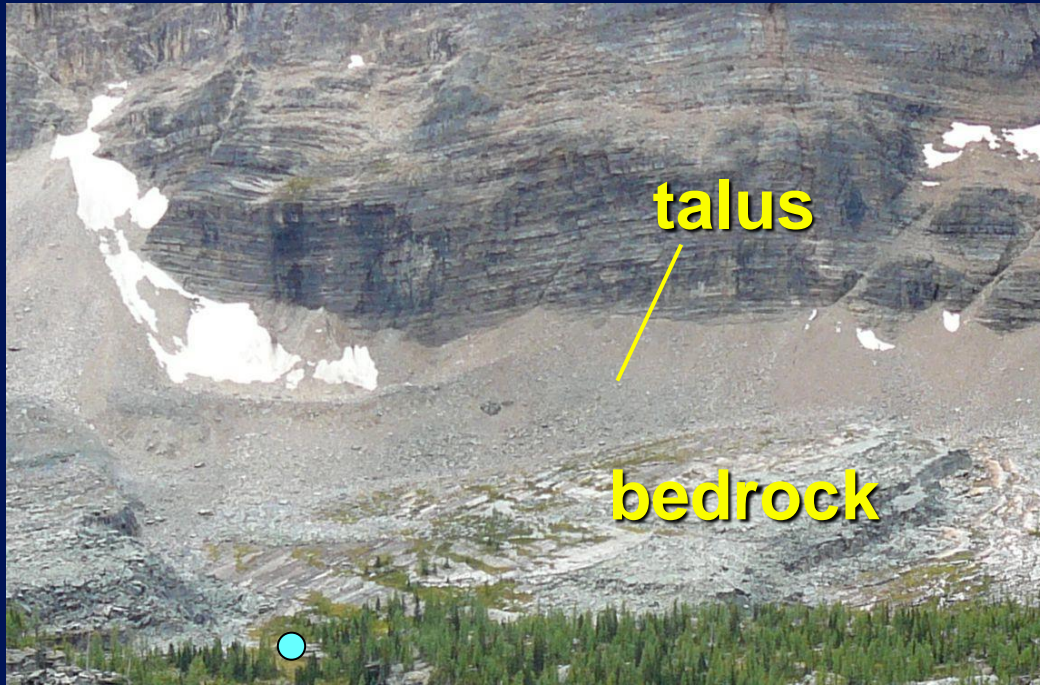


Line 3: Electrical Resistivity



Thickness of low resistivity zone below high resistivity cannot be determined (Hilbich et al., 2009).

Measurement of Talus Discharge



Gauging Station



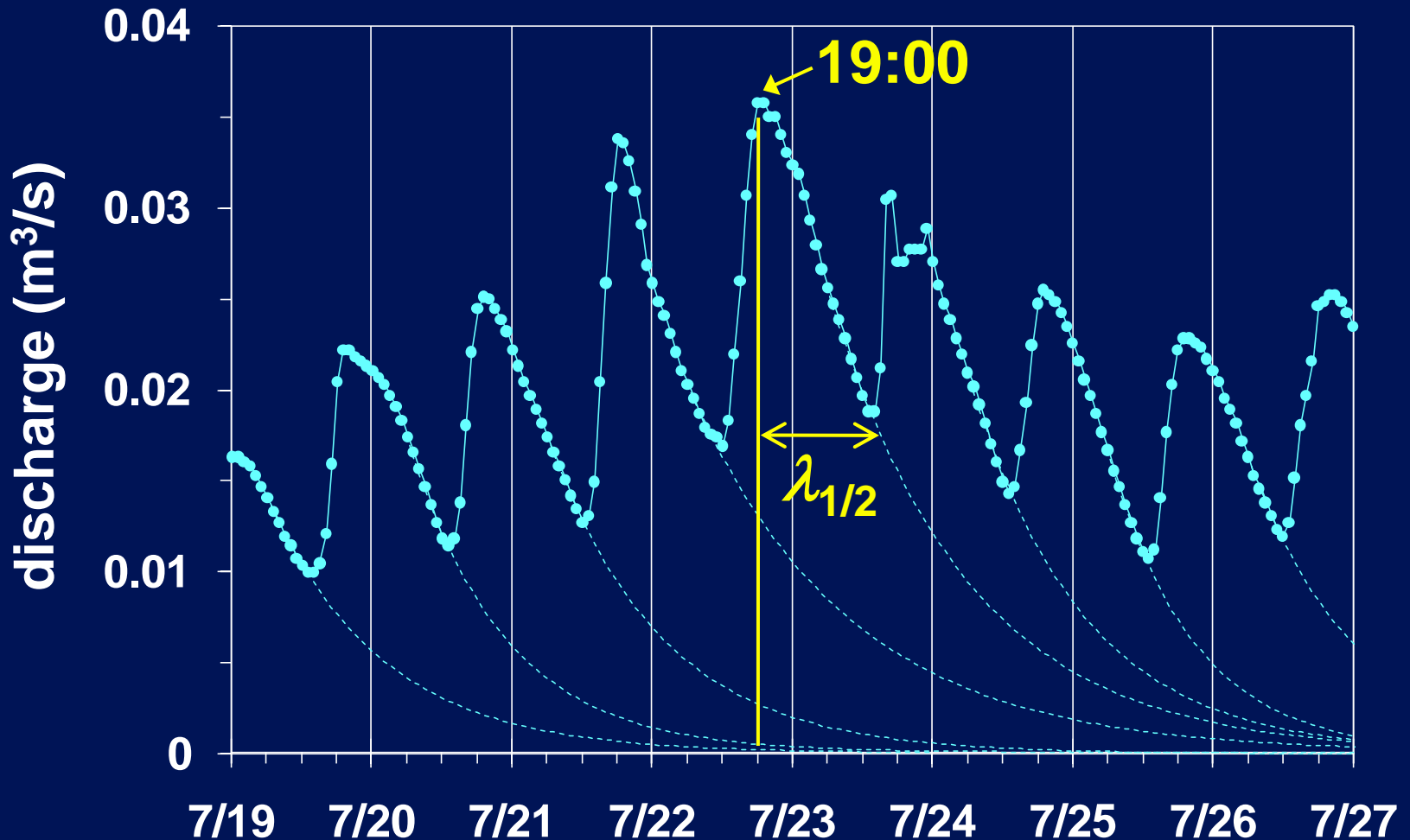
Babylon Creek



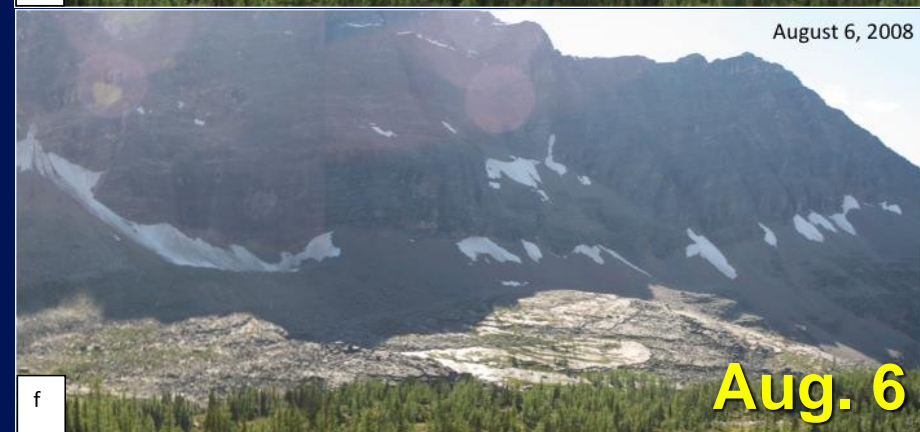
Babylon Creek Discharge, 2008

Diurnal fluctuations, peaking in early evening.

Half life ($\lambda_{1/2}$) of exponential decay < 1 day.

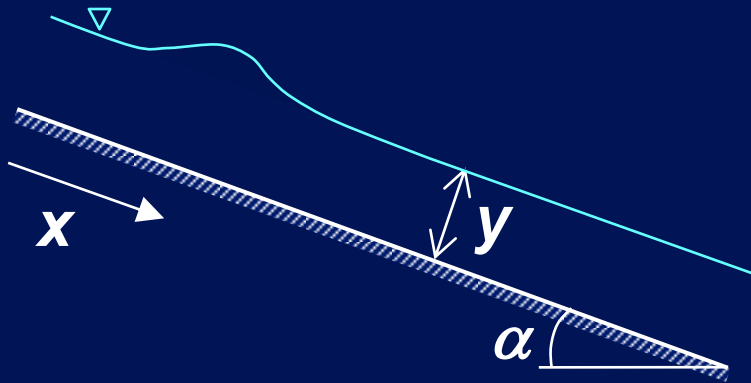


Sequential Photographs, 2008



Hillslope Flow in Unconfined Aquifers

Brutsaert (2005) Approximation: Kinematic Wave



$$Q = Ky \left(-\frac{dy}{dx} + \sin \alpha \right) \cong Ky \sin \alpha$$

Q : flow per unit width (m^2/s)

K : hydraulic conductivity (m/s)

The pulse of water table travels like a wave.

$$c = K \sin \alpha / n_e$$

c : velocity of wave propagation (m/s)

n_e : drainable porosity

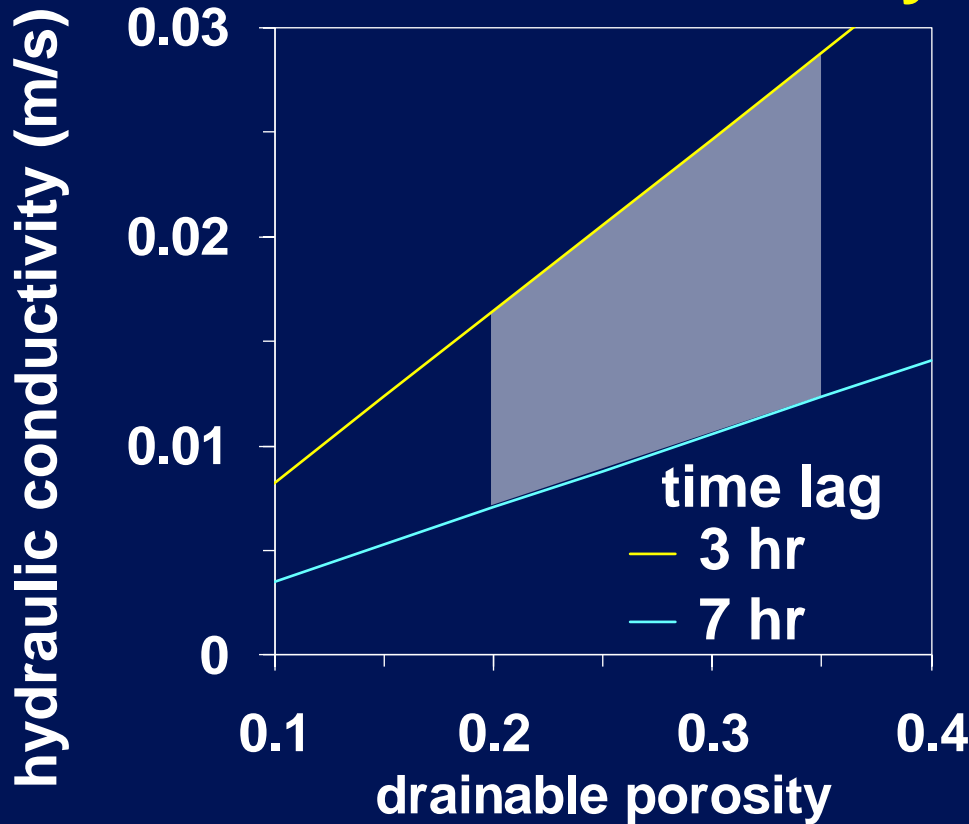
Muir et al. (2011, *Hydrol. Process.*, 25: 2954)

Analysis of Babylon Hydrograph

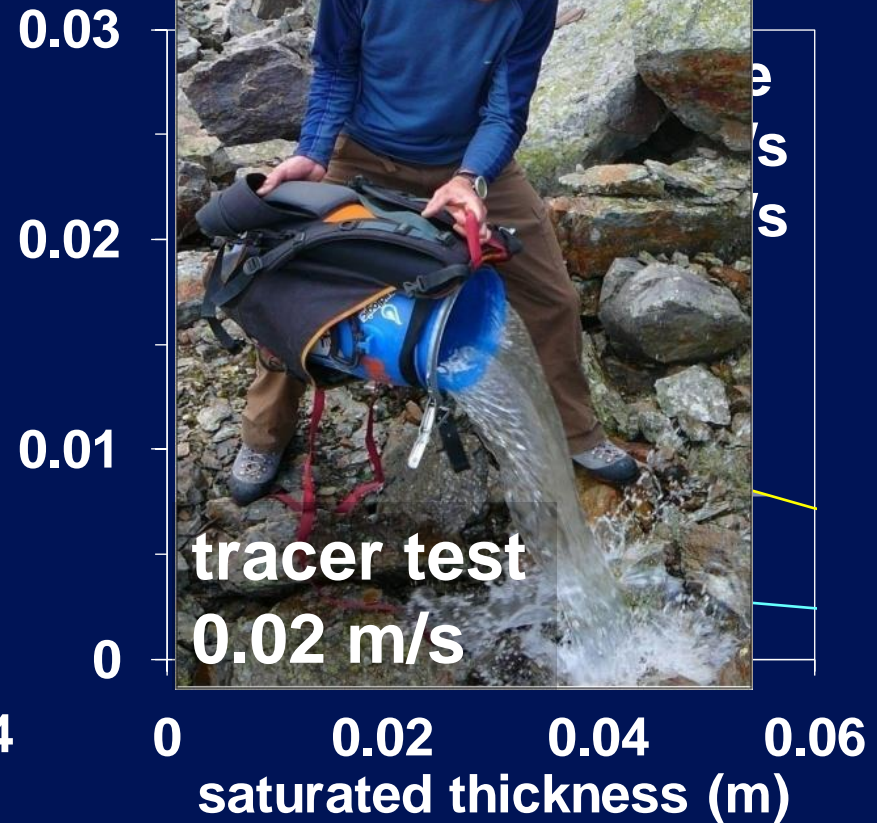
Peak discharge – peak snowmelt = 3 to 7 hrs

Flow rate ranged between 0.01 and 0.03 m³/s

From wave velocity



Flow rate equation

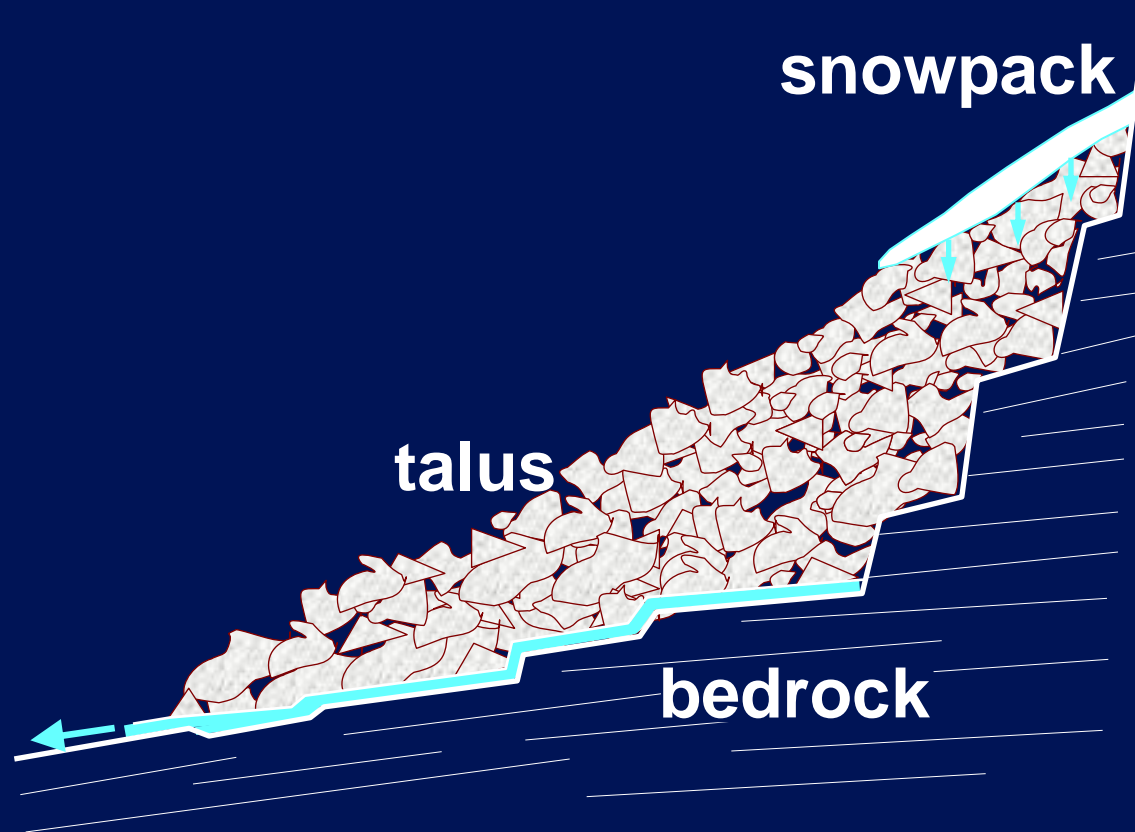


Muir et al. (2011, *Hydrol. Process.*, 25: 2954)

Conceptual Model of Talus Groundwater

Fast hydraulic response time (< 2-3 days).

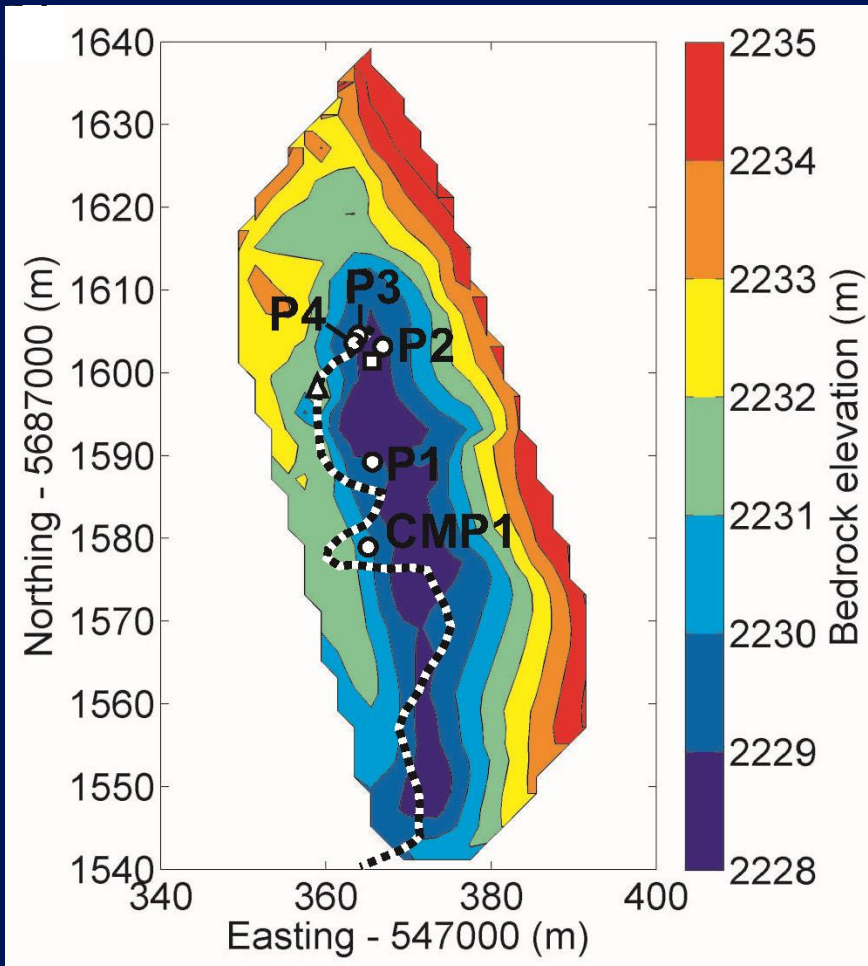
Flow through a thin (< 0.1 m) saturated layer.



Muir et al. (2011, *Hydrol. Process.*, 25: 2954)

Hydrogeological Response Unit: Alpine Meadow

Water table is controlled by fill-spill of bedrock basin.



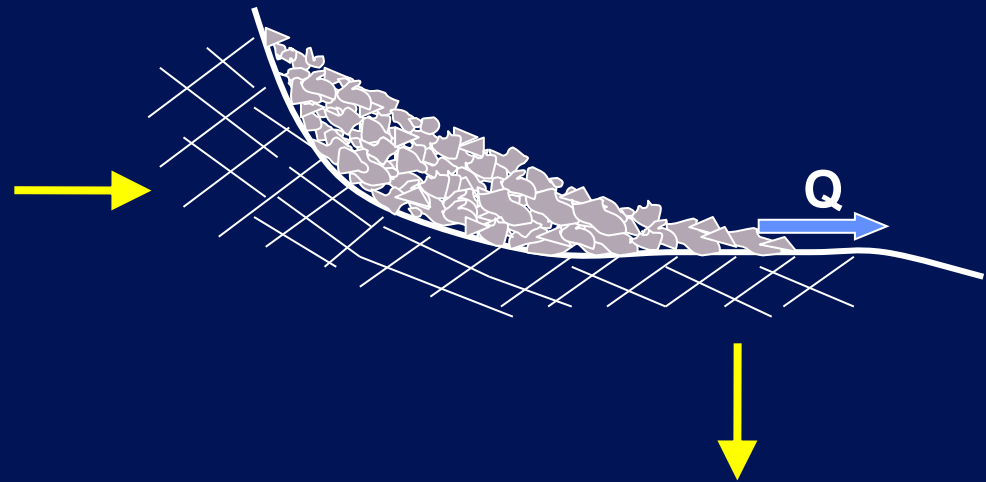
McClymont et al. (2010, *Hydrol. Earth System Sci.*, 14: 849)

From Process Observation to Hydrograph

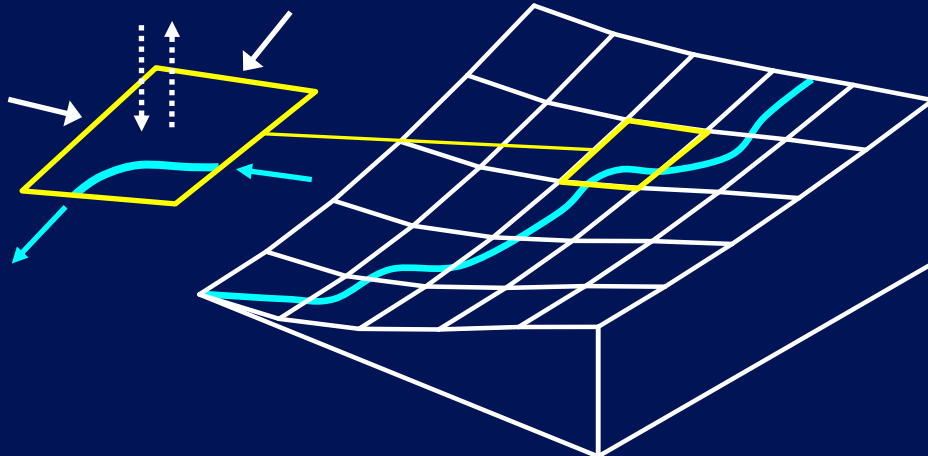
field observation



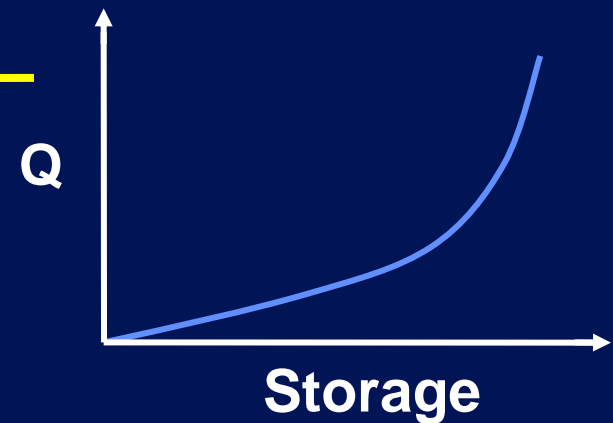
physically-based model



river-basin model

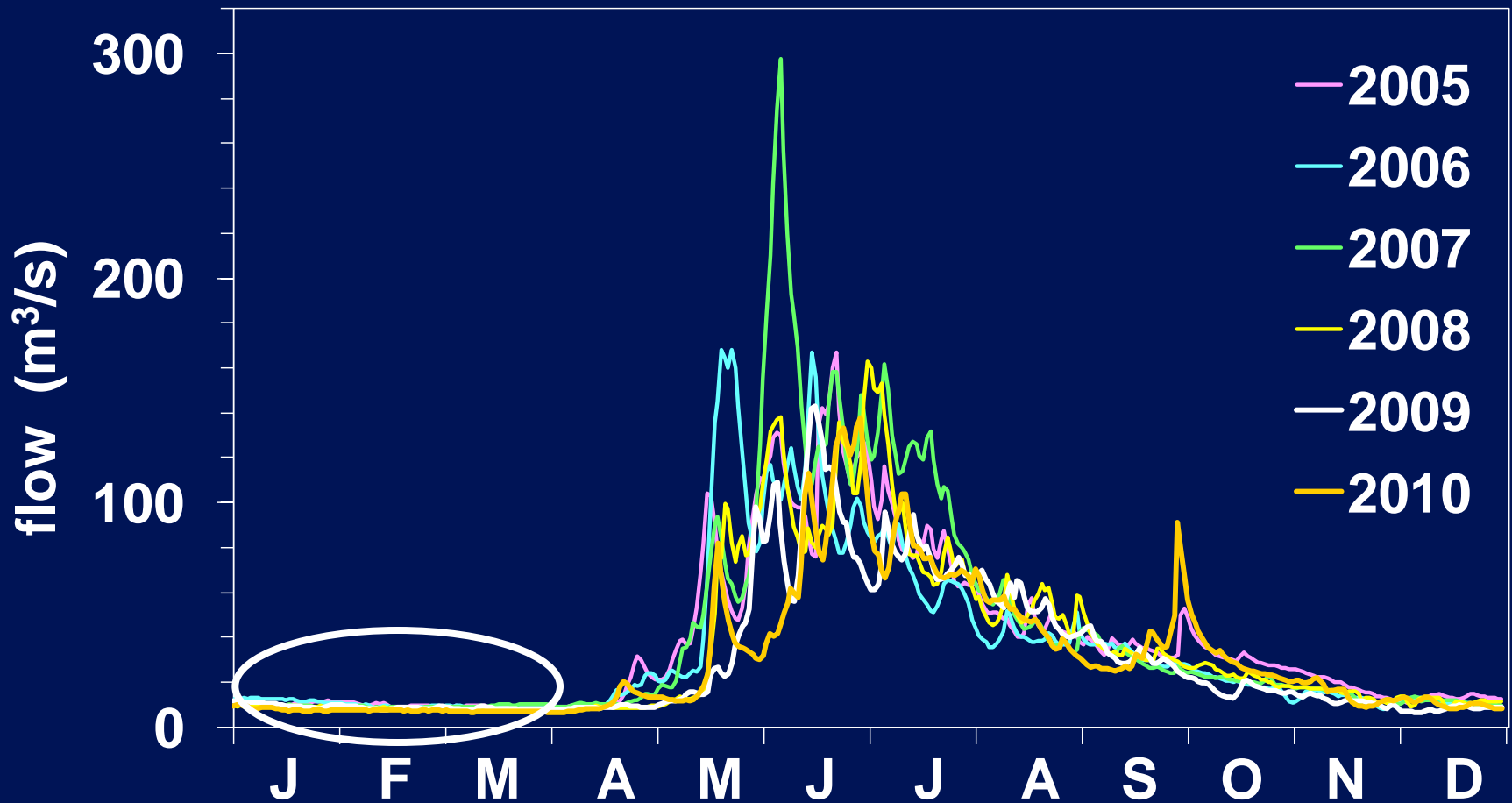


grid-scale function



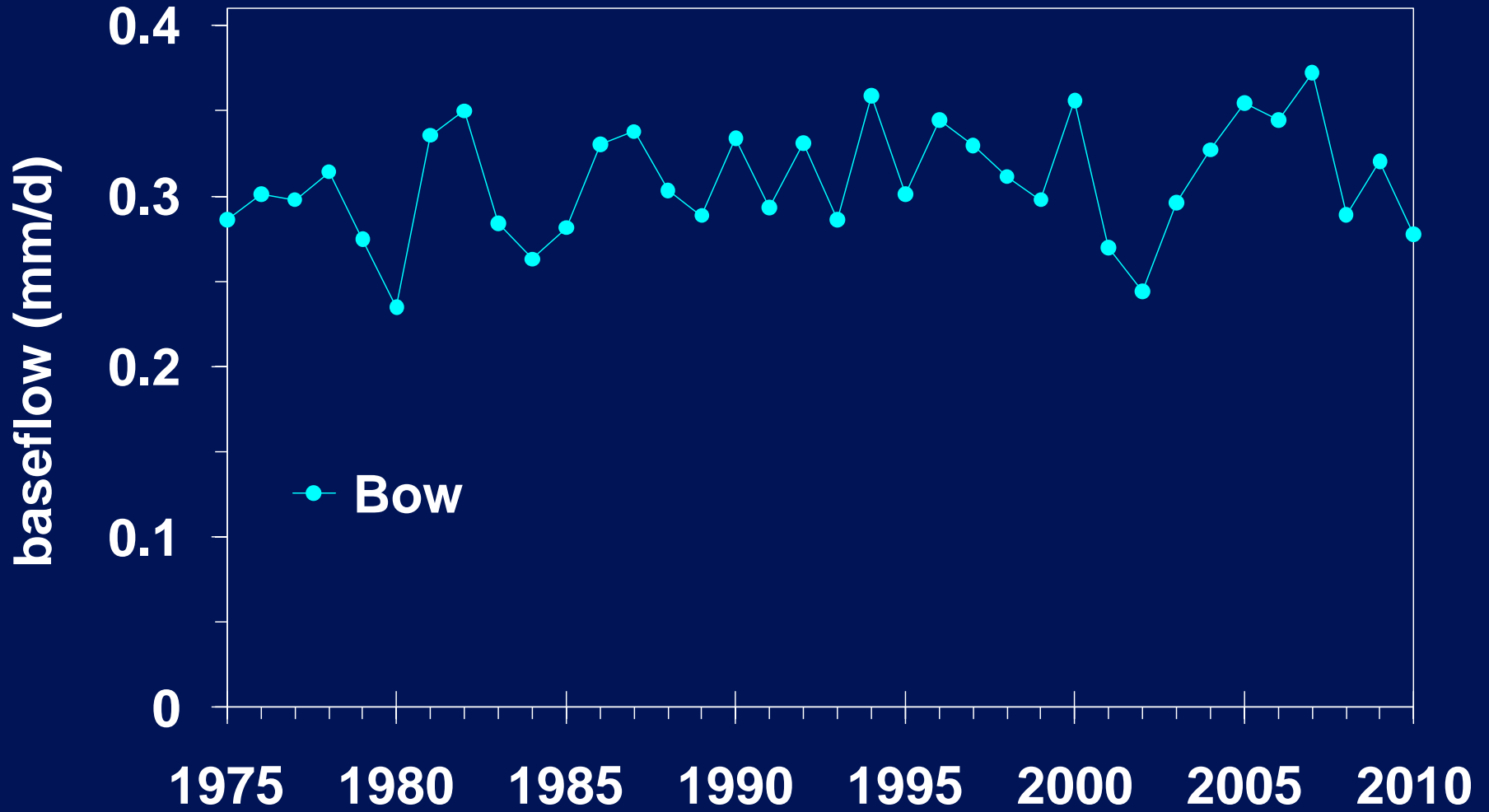
Bow River in the Canadian Rockies

Discharge at Banff (2200 km², unregulated)



Data: Water Survey of Canada

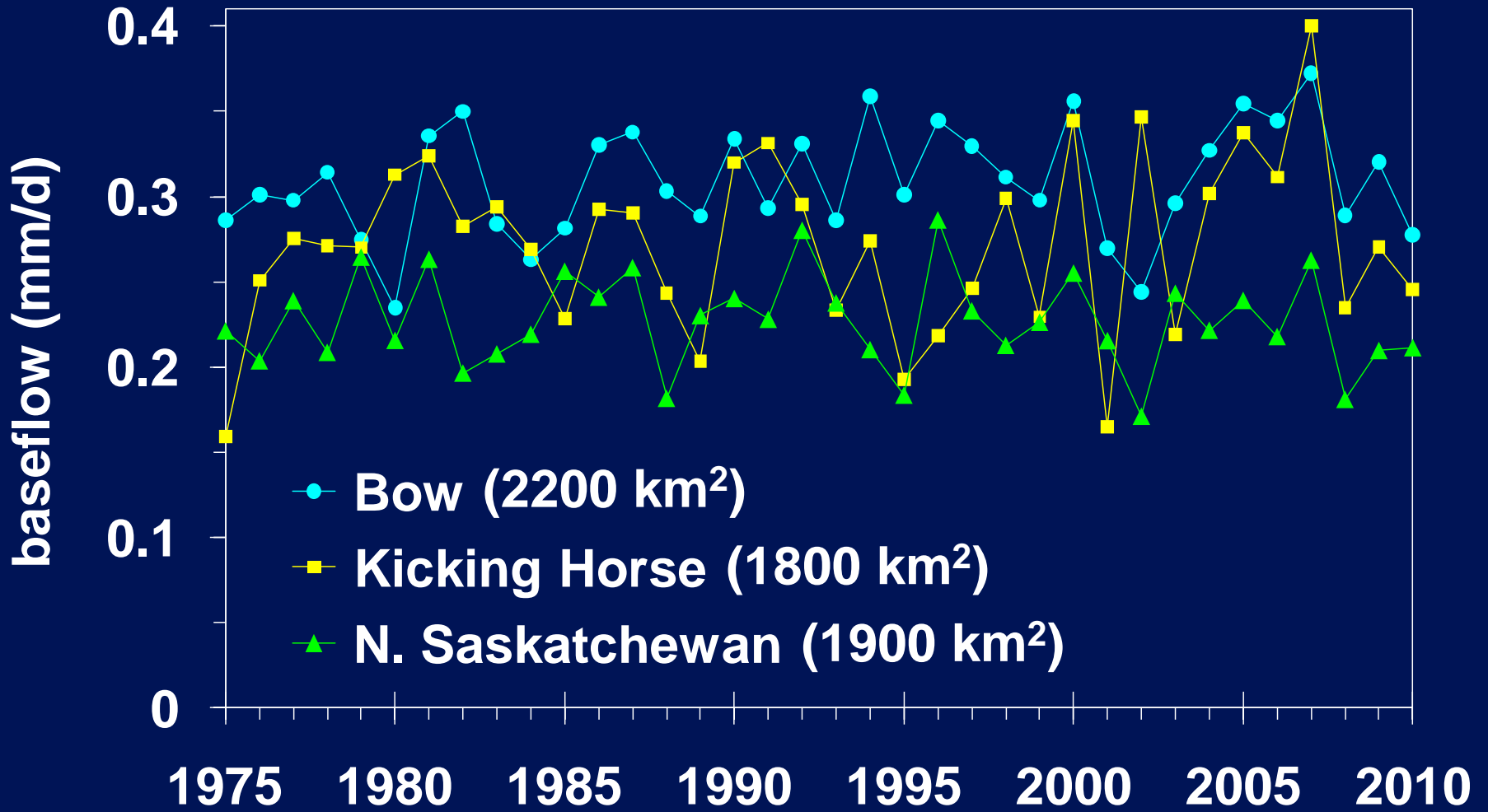
Average Winter Baseflow discharge / drainage area



Data: Water Survey of Canada

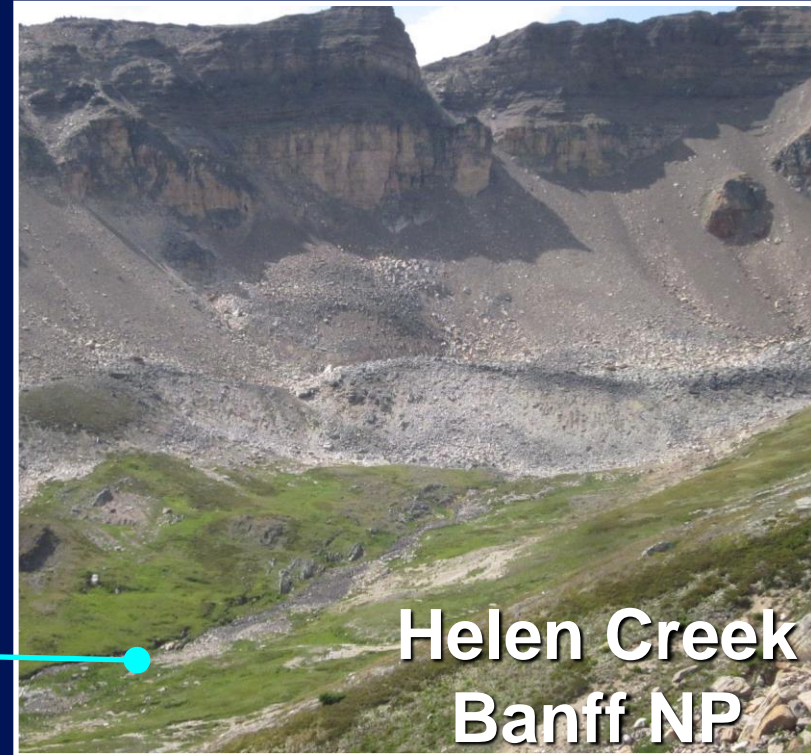
Average Winter Baseflow Runoff

discharge / drainage area



Data: Water Survey of Canada

Groundwater-Dependent Ecosystems



Helen Creek
Banff NP



www.nps.gov



Key Points

- 1. Groundwater stores snowmelt, glacier-melt, and rain water; and releases them slowly.**
- 2. Storage capacity is smaller than snowmelt volume, but significant for baseflow provision.**
- 3. Bedrock topography controls groundwater flow in alpine hydrogeological response units.**
- 4. Challenge remains in scaling up the field-based understanding to river basin models.**

Acknowledgements

Students, Technicians, Post-docs, Colleagues

Greg Langston, Danika Muir, James Roy, Alastair McClymont, Larry Bentley,

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